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Study of the overturning length scales at the Spanish planetary boundary layer

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Abstract

The focus of this paper is to analyze the behaviour of the maximum Thorpe displacement $(d_{\mathrm{T}})_{\max}$ and the Thorpe scale L_{T} at the atmospheric boundary layer (ABL), extending previous research with new data and improving our studies related to the novel use of the Thorpe method applied to ABL. The maximum Thorpe displacements varies between −900 and 950 m for the different field campaigns. The Thorpe scale L_{T} ranges between 0.2 and 680 m for the different data sets which cover different stratified mixing conditions (turbulence sher-driven and convective regions). We analyze the relation between $(d_{\mathrm{T}})_{\max}$ and the Thorpe scale L_{T} and we deduce that they verify a power law. We also deduce that there is a difference in exponents of the power laws for convective conditions and shear-driven conditions. This different power laws could identify overturns created under different mechanisms.

1 Introduction

Atmospheric boundary layer (or ABL) is almost always turbulent. In the absence of turbulence, atmospheric temperature profiles become increasingly monotonic, due to the smoothing effect of molecular diffusion. Turbulence gives rise to an effective eddy diffusivity and as well as other causes (as fluid instabilities or internal wave breaking) makes vertical overturns appear as inversions in measured temperature profiles. These overturns produce small-scale turbulent mixing which is of great relevance for many processes ranging from medium to a local scale. Unfortunately, measuring at small scales is very difficult. To overcome this disadvantage it is interesting to use theories and parameterizations which are based on larger scales. For example, the theories of turbulent stirring which often depend on hypotheses about the length scales of turbulent eddies. Vertical overturns, produced by turbulence in density stratified fluids as lakes or the ABL, can often be quantified by the Thorpe displacements d_{T} and the Thorpe scale L_{T} (Thorpe, 1977).

Next we present the atmospheric data used for the analysis. In Sect. 3 we present the Thorpe method and the definitions of the scale descriptors used. In Sect. 4, the results of Thorpe displacements, the maximum Thorpe displacement and the Thorpe scale L_T at ABL are presented and discussed.

2 Atmospheric data sets and meteorological instrumentation

The results presented in this paper are based on three ABL field campaings made at Spain and called Almaraz94-95, Sables98 and Sables2006. ABL data from 98 zeppelin-shaped tethered balloon soundings ranging from 150 to 1000 m were carried out in Almaraz94-95 field campaign made in Almaraz (Cáceres, Spain). The ABL profiles were obtained from 25 to 29 September 1995 in the time intervals 06:00–12:00 and 15:00–00:00 GMT. And from 5 to 10 June 1994 in the time intervals 05:00–12:00 and 17:00–00:00 GMT. Almaraz94-95 experiment collects data over a whole day and, therefore, covers different stratified conditions and mixing conditions – from shear-driven turbulence to convective regions. For more details see López et al. (2008). Sables98 (Stable Atmospheric Boundary Layer Experiment in Spain) took place over the northern Spanish plateau in the period 10–28 September 1998. The campaign site was the CIBA (Research Centre for the Lower Atmosphere). Two meteorological masts (10 and 100 m) were available at CIBA with high precision meteorological instruments (Cuxart et al., 2000). Additionally, a triangular array of cup anemometers was installed for the purpose of detecting wave events and a tethered balloon was operated at nighttime. A detailed description can be consulted in Cuxart et al. (2000). Sables98 field campaign only collects data over the night and, therefore, under neutral to stable conditions. Sables2006 field campaign took place from 19 June to 5 July 2006 at the CIBA. As in Sables98, different instrumentation was available on a tower of 100 m, a surface triangular array of microbarometers was also deployed and a tethered balloon was used to get vertical profiles of the atmosphere up to 1000 m. As in Sables98, Sables2006 field campaign also collects data over the night. Therefore, Sables98 and Sables2006 experiments let us to analyze the behaviour of overturns under stable conditions while Almaraz94-95 under unstable conditions (and also stable ones). These three sets of data were selected for this analysis because they cover different mixing conditions (turbulence shear-driven and convective regions).

3 Thorpe method and overturn length scales

Thorpe devised an objective technique for evaluating a vertical length scale associated with overturns in a stratified flow (Thorpe, 1977; Itsweire, 1984; Gavrilov et al., 2005). Thorpe's technique consists of rearranging a density profile (which contains gravitationally unstable inversions) so that each fluid particle is statically stable. If the sample at depth z_n must be moved to depth z_m to generate the stable profile, the Thorpe displacement d_T is $z_m - z_n$ (Thorpe, 1977; López et al., 2008, 2015). The Thorpe displacement d_T is not necessarily the real space actually traveled by the fluid sample, is an estimate of the vertical distance from the given vertical profile to the statically stable one that each fluid particle has to move up- or downward to its position in the stable monotonic profile (Thorpe, 1977; Dillon, 1982). Over most of a typical profile, the local stratification will be stable and the Thorpe displacement zero. A turbulent event is, therefore, defined as a region of continuously nonzero d_T, i.e, overturns are defined as a profile section for which $\sum_i d_{\mathrm{T}i} = 0$ while $d_{\mathrm{T}i} \neq 0$ for most i (Dillon, 1982; Peters et al., 1995).

The maximum of the Thorpe displacements scale $(d_\mathrm{T})_\mathrm{max} = \max[d_\mathrm{T}(z)]$ represents the larger overturns which might have ocurred at earlier time when buoyancy effects were negligible (Thorpe, 1977; Dillon, 1982; Itsweire, 1984) and it could be considered as an appropiate measure of the overturning scale.

The Thorpe scale L_T is the root mean square (rms) of the Thorpe displacements $(L_\mathrm{T})_\mathrm{rms} = L_\mathrm{T} = \left\langle d_\mathrm{T}^2(z) \right\rangle^{1/2}$. Therefore, it is a statistical measure of the vertical size of overturning eddies (Thorpe, 1977; Dillon, 1982; Itsweire, 1984; Fer et al., 2004) and is

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proportional to the mean eddy size as long as the mean horizontal potential temperature gradient is much smaller than the vertical gradient. For our field ABL measurements, we can consider that the ABL is horizontally homogenous because the average horizontal temperature gradient is 4×10^{-4} $(\mathrm{K\,m^{-1}})$ and is smaller than the average vertical temperature gradient, which is 2×10^{-2} $(\mathrm{K\,m^{-1}})$ (López et al., 2015).

Because of the expensive nature of collecting data at microscale resolution, there is a great interest to use parameterizations for small-scale dynamics which are based on larger scales – as L_T or $(d_\mathrm{T})_\mathrm{max}$. Therefore, it is very important to analyze the relation between L_T and $(d_\mathrm{T})_\mathrm{max}$ for selecting the most aproppiate overturning scale.

4 Quantitative results

Our methodology is based on reordering 111 measured potential temperature profiles, which may contain inversions, to the corresponding stable monotonic profiles. Then, the vertical profiles of the displacement length scales $d_\mathrm{T}(z)$ or Thorpe displacements profiles can be calculated by using a bubble sort algorithm with ordering beginning at the shallowest depth (Thorpe, 1977; Dillon, 1982; Itsweire, 1984; López et al., 2008, 2015). This simple sorting algorithm works by repeatedly stepping through the data list to be sorted, comparing each pair of adjacent items and swapping them if they are in the wrong order (López et al., 2015).

4.1 Thorpe displacement profiles at ABL

Usually, the signature that might be expected for a large overturning eddy is: sharp upper and lower boundaries with intense mixing inside – displacement fluctuations of a size comparable to the size of the disturbance itself are found in the interior. While common in surface layers strongly forced by the wind, these large features are not always found as in our ABL case (López et al., 2008, 2015). For our ABL studies, Thorpe displacements observed at profiles could be qualitative classified in two groups.

First, the Thorpe displacements under neutral and stable stratification conditions are always zero except in a region with isolated *Z* patterns which would correspond to discrete patches. These isolated overturns are very few well-defined sharp overturns which appear under stability conditions (at sunset, night and sunrise profiles). We really find other features that are smaller, some having an eddylike shape similar to the larger disturbances, some a random mix of small scale fluctuations without sharp boundaries. These are the second group or non-zero Thorpe displacement regions with indistinct and distributed features which appear under convective and/or neutral conditions (at noon, afternoon and evening profiles). These Thorpe displacements are rarely zero for the whole profile. To verify this behaviour see López et al. (2008, 2015).

4.2 Time evolution of maximum Thorpe displacements and Thorpe scale

Figure 1 shows the time evolution of the maximum Thorpe displacement $(d_\mathrm{T})_\mathrm{max}$ along a day for the three field campaigns. The scale $(d_\mathrm{T})_\mathrm{max}$ is very small (approximately zero) under stable conditions from 23:00 to 06:00 GMT (between sunset and sunrise) for all the experiments. From 19:00 GMT, it is observed that scale $(d_\mathrm{T})_\mathrm{max}$ decreases. The greatest values of $(d_\mathrm{T})_\mathrm{max}$ appears under convective conditions from 09:00 to 19:00 GMT being positive and negative. But the positive values of $(d_\mathrm{T})_\mathrm{max}$ are greater than the negative ones. The positive $(d_\mathrm{T})_\mathrm{max}$ has its greatest values about 950 m and the greatest negative $(d_\mathrm{T})_\mathrm{max}$ are about 600 m. These results means as follows. Thorpe displacements were defined as the difference between the final height and the initial height of the fluid particle, i.e. $d_\mathrm{T} = (z_m)_\mathrm{final} - (z_n)_\mathrm{initial}$. If $d_\mathrm{T} > 0((z_m)_\mathrm{final} > (z_n)_\mathrm{initial})$, the fluid particle has to go up to reach its stable position, and if $d_\mathrm{T} < 0((z_m)_\mathrm{final} < (z_n)_\mathrm{initial})$, it has to go down to reach its stable point. From Fig. 1 we can deduce that fluid particles go up and downwards with a high estimation of the vertical distance travelled under convective stratification conditions. Under stable stratification conditions – at night –, the fluid particles also move up and downwards but with small values for the vertical distance travelled. Hence, it is clear that the maximum Thorpe displacement is always greater under convective conditions than under stable ones, independently of its sign.

Therefore, the maximum Thorpe displacements is a parameter which could represent the dynamical behaviour of air particles and its relation with the stratification conditions. Finally, there is a gap in Fig. 1 due to non registered data between 13:00 and 14:00 GMT.

Figure 2 shows the time evolution of the Thorpe scale, L_T during a day for the three field campaigns. The Thorpe scale L_T has small values (and close to zero) under neutral and stable conditions (between sunset and sunrise) from 20:00 to 09:00 GMT. This scale reachs its greatest values under convective conditions from 09:00 to 19:00 GMT. There are two distinct behaviours with high ($L_\mathrm{T} > 100\,\mathrm{m}$) and low ($L_\mathrm{T} < 100\,\mathrm{m}$) Thorpe scales. In most of the turbulent patches, the Thorpe scale does not exceed several tens of meters and they appear under stable and neutral stratification conditions when the Thorpe displacements are also small and related to instantaneous density gradients. In contrast, under convective conditions, Thorpe scales are relatively large, they exceed hundreds of meters and they may be related to convective bursts. Hence, the Thorpe scale L_T is always greater under convective conditions than under stable ones and it is a parameter which also could represent the dynamical behaviour of air particles. As in Fig. 1, there is a gap in Fig. 2 due to non registered data between 13:00 and 14:00 GMT. Both scales, the Thorpe scale L_T and the maximum Thorpe displacement $(d_\mathrm{T})_\mathrm{max}$, have small values (and close to zero) under neutral and stable conditions, and their greatest values appear under convective conditions. Therefore, it is reasonable to think which of the two scales could represent better the dynamical behaviour of turbulent overturns although both scales are alternative length scales that coul be used for characterizing turbulent overturns.

Moreover, it is neccessary to choose an appropiate overturning scale to characterize instabilities leading to turbulent mixing, the turbulent overturning motions themselves and to look for a relation with the Ozmidov scale at ABL data (Dillon, 1982; Lorke and Wüest, 2002; Fer et al., 2004). We could choose the Thorpe scale rather than the maximum Thorpe displacement because we only sample vertically while the turbulence is three dimensional and, therefore, the Thorpe scale is more likely to be a statistically stable representation of the entire feature (Dillon, 1982). But the maximum of the Thorpe displacements is also considered as an appropiate measure of the overturning scale and it is always greater than L_T (better detectable by a limited resolution instrument). Different researchers have found a linear relationship between L_T and $(d_T)_{max}$ for profiles from the equatorial undercurrent (Moum, 1996; Peters et al., 1995) and a high linear correlation computed from the Banyoles99 field data where the ratio $(d_T)_{max}/L_T$ is approximately equal to 3 (Piera Fernández, 2004). It must exist a correlation between L_T and $(d_T)_{max}$ because when computing the rms of a set of Thorpe displacements with high kurtosis distributions, the final result depends on the largest values (Piera Fernández, 2004; Stansfield et al., 2001). A similar linear correlation between L_T and $(d_T)_{max}$ has been found by other researchers: a ratio $(d_T)_{max}/L_T \approx 3.3$ is obtained in the oceanic thermocline (Moum, 1996), a ratio $(d_T)_{max}/L_T \approx 2.4$ is obtained from laboratory experiments (Itsweire et al., 1993) and, finaly, the ratio $(d_T)_{max}/L_T$ is nearly 3 in numerical simulations (Smyth and Moum, 2000). But for microstructure profiles from strongly stratified lakes, a power law – as $(d_T)_{max} \sim (L_T)^{0.85}$ – is found (Lorke and Wüest, 2002). This relation also holds for profiles from other lakes under very different conditions of mixing and stratification with a strong correlation that holds over four orders of magnitude (Lorke and Wüest, 2002).

Hence, we analyze the relation between L_T and $(d_T)_{max}$ scales for our ABL data. Figure 3 shows the maximum Thorpe displacement versus the Thorpe scale from the three field campaigns. The relationship $|(d_T)_{max}| = L_T$ is indicated. We observe that the linear relationship $|(d_T)_{max}| = L_T$ proposes by other authors (Moum, 1996; Peters et al., 1995; Piera Fernández, 2004; Itsweire et al., 1993; Smyth and Moum, 2000) does not verify for our ABL data.

Therefore, we could think that the nearly constant ratio $(d_T)_{max}/L_T$ obtained in a wide range of field and laboratory experiments, does not verify in our ABL data. And, hence, the shape of Thorpe displacements distribution could change at ABL. We also observe a strong correlation which holds over three orders of magnitude as in other researches from profiles in lakes (Lorke and Wüest, 2002). It is the first time that such a relation between this two overturning length scales is found for ABL data.

As other authors, we could state that this string correlation indicates that the Thorpe scale is the determined by the overturns near to the maximum Thorpe displacement. We find the following power law:

$$|(d_\mathrm{T})_\mathrm{max}| \sim (L_\mathrm{T})^{1.14}, \tag{1}$$

which is similar to the one deduced by Lorke and Wüest (2002) from profiles in strongly stratified lakes. We realize a linear simple regression analysis. Of particular interest is the P value associated to the analysis of variance, which tests the statistical significance of the fitted model. For our case the P value is less than 0.05 (operating at the 95 % confidence level) which indicates that a significant linear relationship exists between $|(d_\mathrm{T})_\mathrm{max}|$ and L_T. Moreover, the R-squared coefficient is 96.95 which represents that the linear regression accounts for about 97 % of the variability in the maximum Thorpe displacement $|(d_\mathrm{T})_\mathrm{max}|$ as a function of the Thorpe scale, L_T.

This relation between the maximum Thorpe displacement and the Thorpe scale by a power law has been deduced for the overall data (not separating the data from the three field campaigns). But we have used three different experiments data set with different mixing conditions. SABLES98 and SABLES2006 experiments have been realized at night (turbulence by shear-driven) and ALMARAZ94-95 during a day cycle and, therefore, convective regions have not been excluded. Hence, we consider to analyze if this power law is different from night to day. The objective is to study if it s possible to distinguish between the shear-driven overturns and the convective ones. First, we separate the data from the three experiments in two set: data obtained overnight (from Sables98, Sables2006 and Almaraz94-95 field campaigns) or night data set, and data which have been obtained during the day (only from Almaraz experiment) or day data set. Then we realize a linear simple regression analysis with an adjustment by least squares for the two data sets. And, finally, we realize a comparison of regression lines procedure which is designed to compare the regression lines relating $|(d_\mathrm{T})_\mathrm{max}|$ and L_T at the two levels of our categorical factor (daytime and nighttime).

Figure 4 represents the maximum Thorpe displacement vs. the Thorpe scale only for the daytime data set (from 07:00 to 19:00 GMT). As before, we observe that the linear relationship $|(d_\mathrm{T})_\mathrm{max}| = L_\mathrm{T}$ does not verify for our ABL daytime data. We also observe a strong correlation which holds over three orders of magnitude as it was deduced for the whole data set and in other researches from profiles in lakes (Lorke and Wüest, 2002).

We realize the linear simple regression analysis. The P value associated to the analysis of variance is less than 0.05 (operating at the 95 % confidence level) which indicates that a significant linear relationship exists between $|(d_\mathrm{T})_\mathrm{max}|$ and L_T as before. The R-squared coefficient represents the percentage of the variability in $|(d_\mathrm{T})_\mathrm{max}|$ which has been explained by the fitted linear regression model and is about 97 %.

Figure 5 represents the maximum Thorpe displacement vs. the Thorpe scale only for the nocturnal data set (from 20:00 to 06:00 GMT). The relationship $|(d_\mathrm{T})_\mathrm{max}| = L_\mathrm{T}$ is indicated. As before, we observe that the linear relationship $|(d_\mathrm{T})_\mathrm{max}| = L_\mathrm{T}$ does not verify for these nocturnal ABL data set. We also observe a strong correlation which holds over three orders of magnitude as in the results mentioned before (see Figs. 3 and 4).

Finally, we realize the linear simple regression analysis. The P value associated to the analysis of variance is less than 0.05 (operating at the 95 % confidence level) which indicates that a significant linear relationship exists between $|(d_\mathrm{T})_\mathrm{max}|$ and L_T as before. Moreover, the R-squared coefficient is 95.89 which represents that the linear regression accounts for about 96 % of the variability in the maximum Thorpe displacement $|(d_\mathrm{T})_\mathrm{max}|$.

Therefore, we have deduced that the relation between the maximum Thorpe desplacement $|(d_\mathrm{T})_\mathrm{max}|$ and the Thorpe scale L_T by a power law is different from day to night. For the nighttime data set the power law is:

$$|(d_\mathrm{T})_\mathrm{max}| \sim (L_\mathrm{T})^{1.17}. \quad (2)$$

And for the daytime data set the relation is the following:

$$|(d_\mathrm{T})_\mathrm{max}| \sim (L_\mathrm{T})^{1.12}. \quad (3)$$

We observe that the kind of relation is the same (a power law) but the exponents are different. So we question if these coefficients are statistically different and if there is or not a different behaviour of the overturn length scales between day and night.

These exponents are the slopes of the regression lines fitted to daytime and nighttime data sets (see Figs. 4 and 5). To know if they are statistically different we need to realize a comparison of regression lines. This procedure is a test to determine whether there are significant differences between the intercepts and the slopes at the different levels of our factor (day and night). This test fits two different regression lines to the nighttime data and daytime data sets and realizes two analysis of variance (one for the linear model and secondly for). For the first analysis, the *P* value is less than 0.05, if we operate at the 5 % significance level, and indicates that a significant relationship of the linear form exists between $|(d_\mathrm{T})_\mathrm{max}|$ and L_T for daytime and nighttime data sets (*t*-statistic tests have also been made which *P* values are less than 0.05, 95 % confidence level, and indicate that the model coefficients are significantly different from 0). For the second analysis of variance is performed to determine whether there are significant differences between the slopes of the daytime and nighttime lines are significantly different. We find that the *F* test for slopes tests the null hypotheses slopes of the lines are all equal. Operating at the 1 % significance level, we find a *P* value (for slopes) is less than 0.01, and, therefore, there are significant differences between the slopes of the daytime and nighttime lines (we get the same result for the intercepts).

Finally, we deduce that the two power relation between the maximum Thorpe displacement $|(d_\mathrm{T})_\mathrm{max}|$ and the Thorpe scale L_T for nighttime data (Eq. 2) and daytime data (Eq. 3) are significant different with a 99 % confidence level. Therefore, we could

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classified overturns between day and night ones, i.e., we could distinguish between convective and shear-driven mechanism originating overturns.

As mentioned before, although both scales ($(d_\mathrm{T})_\mathrm{max}$ and L_T) are alternative length scales to characterize turbulent overturns, it is reasonable to choose which of the two scales could represent better overturns. If there is a high linear correlation between the maximum Thorpe displacement $(d_\mathrm{T})_\mathrm{max}$ and the Thorpe scale L_T, this could be considered a better descriptor of the overturn properties although it depends mainly of the values of $(d_\mathrm{T})_\mathrm{max}$ and the relative errors from both scales are approximately equal (Piera, 2004). But we have just deduced that the relation between the maximum Thorpe displacement $(d_\mathrm{T})_\mathrm{max}$ and the Thorpe scale L_T is not linear at our ABL research, unless a power law as other authors (Lorke and Wüest, 2002). Consequently, there would not be a constant ratio $|(d_\mathrm{T})_\mathrm{max}|/L_\mathrm{T}$ which could suggest that the shape of Thorpe displacements distribution could change. Therefore, it is necessary to study the probability density functions (pdf) of the Thorpe displacements to understand better the relation between $(d_\mathrm{T})_\mathrm{max}$ and L_T. Moreover, the Thorpe scale is mainly determined by larger overturns which are not very frequent (Stansfield et al., 2001) and it would be very useful to determine it based on the probability density function of the Thorpe displacements. This pdf study would allow us to decide which of the two overturn length scales is a more representative measure of turbulent overturns.

5 Conclusions

The paper presents results related to the time evolutions of the ABL turbulent parameters L_T and $(d_\mathrm{T})_\mathrm{max}$ during a day with different levels of stability. Secondly, the paper adds insight to the problem of the relationship between these two overturning length scales at ABL.

The Thorpe scale L_T and the maximum Thorpe displacement $(d_\mathrm{T})_\mathrm{max}$, have small values under neutral and stable conditions, and their greatest values appear under convective conditions. The values of the Thorpe scale ranges in (1, 660) m that are greater than effective values in the stratosphere which are $L_T \sim 1$–1.1 m (Gavrilov, 2013), values in mixing surface layers and seasonal thermoclines which are $L_T \sim 0.03$–1.90 m (Dillon, 1982), values in vertical mixing process induced by internal tides which are $L_T \sim 0.2$–4.2 m (Kitade et al., 2003) or values in dense overflow which are $L_T \sim 1$–17 m (Fer et al., 2004). The greater values appear under convective conditions which could generate overturns of larger scale. Under shear-driven conditions, our Thorpe scales are smaller than convective ones, ranging in (1100) m, but they are also greater if we compare them with the scales of other authors. Therefore, we deduce that there would be a relation between the ABL processes which generate mixing and the overturn size and behaviour (for example, the terrain shape interacts with the ability of the ABL to produce local mixing very near the ground and this could be affect to overturns). This theme will need further field work where different conditions are met (combination of the boundary condition effects and of stability combining the 3-D and 2-D characteristics of scale to scale direct and inverse cascades, intermittency of the forcing and scale to scale stratified turbulence cascade, Vindel et al., 2008; Yagüe et al., 2006).

Equations (1) to (3) show that the relationship between the Thorpe scale L_T and the maximum Thorpe displacement $(d_T)_{\max}$ is a power law which has been statiscally demostrated. We must therefore conclude that the linear relation proposes by other authors (Moum, 1996; Peters et al., 1995; Piera Fernández, 2004; Itsweire et al., 1993; Smyth and Moum, 2000) is not adequate for our spanish ABL data. Research will continue on this interesting question which is related to the selection a length scale for characterizing turbulent overturns. This last problem would be better analyzed if we study the probability density function (pdf) of overturning length scales. The objective is to decide if L_T is or not statistically a more appropiate length scale than $(d_T)_{\max}$. Moreover, it is interesting to verify the assumption that the Thorpe scales have a universal probability density function which could be used to verify how accurately the Thorpe scales were computed and also to determine if $(d_T)_{\max}$ is statistically better than L_T as overturning length scale. It is very likely that the pdf parameters depend on the governing background conditions generating Thorpe displacements, which are

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different in the boundary layers from those in the interior layers with intermittent mixing, or in convective conditions from shear-driven conditions. We also would like to verify if the density probability function is decaying exponentially for increasing displacement length with a separate cut-off before $(d_{\mathrm{T}})_{\mathrm{max}}$.

In the future, we will go on studying the power relationship between the maximum Thorpe displacement and the Thorpe scale corresponding to ABL data to verify the power law deduced at this paper. For this purpose, we will use new set of spanish ABL data from new field campaigns. We will analyze the probability density function of overturning length scales to clarify better the relation between $(d_{\mathrm{T}})_{\mathrm{max}}$ and L_{T} and as a tool to choose the more appropiate turbulent patch length scale. Moreover, we would like to study the following hypothesis if the Thorpe scale is greater than the integral scale there would be a local convective process and if it is not, there would be stratification.

Finally, there is another subject which is important to mention. At future researches, we need to study better the overturn identification as Piera et al. (2002). They propose a new method based on wavelet denoising and the analysis of Thorpe displacements profiles for turbulent patch identification. Although their method is for microstructure profiles (that is not our case), it reduces most of the noise present in the measured profiles (increasing the resolution of the overturn identification) and it is very efficient even at very low-density gradients for turbulent patch identification. Another way to get overturn identification would be, for example, to use a 3 or 4 dimensional parameter space formed by $(L_{\mathrm{O}}, L_{\mathrm{T}}, L_{\mathrm{MO}})$ to locate mixing events and also to study the evolution of the processess.

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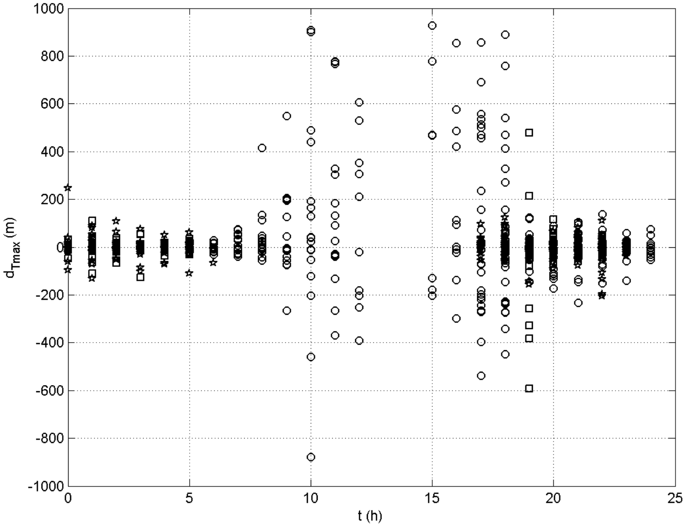

Figure 1. Time evolution of the maximum Thorpe displacements during a day cycle. The symbols are as follows: ∘ is for Almaraz94-95 data, ★ is for Sables98 data and ☐ is for Sables2006 data.

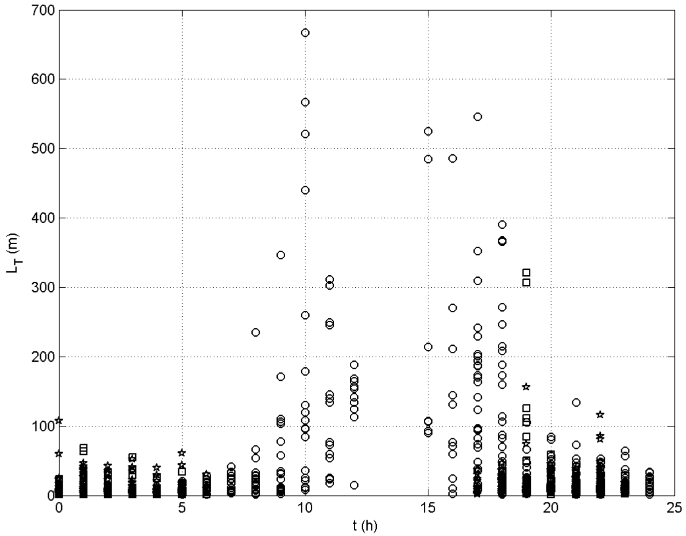

Figure 2. Time evolution of the Thorpe scale during a day cycle. The symbols are as follows: ∘ is for Almaraz94-95 data, ★ is for Sables98 data and ☐ is for Sables2006 data.

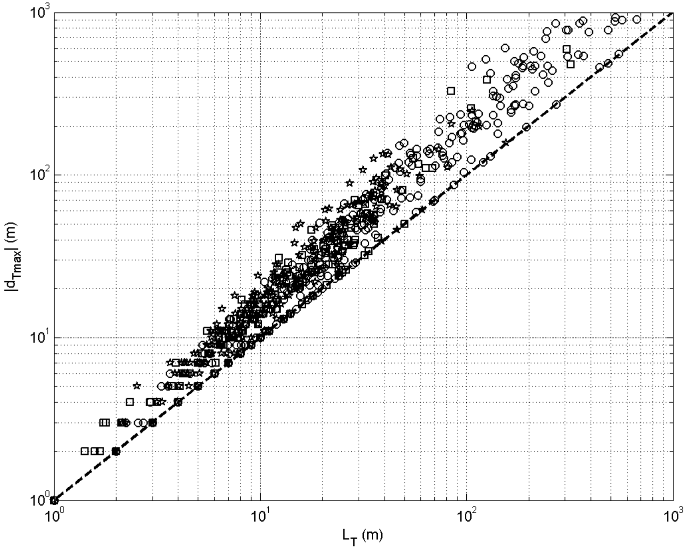

Figure 3. Absolut value of the maximum Thorpe displacement vs. Thorpe scale. The symbols are as follows: ∘ is for Almaraz94-95 data, ★ is for Sables98 data and ☐ is for Sables2006 data. The relationship $|(d_T)_{max}| = L_T$ is indicated by the discontinous line.

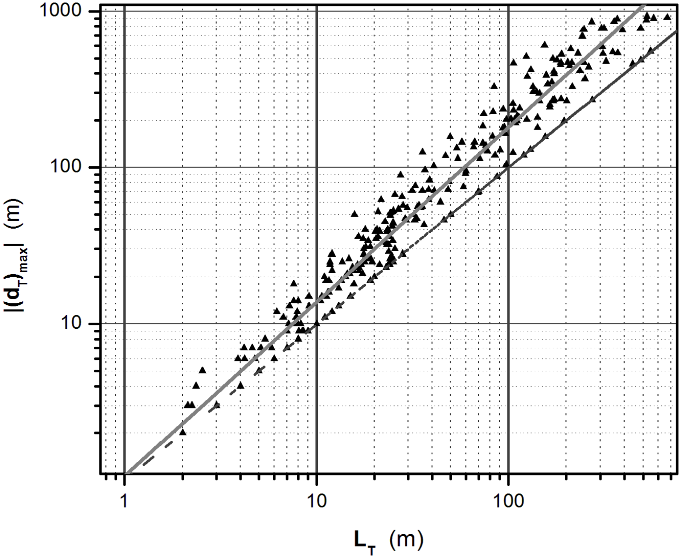

Figure 4. Absolut value of the maximum Thorpe displacement vs. Thorpe scale for the daytime data set (▲). The relationship $|(d_T)_{max}| = L_T$ is indicated by the discontinous darkgray line. The linear fit is indicated by the continuous light gray line.

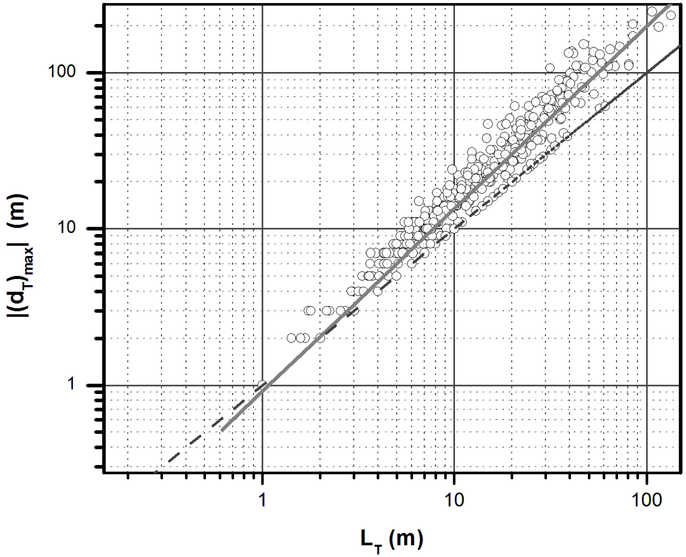

Figure 5. Absolut value of the maximum Thorpe displacement vs. Thorpe scale for the nighttime data set (∘). The relationship $|(d_\mathrm{T})_\mathrm{max}| = L_\mathrm{T}$ is indicated by the discontinous darkgray line. The linear fit is indicated by the continuous light gray line.