



Recent trends in daily temperature extremes over southern Montenegro (1951–2010)

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Abstract. Montenegro so far has been poorly investigated in terms of climate extremes. The aim of this paper was to analyse the extreme ETCCDI (Expert Team on Climate Change Detection and Indices) temperature indices in the Mediterranean region of Montenegro for the period of 1951–2010. Four stations in the coastal area of Montenegro have been analysed: Herceg Novi, Ulcinj, Budva and Bar. Two periods (before 1980 and after 1980) were separately investigated in this study due to a well-known climate shift that occurred in the late 1970s. Seven indices of temperature extremes have been chosen. The trend was analysed using a Mann–Kendall non-parametric test, while the slope was estimated using Sen’s slope estimator. A negative trend has been calculated for cold nights and cold days at almost all stations. The most significant positive trends were obtained for warm conditions. The two separately investigated periods have shown contrasting temperature trends.

1 Introduction

There has been considerable interest in the study of temperature and precipitation extremes at different scales (Zolina, 2008; Durao et al., 2010; Ceasar et al., 2011) due to their great impact on agriculture, water resources, industry, environment, and society. Since temperature records show an increase in the global mean temperature between 0.4 and 0.8 °C in the last 100 yr (IPCC, 2007), studies of extreme temperature events have raised the interest of the scientific community. Many authors have made an effort to

examine extreme temperature trends in the Mediterranean region (Klein Tank and Können, 2003; Kostopoulou and Jones, 2005; Della-Marta et al., 2007; Kuglitsch et al., 2010; Efthymiadis et al., 2011). Most of their results suggest an increased frequency and duration of warm events. Warmer conditions over the Mediterranean may impact the environment and increase the burden on societies that are currently under economic turmoil. However, due to a non-uniform magnitude of extreme temperature changes identified over the entire Mediterranean basin, many studies focused on particular countries. In Spain, numerous studies (Brunet et al., 2007; Rodriguez-Puebla et al., 2010) assessed trends in temperature extremes, and Brunet et al. (2007) noted larger changes in high temperature extremes than in low temperature extremes over the 20th century. While analysing temperature extreme indices in northeast Spain, El Kenawy et al. (2011) also observed more prevalent changes in hot extremes than in cold extremes. Toreti and Desiato (2008) analysed changes in temperature extremes over Italy and showed a cooling trend until the end of the 1970s followed by a warming trend over the last 25 yr. The most significant results they obtained were related to an increased number of summer days and tropical nights.

The regions of Serbia and Montenegro were not always covered within European studies on temperature extremes. Recent analysis of trends in temperature indices over Serbia and its relationship to large-scale circulations conducted by Unkašević and Tošić (2013) suggested that the Serbian climate generally became warmer over the last 61 yr. They showed the most significant temperature trends for the

Table 1. List of the stations.

Station name	WMO	Period	Latitude (N)	Longitude (E)	Elevation (m)
Budva	13458	1951–2010	42°17′	18°50′	2.0
Herceg Novi	13455	1951–2010	42°27′	18°31′	10.0
Bar	13461	1951–2010	42°06′	19°05′	5.7
Ulcinj	13464	1951–2010	41°55′	19°17′	3.6

summer season. However, the neighbouring Mediterranean region of Montenegro has been poorly studied with regards to temperature extremes. This highlights the interest for our study, as it is the first one over the investigated area and may contribute to a better understanding of temperature changes in Montenegro.

The structure of the paper is as follows: Sect. 2 provides a description of the study area; Sect. 3 describes station data, indices and methods; Sect. 4 includes a discussion of the results and conclusions are summarized in Sect. 5.

2 Study area

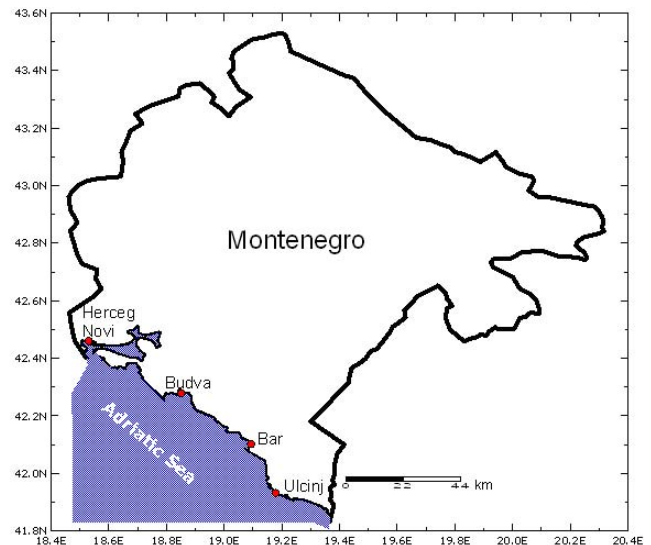
The study area comprises the Mediterranean region of Montenegro, which covers an area of 1500 km². It is surrounded by Orijen Mountain (1895 m), Lovćen Mountain (1749 m), Sutorman Mountain (1180 m) and Rumija Mountain (1595 m). Montenegro has a varied topography with a narrow Adriatic coastline stretching along the southwestern part, a Karst region in the central part and high Dinaric Mountains (above 2000 m) in the northern parts. The coastal region is characterized by a typical Mediterranean climate while a moderate continental climate is present in the northern parts of the country. The average annual rainfall is 2120 mm and the totals in the country generally decrease from the southwest towards the northeast. The mountains in the southwest, receiving above 4000 mm of precipitation, are the wettest part of the Mediterranean region (Ducić et al., 2012).

3 Data and methodology

3.1 Data

In this paper, a daily temperature time series was used from meteorological stations Ulcinj, Bar, Herceg Novi and Budva (Table 1, Fig. 1) covering the period of 1951–2010. Data was obtained from the Hydro-meteorological Service of Montenegro.

The data were quality controlled for all stations following standards recommended by WMO (2004). Metadata was analysed regarding relocation and maloperation. The locations of stations did not change throughout the investigated period and measurements were performed without a break using the same type of instruments. A time series homo-

**Fig. 1.** Locations of the stations.

geneity is a basic requirement when assessing changes in temperature extremes. Unfortunately, an advanced correction method for a daily temperature time series is not developed yet (Toreti et al., 2010). In this paper, data is carefully evaluated by applying a Multiple Analysis of Series for Homogenization (MASH v3.02) method. The MASH method was developed in the Hungarian Meteorological Service (Szentimrey, 1994, 1999, 2003); it is a relative homogeneity test that does not presume that the reference series are homogeneous (Costa and Soares, 2009). In this study a version of MASHv3.02 extended for homogenization of daily temperature data has been applied. Portions of the time series with inhomogeneities were excluded from the analysis.

An Expert Team on Climate Change Detection and Indices (ETCCDI), which is supported by the World Meteorological Organization (WMO) Commission for Climatology, the Joint Commission for Oceanography and Marine Meteorology (JCOMM) and the Research Program on Climate Variability and Predictability (CLIVAR) developed the list of temperature and precipitation indices used in this study.

In this paper, we selected a set of seven indices of temperature extremes (Table 2). Four indices are associated with warm temperature (SU, TR, Tn90p, Tx90p) and three indices with cold temperature (FD, Tx10p, Tn10p). All indices are defined in terms of the numbers of days that exceed either absolute or percentile thresholds.

Percentile-based indices allow spatial comparisons, as they sample the same part of the probability distribution of temperature at each location. Day-count indices based on absolute thresholds are less suitable for spatial comparisons of extremes. The reason is that, over large areas, day-count indices based on absolute thresholds may sample very different parts of the temperature distributions (Peterson et al., 2001.; WMO, 2009; Vincent and Mekis, 2006).

Table 2. Definitions and abbreviations of the ETCCDI temperature indices used in this study.

Index	Definition	Units
FD	Total number of frost days (days with absolute Tn (daily minimum temperature) < 0°C)	Days
SU	Number of summer days: Annual count of days when Tx (daily maximum temperature) > 25°C	Days
TR	Number of tropical nights: Annual count of days when Tn > 20°C	Days
Tn10p	Days with Tn < 10th percentile of daily Tn of the base period	Days
Tx10p	Days with Tx < 10th percentile of daily Tx of the base period	Days
Tn90p	Days with Tn > 90th percentile of daily Tn of the base period	Days
Tx90p	Days with Tx > 90th percentile of daily Tx of the base period	Days

Table 3. Trends in temperature indices in the period 1951–2010 in days per decade (with statistical significance values: *** 0.1 % level, ** 1 % level, * 5 % level and +10 % level).

Index	Ulcinj	Bar	Budva	Herceg Novi	Region (1951–2010)	Region (1951–1980)	Region (1981–2010)
FD	−0.33	−1.25***	0.00	0.00	−0.51*	−0.36	0.33
Tx10p	−1.72*	−2.50**	−2.41**	−1.67*	−2.16*	0.50	−3.88*
Tn10p	0.00	−5.00***	−4.29***	−1.41	−2.71***	1.25	−2.20
SU	2.37**	4.38***	3.98***	1.25	2.93**	−3.25+	7.67**
TR	2.37	5.56***	7.54***	5.00**	5.44***	−4.86*	10.00***
Tx90p	7.98***	9.66***	8.59***	5.65***	7.99***	−2.75	24.29***
Tn90p	4.49**	7.86***	6.36***	6.58**	6.80***	−7.25**	20.68***

3.2 Methodology

Percentile thresholds are determined empirically from the observed data series in the standard period of 1961–1990. The percentiles are calculated from five-day windows centred on each calendar day to account for the mean annual cycle. A five-day window is chosen to yield a total sample size of 30 yr × 5 days = 150 data points for each calendar day, which results in a relatively smooth annual cycle of percentile thresholds. The procedure ensures that extreme temperature events can occur with equal probability throughout the year (WMO, 2009).

The statistical significance of the calculated trends of the indices was tested using a Mann–Kendall test. This method has been applied because it is more suitable for non-parametric distributions. The Mann–Kendall test is used for trend analysis in ETCCDI workshops (Zhang et al., 2005). Sen's (1968) slope estimator was used for estimating trends within the indices. This is a more robust approach for the estimation of trends in indices based on daily data.

4 Results and discussion

Seven temperature indices have been assessed for the coastal region of Montenegro. Stations have been analysed individually and then the trend for the whole region was calculated. All stations indicate statistically significant increases in the

Tn90p index over the period of 1951–2010 showing the most consistent pattern of trends. There has been a temperature increase in the whole region as well (Table 3).

A very strong increase is also noted in Tx90p, as all stations indicated significant increases. Percentile-based indices have shown an agreement in sign and significance of trend for stations individually and as a whole region when compared with global results. In coastal Montenegro, the frequency of warm nights and warm days increased while the frequency of cold nights and cold days decreased. Percentile indices are more robust over large areas as they account for local climate effects (Caesar et al., 2011). Cold nights (Tn10p) have shown a negative trend that is significant for Bar, Budva and the whole region in 1951–2010. Similar to this, small and negative trends are calculated for cold days (Tx10p) and are significant for all stations.

The number of frost days (FD) recorded decreased over all stations except for Budva and Herceg Novi. Most trends are not significant, with the exception of the Bar station and its corresponding region for the entire period. Frich et al. (2002) found uniform decreases in FD in the second half of the 20th century. Results for the coastal region of Montenegro are consistent with those reported by Frich et al. (2002). However, in contrast, Kostopoulou and Jones (2005) found an increase in the number of frost days over the area of Serbia and Montenegro at annual and seasonal scale for the period 1958–2000.

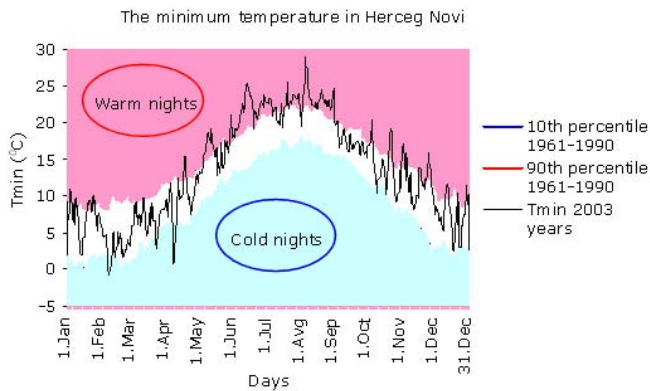


Fig. 2. Tn10p and Tn90p in Herceg Novi in 2003.

Summer days have shown strong statistically significant temperature increases for all stations with the exception of an insignificant trend for the Herceg Novi station. This reveals an increased number of days with a daily maximum temperature above 25 °C over the study area during the warm season.

Tropical nights have also shown strong statistically significant increases at all stations except for Ulcinj, indicating a tendency towards warmer conditions.

According to values of temperature indices, years 2003, 2007 and 2008 were particularly warm. For example, in 2003 in Herceg Novi there were 83 Tx90p, 133 SU, 117 Tn90p, 85 TR, only 17 Tx10p and 10 Tn10p with 1 FD. Figure 2 shows values of Tn10p and Tn90p in Herceg Novi in 2003. As can be seen in the figure, daily minimum temperatures were above the 10th percentile throughout the year. Atypical warmer conditions were detected in the capital city Podgorica in 2003, where daily maximum temperatures were 30 °C and above from 27 May until 1 September (Burić et al., 2011). This heat wave was detected over western and central Europe during the summer of 2003 and is an example of an exceptional recent extreme. That summer was the hottest in instrumental records (IPCC, 2007).

Trends calculated for the Mediterranean part of Montenegro have generally shown a tendency towards warmer conditions. Changes in warm indices are higher than changes in cold indices. Negative trends are calculated for cold nights, but the most significant trends are obtained in terms of warm conditions over the study area. The Bar station had the most pronounced shift with the strongest significance in general. This could be due to several reasons and include the fact that the process of urbanization has extended rapidly in the area over the past few decades, which may lead to urban islands of heat. Results for Ulcinj and Herceg Novi may not indicate the impact of urban heat island. These meteorological stations are located far from the town centre and the trend change for them is showing less significant results when compared to urban stations in Bar and Budva.

However, there has been a statistically significant increase in temperature in Montenegro over the second half of the

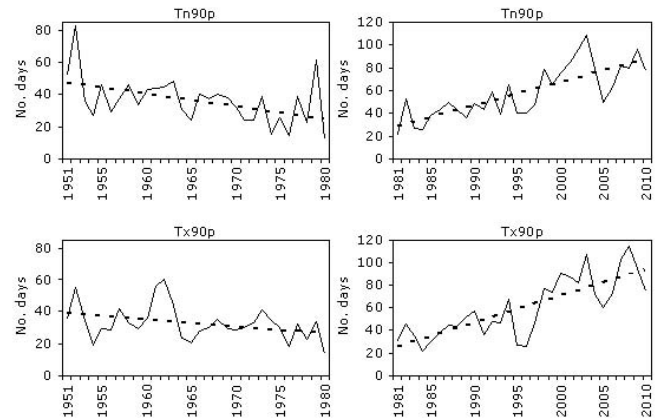


Fig. 3. Regional time series for warm nights and warm days (units: days) in (left) the 1951–1980 period and (right) the 1981–2010 period. Dashed line indicates Sen's slope estimator.

20th century (Burić, 2011). So far, there is no unique scientific opinion regarding the cause of such temperature increases. The greenhouse effect is not always responsible for temperature and precipitation variations, especially when internal oscillations may also be a contributing factor (Luković et al., 2010).

Due to a well-known great climatic shift that occurred in the late 1970s, two periods (before 1980 and after 1980) were separately investigated (Table 3) on a regional scale. In Europe, a cooling period followed by a warming period starting in the mid-1970s has been found (Balling et al., 1998; Klein Tank et al., 2002). Gurevich et al. (2011) found that the change in the temperatures in Israel occurred approximately between the mid-1980s and the mid-1990s. Toreti and Desiato (2008) have noted the increase in high temperature events in Italy from the late 1970s/early 1980s with slight differences between the gradients slopes of warming trends. For the period 1951–1980, all indices have shown decreases except for cold days and cold nights. Statistically significant decreases are calculated for summer days, tropical nights and warm nights. However, in the period of 1981–2010, results show significant increases in warm days and warm nights, including summer days and tropical nights. A significant decrease is calculated in cold days and is non-significant for cold nights. Regional indices calculated for the period of 1981–2010 are in accordance with those calculated for the whole period as well as with the global analysis of Alexander et al. (2006). Figure 3 shows opposite trends for warm nights and warm days in 1951–1980 and 1981–2010. For the Tn90p index, there is no agreement in the sign and the significance of the trend in 1951–1980 when compared with the trend in 1981–2010 and global results.

5 Conclusions

In this paper, seven indices have been analysed in order to assess temperature extremes in the Mediterranean part of Montenegro. The whole country, to date, has been poorly documented in terms of climate extremes.

Results presented in this paper are in agreement with the global analysis of Alexander et al. (2006). A negative trend has been calculated for cold nights and cold days at almost all stations. The most significant trends are obtained for warm conditions over the investigated area during the period of 1951–2010. This is also in accordance with other regional analyses (Klein Tank et al., 2006).

The trend found in this study can be associated with the positive phase of the North Atlantic Oscillation (NAO) since this pattern is recognized as the main mode of climate variability in the extratropical Northern Hemisphere (e.g. Hurrell, 1995; Sun et al., 2009; del Río et al., 2011).

However, a separately investigated period of 1951–1980 for the region has shown opposite tendencies and a contrasting trend to the period of 1951–2010 as well as 1981–2010. This result is possibly due to a well-known climate shift that occurred in the late 1970s where there was a change in the sign of trend for warm days and warm nights. These two separately investigated periods have shown contrasting temperature trends.

Future investigation should take into account other temperature indices as well as seasonal scales of the investigation.

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