
Impacts of the North Atlantic Oscillation on stream water chemistry in mid-Wales

L. Ness¹, C. Neal², T.D. Davies¹ and B. Reynolds³

¹School of Environmental Sciences, University of East Anglia, Norwich, NR4 7TJ, UK

²Centre for Ecology and Hydrology, Maclean Building, Crowmarsh Gifford, Wallingford, Oxon, OX10 8BB, UK

³Centre for Ecology and Hydrology, Bangor Research Unit, UWB, Deiniol Road, Bangor, Gwynedd, LL57 2UP, UK

Email for corresponding author: L.Ness@uea.ac.uk

Abstract

Analysis of winter stream chemistry data from the Afon Hafren in mid-Wales reveals links between stream chemistry and the North Atlantic Oscillation (NAO). K, Y, Al and dissolved organic carbon (DOC) concentrations increase during high NAO index months (relatively warm and wet weather), while Ca, Mg and NO₃ concentrations increase during low NAO months (relatively cold and dry conditions) with the increased concentrations lasting into the next month. The cause of the concentration changes varies between constituents: Y, Al and DOC are impacted by short term variations in rainfall; Mg and Ca are influenced by flow conditions; and NO₃ is temperature controlled. The dominant control mechanism for K concentration seems to be related to rainfall after tree felling has taken place but, prior to felling, the relationships are less certain.

Keywords: nitrate, potassium, yttrium, aluminium, Dissolved Organic Carbon, magnesium, calcium, North Atlantic Oscillation, stream chemistry, Plynlimon, Hafren

Introduction

Many of the varied mechanisms controlling stream chemistry are influenced by climatic conditions affecting either catchment inputs or within-catchment processes (Monteith *et al.*, 2000; Neal *et al.*, 2001). Variations in the climate of the North Atlantic can be characterised through the use of the North Atlantic Oscillation (NAO) index which is the difference in sea-level atmospheric pressure between the Icelandic low and Azores high pressure systems. The pressure difference influences the climatic conditions experienced across north-western Europe, especially during the winter months. A positive NAO index in the winter translates to relatively warm and wet conditions as Atlantic frontal systems are driven across the region. Correspondingly, negative NAO winters are cold, dry and dominated by air flow from the north-east (Hurrell, 1995). Changes in the strength of the prevailing wind conditions can cause variations in the total atmospheric deposition of sea-salts, affecting fresh water chemistry (Evans *et al.*, 2001) and associated variations in temperature have been linked

to changes in stream NO₃ concentrations (Monteith *et al.*, 2000). Recently, the NAO index has been showing a trend towards the positive phase, which may be due partly to increasing atmospheric greenhouse gas concentrations (Gillett *et al.*, 2003), and may well continue into the future. The change in winter conditions and increased frequency of extremes may lead to greater variability in stream chemistry and impact stream ecosystem function (Bradley and Ormerod, 2001). However, these changes will be predictable only if the characters of the impacts of the different aspects of the NAO are understood.

This study investigates the links between streamwater NO₃ concentrations at Plynlimon, mid-Wales, and the NAO. Previous work at this site has considered only the responses of stream NO₃ concentrations in the moorland areas over the December–March time period. The analysis has now been extended to consider the responses of K, Y, Al, DOC, Ca and Mg concentrations to the NAO in both the moorland and the forested part of the catchment on a monthly timescale. The Afon Hafren water quality data set used here

is one of the longest series of its kind in the UK, with ongoing data being collected since 1983 (Neal *et al.*, 1997, 2001).

Study area and methods

The Afon Hafren, one of the main tributaries of the River Severn, is located on the south-eastern slopes of Plynlimon in mid-Wales, about 30 km inland from the Irish Sea. The stream drains a catchment of 347 ha, spanning an altitude range of 350–690 m (Neal *et al.*, 2004). The catchment's bedrock consists of lower Palaeozoic mudstones, greywackes, sandstones and grits while the soils in the catchment are thin (around 70 cm deep) and acidic, consisting of mainly stagnopodzols, peats, brown earths and gleys, apart from on the mountain-top plateau where deep peat predominates. The vegetation within the catchment changes from semi-natural acidic moorland near the stream's source to Sitka spruce plantation forestry further down the catchment. The forest was planted between 1937 and 1964 and remained intact until the late 1980s when phased harvesting began.

Air temperature in the catchment (measured at the Moel Cynnedd meteorological station at 350 m altitude until 1999) ranged from -13.7°C to 30.7°C over the period 1983–1999, with a mean daily minimum of 4.0°C and a mean daily maximum of 11.3°C . The catchment receives around 2500 mm of rainfall per year and losses through evapotranspiration are 500–700 mm per year. The stream flow is 'flashy', responding rapidly to rainfall inputs. Streamwater chemistry has been shown to vary with the proportion of waters from soil and groundwater sources entering the stream (Neal *et al.*, 1990; Foster *et al.*, 2001).

The hydrochemistry of the Afon Hafren has been studied since May 1983, with weekly instantaneous grab sampling of the stream just upstream from the confluence with the Afon Hore (Lower Hafren site). In July 1990, further grab sampling was started at a site about 200 m above the upper edge of the forest (Upper Hafren site). Rainfall has been sampled weekly since May 1983 using two continuously open bulk collectors which are located at 570 m in the Hafren catchment (Fig. 1) and at 390 m in the neighbouring Gwy catchment, 3 km away (Kirby *et al.*, 1991). The stream and rainwater samples were analysed in the same manner for a wide range of chemical constituents including NO_3 , K, Y, Al, DOC, Ca and Mg (Neal *et al.*, 1997; Wilkinson *et al.*, 1997). After an initial trial period, the rainfall samples from different altitudes were combined and analysed together as their chemical concentrations were very similar (Wilkinson *et al.*, 1997).

Using monthly NAO indices (<http://www.cru.uea.ac.uk/cru/data/nao.htm>), the mean NAO index was calculated for

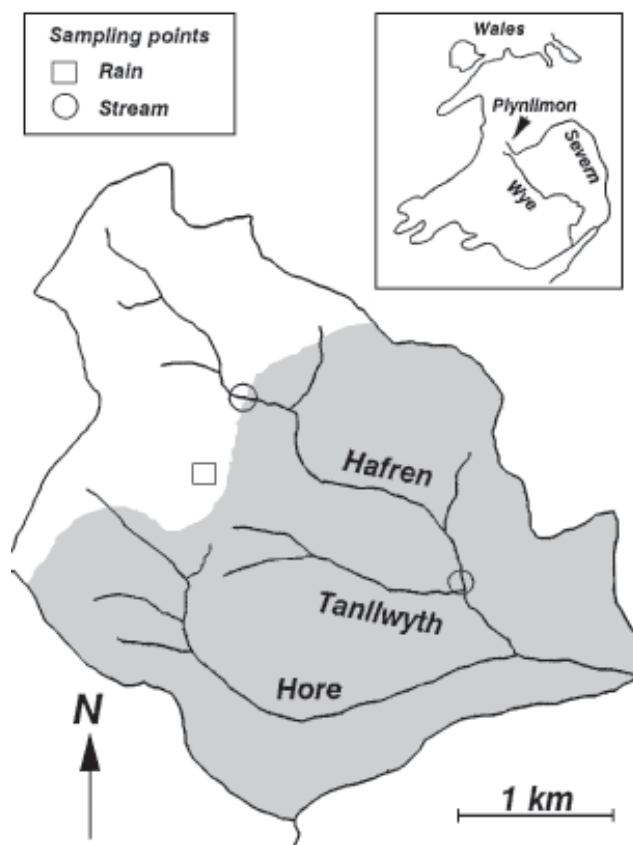


Fig. 1. The Plynlimon study area.

the periods December–February, December–March and January–March of each year from 1983–2002 (with the year referring to when the winter season finished). These periods were selected as they cover the time of annual minimum temperature and the part of the year when the NAO index is most clearly defined (Portis *et al.*, 2001). Monthly rainfall totals were obtained from the Carreg Wen meteorological station within the Hafren catchment. Mean monthly temperature data were available from RAF Valley (located 100 km away on Anglesey) and the CRU TS 2.0 temperature series (Mitchell *et al.*, 2003). The first of these was site-specific and the second a regional generalisation (on a 0.5° grid). However, the data sets were highly correlated with one another (Pearson correlation, $r^2 = 0.979$, $p < 0.01$) so the CRU TS 2.0 data were selected because this study was trying to link general climate patterns to local stream events.

Monthly and December–February, December–March and January–March mean stream and rain concentrations were calculated for each chemical constituent. The corresponding atmospheric deposition totals were calculated using bulk rainfall concentration data and meteorological station rainfall totals.

Non-parametric (Spearman) correlation analysis was used

to determine if there was a significant relationship between the monthly mean streamwater concentrations of different species, monthly mean air temperature, monthly rainfall total and monthly NAO index for each winter month between December 1983 and March 2002. The same analysis was used for the December-February, December-March and January-March mean concentrations and indices. If the stream concentrations of any constituent showed a trend of increasing/decreasing concentration over time, regression analysis was used to test its significance. If the trend was significant at the $p < 0.05$ level, a correlation analysis was done

- (i) between the concentrations and the NAO for the period the trend was occurring;
- (ii) between the residuals from the regression and the NAO indices.

Results and discussion

The individual monthly NAO indices were highly significantly correlated with the corresponding month's rainfall total and mean monthly temperature for the CRU TS 2.0 grid square covering Plynlimon (Spearman's Rank correlation, $r^2 = 0.631$ and 0.546 , respectively, $p < 0.01$ for both).

NITRATE

Monthly mean NO_3 concentrations in the Upper Hafren ranged from 0.70 to $2.67 \text{ mg NO}_3 \text{ l}^{-1}$ between 1990 and 2002 (Fig. 2). Over this time, the mean stream concentrations declined slightly (linear regression: $r^2 = 0.344$, $p < 0.05$ for December-March concentrations, also commented on by Neal *et al.*, 2001). The period of elevated concentrations in the winter of 1996 was also observed at several other UK locations (Evans and Monteith, 2001). The total rainfall in each winter month was significantly inversely related to the mean stream NO_3 concentration (Table 1) but NO_3 concentrations were not related significantly to either the temperature or NAO index of the same month. However, if a lag time of a month was introduced, a significant link between the NO_3 concentration of one month and the total rainfall, mean temperature and NAO index of the previous month became apparent.

Decreases in atmospheric deposition totals could be the cause of the long-term decline in NO_3 concentrations in the Upper Hafren. Small declines in NO_3 deposition have been observed in some European locations, but in many places these are not statistically significant due to 'noise' in the signal (Wright *et al.*, 2001). However, Irwin *et al.* (2002) observed an average annual decrease in rainfall NO_3 concentration of 0.1% between 1986 and 1999 at Plynlimon, while Reynolds *et al.* (1999) found a small decrease in wet deposited $\text{NO}_3\text{-N}$ between 1984 and 1995 over Wales. From

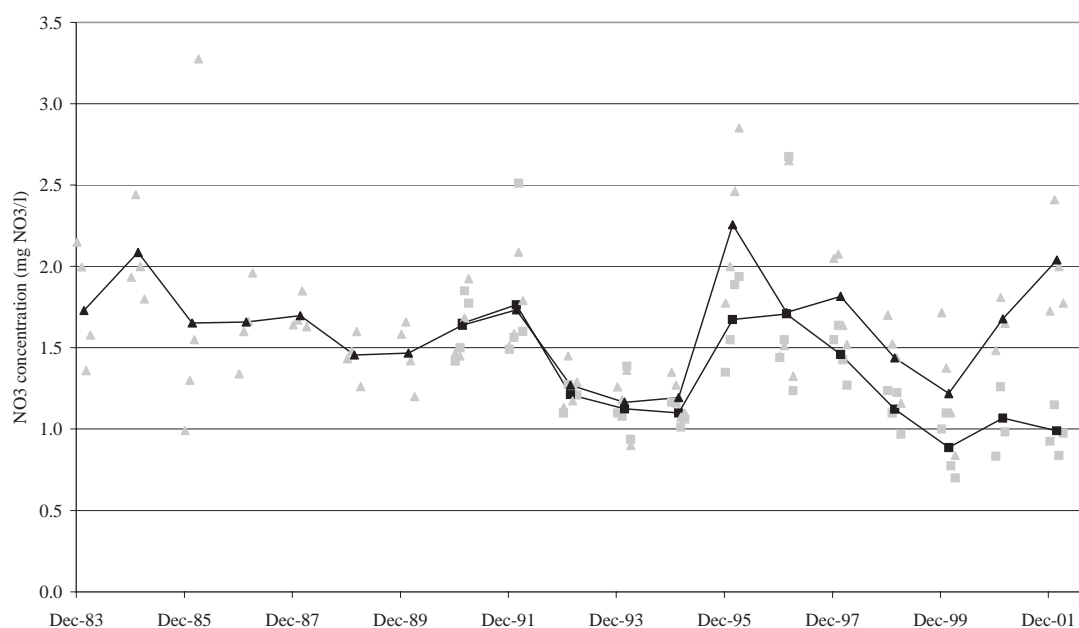


Fig. 2. NO_3 concentrations in the Upper and Lower Hafren. Individual points show the mean concentrations for December, January, February and March in each year while the lines represent the mean December-March concentrations for each year. The Upper Hafren concentrations are represented by squares, the Lower Hafren values by triangles.

Table 1. Spearman's correlation coefficients (r^2) for NO_3 with climatic variables. UH=Upper Hafren, LH=Lower Hafren. Significance levels: ** = $p < 0.01$, * = $p < 0.05$, ns = not significant ($p > 0.05$).

| NO_3 concentration (mg l^{-1}) | Independent variable | | | Coefficient when dependent variable lags the independent variable by a month | | |
|--|-----------------------------|---|-------------------|--|---|-------------------|
| | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index |
| | UH | -0.414 ** | ns | ns | -0.473 ** | -0.380 * |
| LH | -0.265 * | -0.288 * | -0.272 * | -0.494 ** | ns | -0.509 ** |
| LH 1984-1993 | ns | ns | ns | -0.512 ** | -0.339 * | -0.410 * |
| LH 1994-2002 | ns | ns | -0.373 * | -0.462 ** | ns | -0.624 ** |

the bulk rainfall data collected in the Hafren catchment, the annual NO_3 deposition declined from over $20 \text{ kg ha}^{-1} \text{ yr}^{-1}$ in early 1990s to around $15 \text{ kg ha}^{-1} \text{ yr}^{-1}$ in 2001. However, none of these estimates took any account of occult or dry deposition other than through the unknown and variable amount of both contained in the wet deposition sampled by open collectors.

Rainfall NO_3 concentrations over Plynlimon were affected significantly by the NAO, with high NO_3 concentrations occurring in low NAO index months when there was an increased probability of air flow from Europe. However, NO_3 wet deposition totals were not significantly related to

the NAO index in any way. The residuals of the regression of monthly Upper Hafren NO_3 concentration were significantly negatively correlated with the NAO of the previous month (Spearman's Rank correlation, $r^2 = -0.513$, $p < 0.01$, compared to $r^2 = -0.391$, $p < 0.01$ for measured stream NO_3 concentrations). The pattern of NO_3 concentration change seen in the Upper Hafren seems, therefore, to represent a general decline in stream NO_3 concentrations over time with climatically influenced NO_3 changes superimposed.

Over the study period, 1996 was the first negative NAO winter since 1985 (Fig. 3) and the most negative NAO year

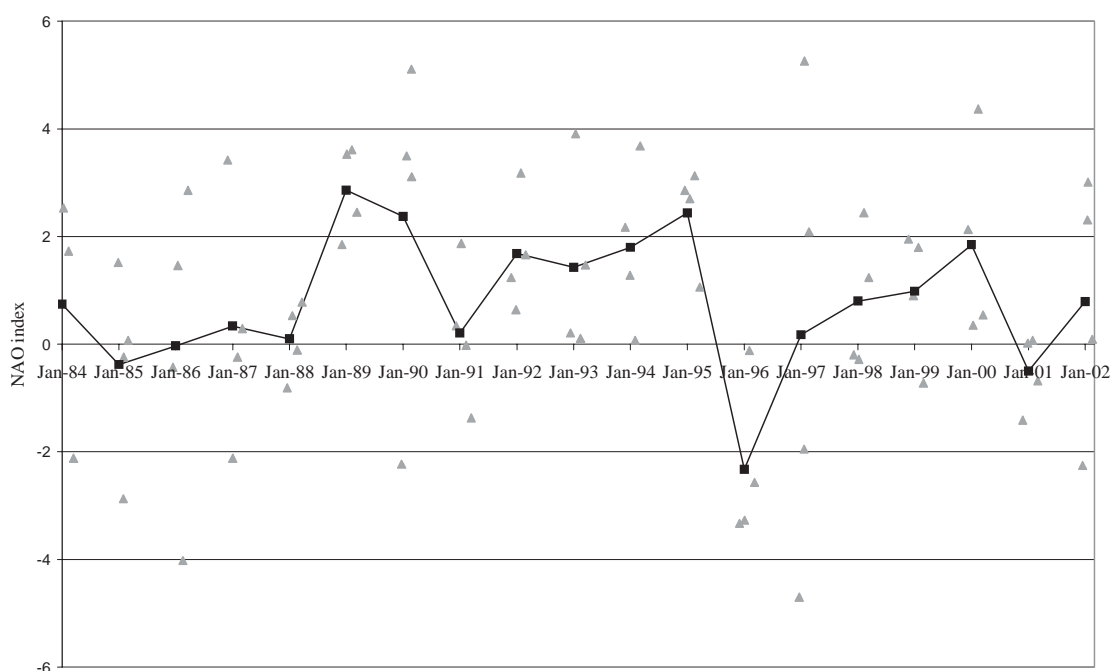


Fig. 3. Winter (December-March) North Atlantic Oscillation Index values 1983-2002. Triangles represent the North Atlantic Oscillation Index values for the individual winter months; squares represent the December-March mean value.

since the 1820s. This winter had very high stream NO_3 concentrations, probably caused by the extreme climate conditions (the mean minimum and maximum air temperatures recorded at the Moel Cynnedd meteorological station for winter 1996 were -0.94°C and 4.16°C , with an absolute minimum of -10.4°C whereas the mean minimum and maximum air temperatures for all the winters 1983–1998 were 0.42°C and 5.91°C , with only eight days overall recording lower temperatures than the minimum seen in 1996). High stream NO_3 concentrations in other locations have been linked to freezing events killing fine plant roots and disrupting microbe, soil and plant interactions. This results in the release of NO_3 into the soil, and as uptake by damaged roots is limited, allows NO_3 to be washed into the streams (Monteith *et al.*, 2000; Fitzhugh *et al.*, 2003). Monteith *et al.* (2000) reported a strong link between the winter peak NO_3 concentration and the December-March NAO index in the Afon Gwy, another Plynlimon stream, and linked this to temperature. Therefore, as high NO_3 concentrations in the Upper Hafren were associated with low NAO events (i.e. cold and dry conditions) this temperature dependent mechanism was probably operating in the Hafren's catchment as well. The present study shows that there are also links between the NAO and stream chemistry on time-scales shorter than the four month period considered by Monteith *et al.* (2000). A weak negative relationship was seen between the monthly mean NO_3 concentration in the Lower Hafren and the NAO index. However, correlations between monthly mean stream NO_3 concentrations and monthly NAO indices were stronger in both the Upper and Lower Hafren when a time lag of one month was introduced between the NAO index month and that of the mean concentration (Table 1). For example, the correlation between Upper Hafren NO_3 concentration and the monthly NAO index was not significant with no lag, but became significant at the $p < 0.01$ level ($r^2 = 0.391$) when a lag of 1 month was introduced; the correlation between Lower Hafren NO_3 concentration and the monthly NAO index increased from $r^2 = 0.272$ ($p < 0.05$) to $r^2 = 0.509$ ($p < 0.01$) when the monthly lag was introduced. These time lags could possibly be caused by the various different biological and hydrological processes involved in NO_3 transport.

In the Lower Hafren, mean monthly NO_3 concentrations ranged from 0.84 to $3.28 \text{ mg NO}_3 \text{ l}^{-1}$ and mainly remained within the range 1 – $2 \text{ mg NO}_3 \text{ l}^{-1}$ (Fig. 2). Between 1983 and 1994, NO_3 concentrations also appeared to decline in the Lower Hafren (linear regression, $r^2 = 0.403$, $p < 0.05$ for December-March mean concentrations). From 1991 to 1993, NO_3 concentrations at the two Hafren sites were very similar. However, by 1994, NO_3 concentrations in the Lower Hafren

were beginning to increase above those in the Upper Hafren. The relative increase in NO_3 concentrations was due to forest harvesting between the two measuring stations. In the Hafren Forest, the main felling technique leaves the tree stumps, needles and branches *in situ* (Neal, 2002a,b). When felling is carried out, biological NO_3 demand from the trees is eliminated and the damaged and decaying vegetation releases NO_3 , allowing more NO_3 to enter streams, with the effects lasting for a few years (Neal *et al.*, 2001).

The data for the Lower Hafren were split into two periods (1984–1993 and 1994–2002) to see whether the relationship between stream NO_3 concentrations and the NAO altered after felling. Between 1984 and 1993, the Lower Hafren NO_3 concentrations showed the most significant relationship with the previous month's rainfall and weaker relationships with the previous month's NAO index and temperature (Table 1). However, in the period 1994 to 2002, the most significant relationships were with the NAO index and rain of the previous month. Correspondingly, there was no significant relationship with temperature, suggesting that the temperature dependent freezing process may have decreased in importance as the catchment conditions changed. Variations in stream NO_3 concentration post-felling may have been caused by NO_3 released from decaying felled material collecting in the soil until it was washed out by rainfall, so high concentration pulses reached the stream during rainfall following dry months. The NAO may have become a more important factor due to its greater correlation with rainfall than with temperature. Therefore, in the Hafren, it appears that NO_3 is significantly affected by the NAO but the effect can be altered by changes in catchment management.

Over timescales greater than a month, the link between mean NO_3 concentration and the NAO index was still apparent at times. The long lasting effects were visible as the mean January-March concentration in the Lower Hafren post-felling showed no significant correlation with the January-March NAO but was significantly correlated with the December-February NAO for this period (Spearman's Rank correlation $r^2 = -0.817$, $p < 0.01$).

POTASSIUM

Between 1991 and 2002, mean monthly K concentrations in the Upper Hafren varied between 0.13 and 0.28 mg l^{-1} (Fig. 4). There was no clear trend over time and a majority of the monthly mean concentrations were between 0.15 and 0.20 mg l^{-1} . Potassium concentrations in the stream were only significantly correlated with rainfall totals (Table 2). In the Lower Hafren, K concentrations remained between 0.15 mg l^{-1} and 0.20 mg l^{-1} until 1994, when they began to

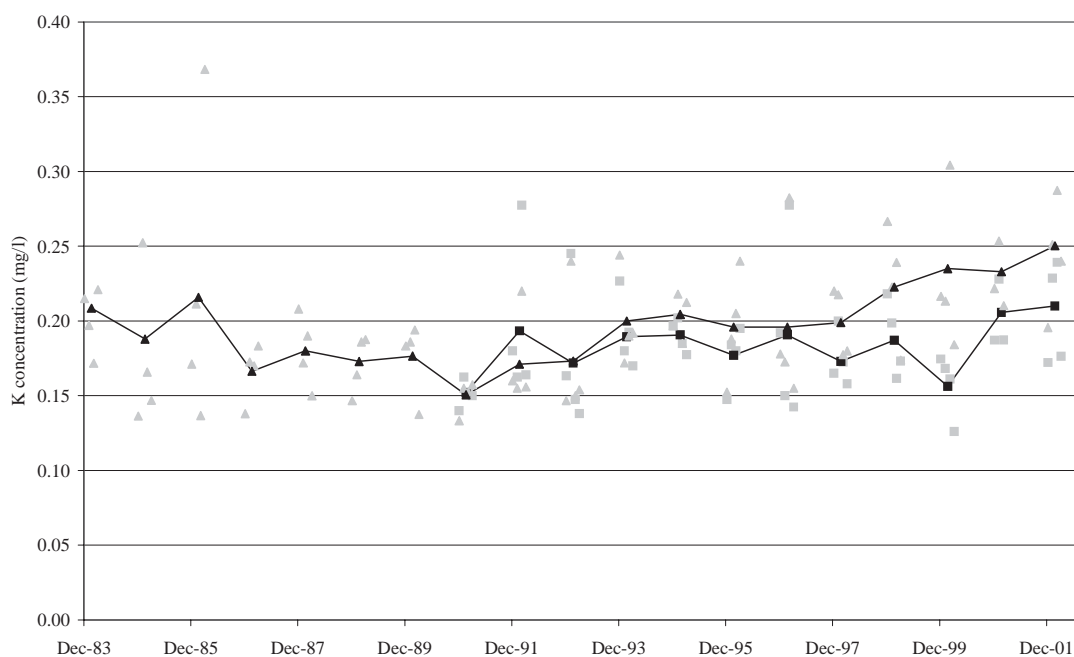


Fig. 4. K concentrations in the Upper and Lower Hafren. Key as for Fig. 2.

Table 2. Spearman’s correlation coefficients (r^2) for K with climatic variables. UH=Upper Hafren, LH=Lower Hafren. Significance levels: ** = $p < 0.01$, * = $p < 0.05$, ns = not significant ($p > 0.05$).

| K concentration (mg l ⁻¹) | Independent variable | | | Coefficient when dependent variable lags the independent variable by a month | | |
|--|--------------------------------|----------------------------------|----------------------|---|-------------------------------------|-------------------------|
| | Monthly rainfall total (mm) | Mean monthly temperature (°C) | Monthly NAO index | Monthly rainfall total (mm) | Mean monthly temperature (°C) | Monthly NAO index |
| | UH | 0.400 ** | ns | ns | ns | ns |
| LH | 0.292 * | ns | ns | ns | ns | ns |
| LH 1983-1992 | ns | ns | ns | ns | ns | ns |
| LH 1993-2002 | 0.421 ** | ns | ns | ns | ns | ns |

increase steadily. The increase in stream water K concentration was caused by felling as K is released rapidly into the soil by damaged and decaying vegetation (Neal *et al.*, 1992). Prior to 1993, K concentration in the Lower Hafren showed no significant correlation with the NAO index, rainfall totals or temperature but for the period 1993 to 2002, K became significantly correlated with rainfall totals from the same month (Table 2). This again indicated that felling has altered the effect that climate has on stream chemistry variations.

Unlike NO₃, total monthly bulk atmospheric deposition of K over the Hafren catchment was strongly related to the

month’s NAO index (Spearman’s rank correlation, $r^2 = 0.649$, $p < 0.01$). During high NAO index months, the rainfall totals at Plynlimon were larger and the rain came mainly from airstreams crossing the Atlantic, a significant source of K to the atmosphere. The K concentrations in the Upper Hafren were linked significantly to atmospheric deposition total (Spearman’s rank correlation, $r^2 = 0.399$, $p < 0.01$). However, stream K concentrations were not significantly linked to the NAO index. In the Lower Hafren, the concentration of K before 1993 showed no significant relationship with bulk deposition or rainfall K concentration, possibly because the ecosystem is K limited (Emmett *et al.*,

1995) and K is tightly cycled by the plants in the soil-plant system (Hornung *et al.*, 1986). After felling, K concentration in the Lower Hafren showed a positive relationship with both deposition and total rainfall (Spearman's Rank correlation $r^2 = 0.372$, $p < 0.05$ and $r^2 = 0.421$, $p < 0.01$) but, again, not with the NAO index. As rainfall is a major driver of bulk deposition total, the significant correlation between deposition and stream K concentration is probably a result of auto-correlation and it is therefore possible that the link between deposition and stream concentrations at this site is apparent only because rainfall, which is important for deposition totals, washes out the K which is already present in the soil. Elsewhere in the catchment, K concentrations in stream water have increased during storm events as storm flow becomes enriched in K by its passage through the soil organic horizon (Chapman, 1994). Thus:

- (i) in the Upper Hafren and after felling in the Lower Hafren (when the availability of K in the soil would be greater than before) such rainwater flushing events are probably occurring;
- (ii) the predominant control on K concentrations in stream water appears to be hydrological in nature;
- (iii) although variations in K concentration in the Upper and Lower Hafren could not be linked directly to the NAO index, atmospheric circulation is still important as rainfall variations in Europe are linked to NAO changes.

The same frost damage mechanism causing NO_3 release from plants has been proposed to have a significant effect

on stream K concentrations (Fitzhugh *et al.*, 2003) but there is little evidence that this occurred at Plynlimon. Hubbard Brook (Boormann and Likens, 1979; Likens and Boormann, 1995) in the north-eastern USA, where the frost mechanism was identified, is mostly second-growth deciduous northern hardwood forest and the region suffers more extreme winter conditions than Plynlimon. The difference in the catchments' climate and ecosystems could account for the very different relationships between K and climate in the two locations. This shows that stream chemistry behaviour in one catchment cannot be used easily to predict behaviour in another when there are major differences in catchment characteristics and climate.

YTTRIUM, ALUMINIUM AND DISSOLVED ORGANIC CARBON

Yttrium, Al and DOC all showed very similar behaviour to each other and at both sampling sites on the Hafren. Yttrium concentrations in the Upper Hafren varied between 0.12 and 0.40 $\mu\text{g l}^{-1}$ from 1990 to 1998. The range of Y concentrations in the Lower Hafren over the study period was similar and although they showed no trend, any month to month concentration variability was reflected at both stream sites (Fig. 5). Aluminium concentrations in the Upper Hafren fluctuated between 19 and 215 $\mu\text{g l}^{-1}$. Those in the Lower Hafren were consistently higher, varying between 56 and 380 $\mu\text{g l}^{-1}$ (Fig. 6). Again, there were considerable changes from month to month, but the changes were always

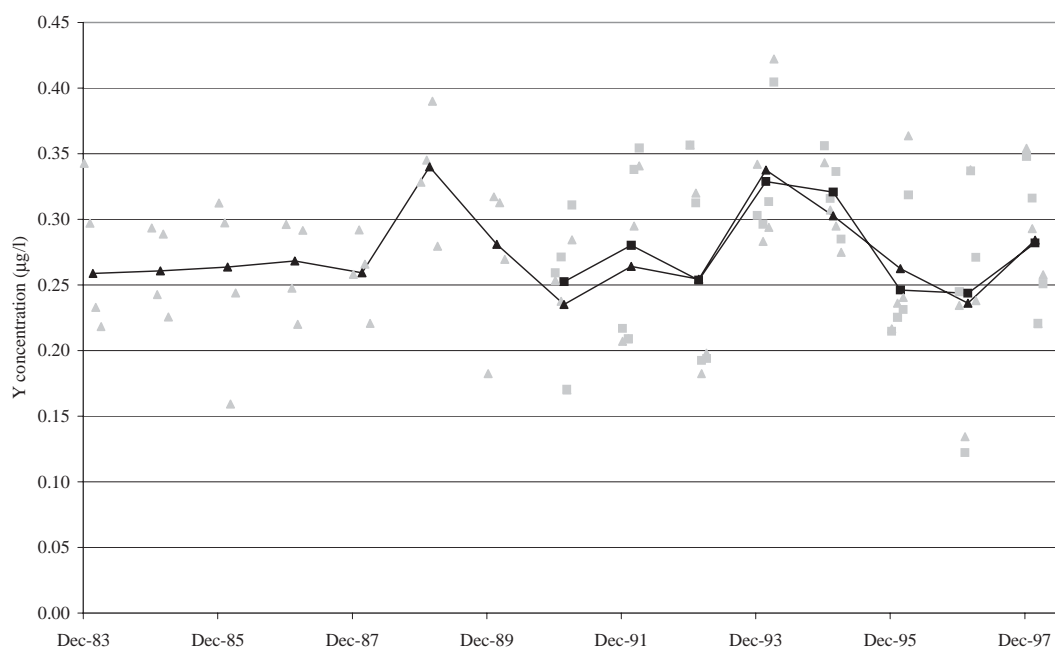


Fig. 5. Y concentrations in the Upper and Lower Hafren. Key as for Fig. 2.

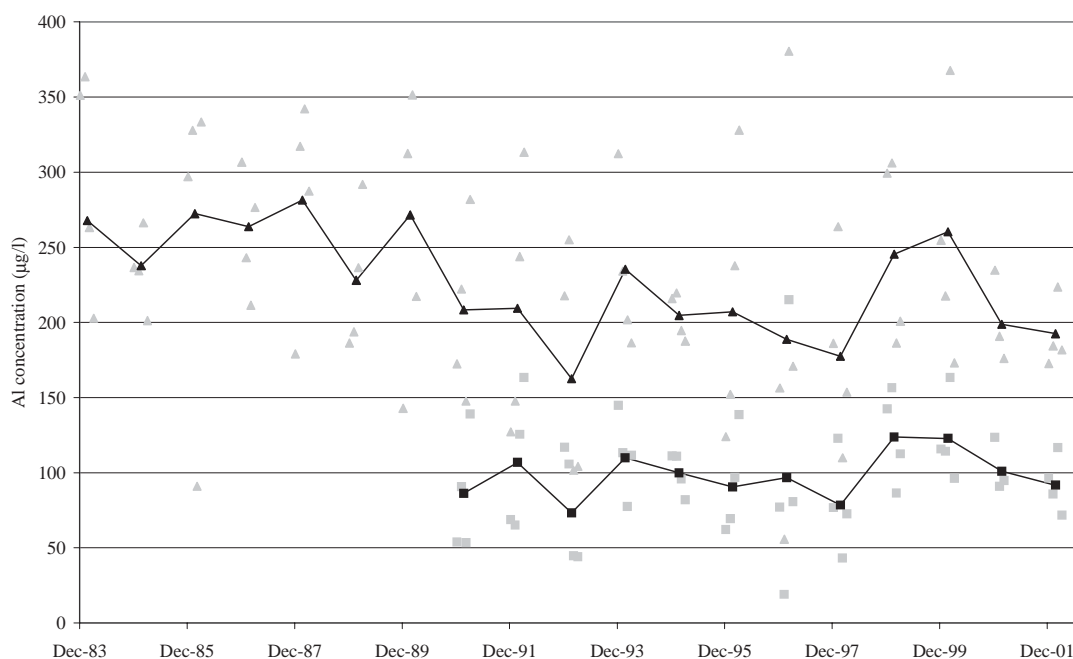


Fig. 6. Al concentrations in the Upper and Lower Hafren. Key as for Fig. 2.

reflected at both sites. Dissolved Organic Carbon concentrations in the Upper Hafren ranged between 0.3 and 1.9 mg l⁻¹ and the Lower Hafren showed a very similar pattern of change, except the concentrations were slightly higher (Fig. 7). Yttrium, Al and DOC are linked in that they occur at their highest concentrations in the Plynlimon soil waters, although there is a delineation between acidic, organic rich soil waters (but Al and Y depleted) in their upper horizons where organic matter predominates and less acidic (but Al and Y enriched waters) in the lower soils due to aluminium oxide/hydroxide rich materials in the soil matrix of the lower horizons. Both Y and Al are derived from the lower soil horizons by acid dissolution of oxide and hydroxide phases in the lower soils. Yttrium, Al and DOC concentrations are highest in the streams at high flows as it is at this time when the waters are derived mainly from both the upper and lower soils zones (Neal *et al.*, 1997): baseflow waters are of low Y, Al and DOC concentrations due to decomposition processes in the case of DOC and hydrolysis reactions for Y and Al under the less acidic conditions that appertain (Al and Y hydroxides can precipitate under higher pH conditions). Al and Y are approximately linearly related, and their correlation with DOC relates to the mixing of soil and groundwaters rather than a major linkage over speciation (for example, organic matter complexes with Al are relatively small (Neal, 1995)).

The mean monthly concentrations of Y, DOC and Al were all closely correlated ($p < 0.01$) in the Upper Hafren and

showed significant correlations with the month's NAO index (Table 3). The relationship between concentrations and the NAO index was of a very similar significance in the Lower Hafren. The strongest correlations for solute concentrations were with rainfall totals (Table 3), with the correlations with temperature being only slightly weaker. There was no correlation between monthly bulk deposition totals and stream concentration for these substances. The most significant correlations were those without a time-lag, suggesting a rapid response mechanism. The rainfall aspect of the NAO index was probably the controlling variable as storm flow waters, created by runoff originating from soil water and surface horizons, have higher Y, Al and H⁺ concentrations than base flow waters as mentioned above. Al and Y are known to undergo rapid flushing during rainfall events (Soulsby and Reynolds, 1993); this possibly helps to explain the lack of correlation between the concentrations of one month and the conditions of the previous month. Storm flow is also known to flush DOC out of the soil, creating high concentration events, especially when the flow path of the storm water avoids mineral horizons within the soil and antecedent conditions have been relatively dry (Eatherall *et al.*, 1998).

If the longer term December-March indices were used, there was no significant correlation between the NAO index and Al or DOC concentrations. However, the December-March mean Y concentration in the Upper Hafren was correlated significantly with the December-March NAO

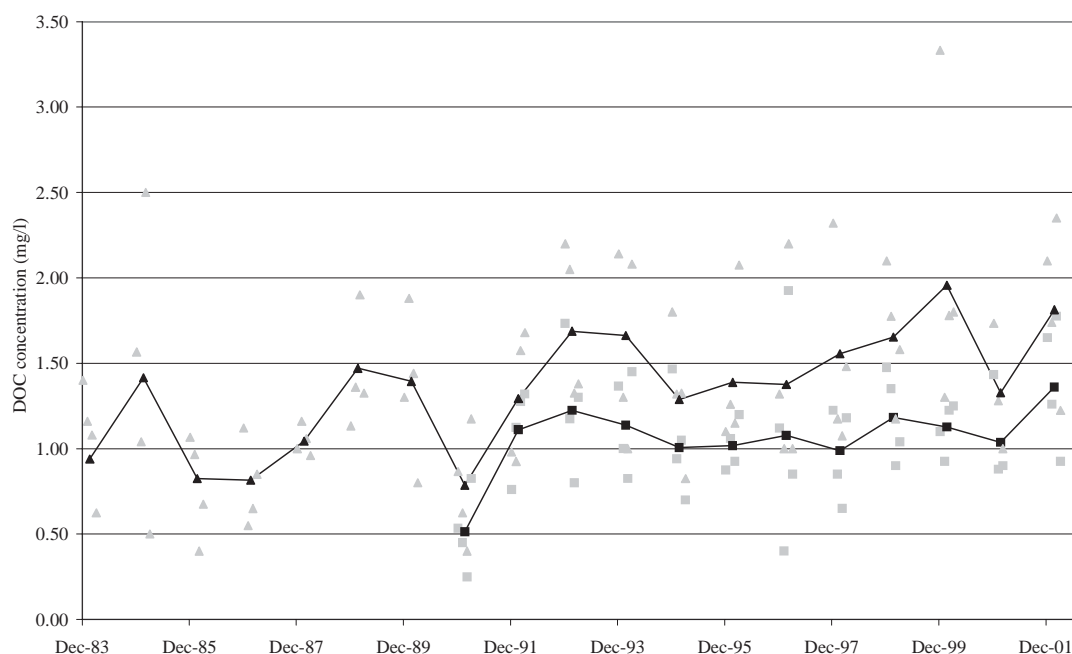


Fig. 7. DOC concentrations in the Upper and Lower Hafren. Key as for Fig. 2.

Table 3. Spearman's correlation coefficients (r^2) for Y, Al and DOC with climatic variables. UH=Upper Hafren, LH=Lower Hafren. Significance levels: ** = $p < 0.01$, * = $p < 0.05$, ns = not significant ($p > 0.05$).

| Y, DOC (mg l^{-1}) and Al ($\mu\text{g l}^{-1}$) conc. | Independent variable | | | Coefficient when dependent variable lags the independent variable by a month | | |
|---|-----------------------------|---|-------------------|--|---|-------------------|
| | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index |
| | UH Y | 0.635 ** | 0.491 ** | 0.442 * | ns | ns |
| UH Al | 0.599 ** | 0.371 * | 0.296 * | ns | ns | ns |
| UH DOC | 0.476 ** | 0.451 ** | 0.318 * | ns | ns | ns |
| LH Y | 0.610 ** | 0.466 ** | 0.494 ** | ns | 0.255 * | ns |
| LH Al | 0.460 ** | ns | 0.352 ** | ns | ns | ns |
| LH DOC | 0.399 ** | 0.456 ** | 0.284* | ns | 0.252* | ns |

index (Spearman's Rank correlation, $r^2 = 0.717$, $p < 0.05$) and was significant in the Lower Hafren too (Spearman's Rank correlation $r^2 = 0.623$, $p < 0.05$). Yttrium concentrations were much lower than those of Al or DOC, so small stream concentration changes caused by rainfall totals may have been more visible and significant. This difference in relationship between long and shorter term periods shows the importance of considering processes at different timescales.

CALCIUM AND MAGNESIUM

Between 1990 and 2002, Ca concentrations in the Upper Hafren ranged from 0.37 to 0.71 mg l^{-1} (Fig. 8). There was no trend in the data, but mean concentrations varied substantially from month to month and were consistently around 0.3 mg l^{-1} greater in the Lower Hafren than the Upper Hafren. The Ca concentration in the Upper Hafren was significantly negatively related to that month's rainfall total, NAO index and mean temperature and the previous month's NAO (Table 4, Spearman's Rank correlations, $r^2 = -0.776$,

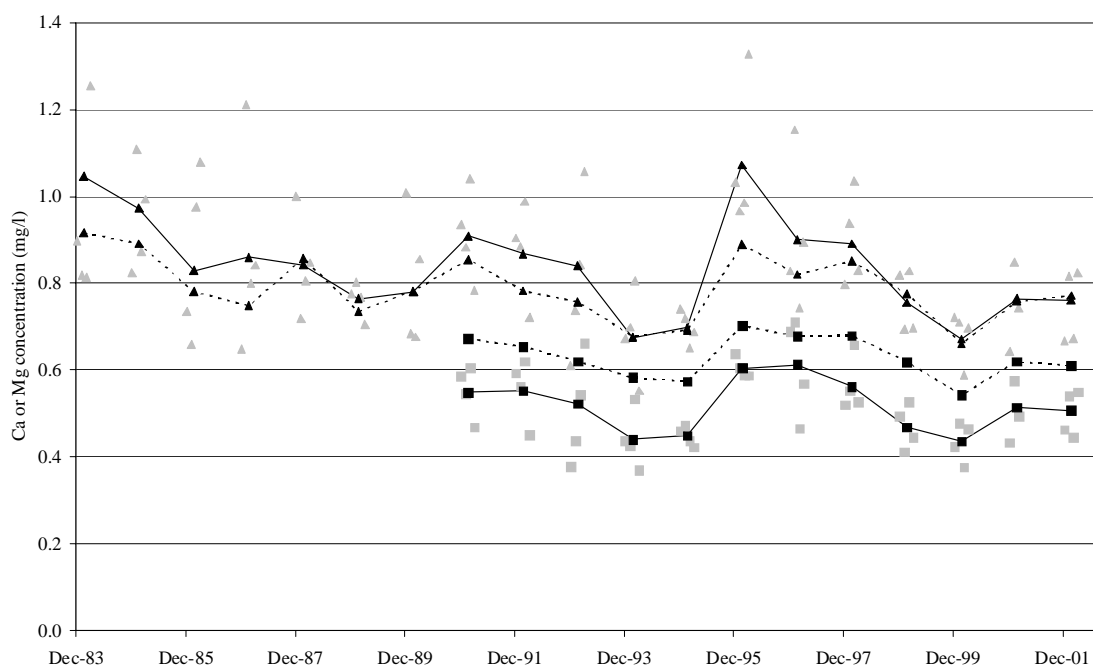


Fig. 8. Ca and Mg concentrations in the Upper and Lower Hafren. Individual points show the mean Ca concentrations for December, January, February and March in each year while the solid lines represent the mean December-March Ca concentrations for each year. The Upper Hafren Ca concentrations are represented by squares, the Lower Hafren Ca values by triangles. The Upper Hafren mean December-March Mg concentrations are represented by squares joined by a dashed line and the Lower Hafren mean December-March Mg concentrations for each year are represented by triangles joined by a dashed line. Individual monthly Mg concentrations are not shown.

Table 4. Spearman's correlation coefficients (r^2) for Ca and Mg with climatic variables. UH=Upper Hafren, LH=Lower Hafren. Significance levels: ** = $p < 0.01$, * = $p < 0.05$, ns = not significant ($p > 0.05$).

| Ca or Mg concentration (mg l^{-1}) | Independent variable | | | Coefficient when dependent variable lags the independent variable by a month | | |
|---|-----------------------------|---|-------------------|--|---|-------------------|
| | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index | Monthly rainfall total (mm) | Mean monthly temperature ($^{\circ}\text{C}$) | Monthly NAO index |
| | UH Ca | -0.776 ** | -0.392 ** | -0.395 ** | ns | ns |
| UH Mg | -0.561 ** | -0.346 * | ns | -0.351 * | -0.310 * | -0.483 ** |
| LH Ca | -0.690 ** | -0.416 ** | -0.428 ** | -0.363 ** | ns | -0.304 ** |
| LH Mg | -0.527 ** | -0.396 ** | -0.316 ** | -0.356 ** | -0.527 ** | -0.352 ** |

-0.395, -0.392 and -0.410, $p < 0.01$). In the Lower Hafren, the significant correlations were the same, with the addition of a significant link between the Ca concentration and the rainfall total of the previous month (Table 4). Magnesium concentrations at both sites followed those of Ca very closely, suggesting similar controls and transport. In the Upper Hafren, the mean Mg concentration ranged between 0.47 and 0.74 mg l^{-1} . Again, any concentration changes were reflected at both sites with concentrations around 0.15 mg l^{-1} higher for the Lower Hafren. The Mg concentrations

in the Upper Hafren were significantly negatively correlated with the rain total, mean temperature and NAO index of the previous month and the rainfall and mean temperature of the same month (Table 4, Spearman's Rank correlations, $r^2 = -0.351, -0.310, p < 0.05, r^2 = -0.483, -0.561, p < 0.01, r^2 = -0.346, p < 0.05$). Magnesium concentrations in the Lower Hafren showed the same relationships and also were significantly correlated with the NAO index of the same month.

Felling activity increases Ca concentrations in the Afon

Hore, also on Plynlimon (Neal, 2002a,b). Ca concentrations in the Lower Hafren, however, do not appear to have been affected by felling. This difference may be due to the Afon Hafren catchment containing less Ca rich bedrock than the Afon Hore catchment (Reynolds *et al.*, 1987; Neal *et al.*, 1997), the hydrology of the Hafren catchment allowing less dominance of base rich waters than the Hore or the felling activity in the Hafren catchment being less intense.

The primary cause of the variation of Ca and Mg concentrations with the NAO index was not atmospheric deposition because this was positively correlated with the NAO while the stream concentrations were inversely related to the NAO. Base flow waters in the Hafren have higher Ca concentrations than storm flow waters as they contain proportionally more water from the lower soil and bedrock where the weathering of calcium carbonate and calcium bearing silicates has been occurring (Reynolds *et al.*, 1987; Neal *et al.*, 1997). Changes in the proportion of the stream's total flow coming from base flow are therefore a major control on the concentration of Ca in stream water (Reynolds *et al.*, 1987). Months with a low NAO index were relatively dry, so base flow would have dominated the stream flow and the concentration-flow diagram for Ca and Mg showed the pattern typical of a component controlled by a groundwater source (Fig. 9), suggesting that their concentrations are controlled indirectly by the NAO through the volume of water from different sources present in the stream. Calcium (and to some extent Mg) concentrations

Table 5. Spearman's correlation coefficients (r^2) for Ca, Mg, K, Y, Al and DOC stream water concentrations in the Lower Hafren. Significance levels: ** = $p < 0.01$, * = $p < 0.05$, ns = not significant ($p > 0.05$).

| | Ca | Mg |
|-----|----------|----------|
| K | -0.20 * | ns |
| Y | -0.66 ** | -0.55 ** |
| Al | -0.28 ** | ns |
| DOC | -0.49 ** | -0.40 ** |

were also significantly negatively correlated with Y, Al and DOC (Table 5), as would be expected if their sources were different waters. The concentration-flow diagrams of Y, Al and DOC also appeared typical of those of constituents controlled by run-off sources (data not shown). The negative relationship of Ca and K concentrations was weaker (and even insignificant in the case of Mg), indicating that K concentrations may be influenced by processes more akin to those controlling Ca and Mg. As there are weathering sources of K in the catchment in the form of illite (a K rich mica) in the bedrock (Reynolds *et al.*, 1997), this may account for the difference in relationship strength as base flow and storm flow waters (through the flushing of excess K from soils) would both be sources of K.

Some of the variation in stream Ca and Mg concentrations

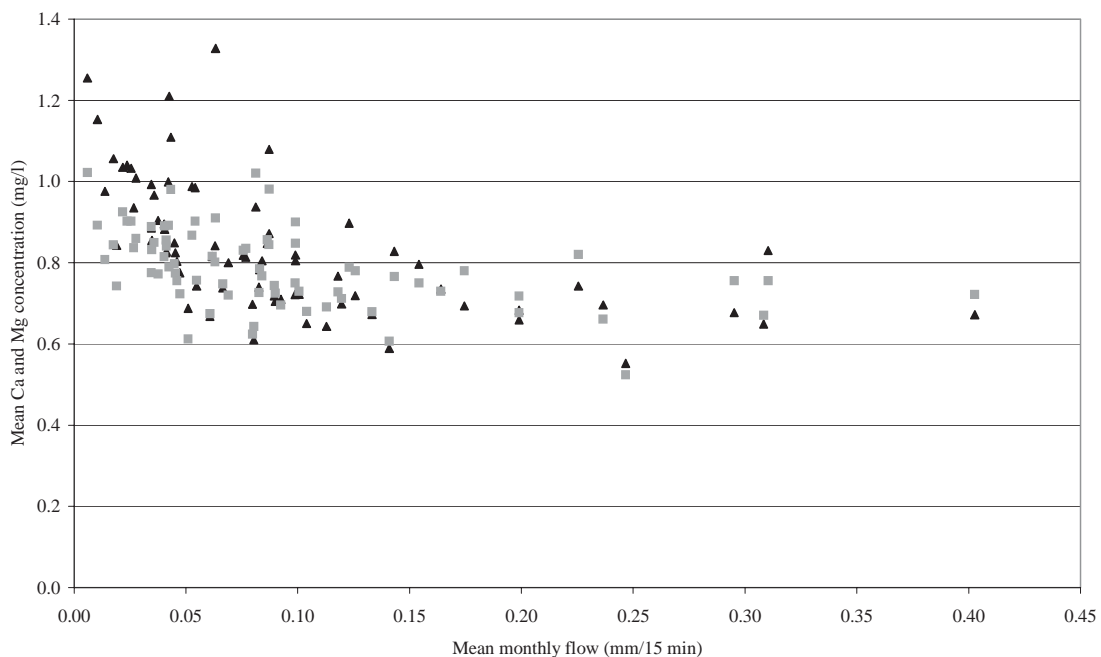


Fig. 9. Mean Ca and Mg concentration variations with flow volume in the Lower Hafren. Ca concentrations are represented by triangles and Mg by squares.

elsewhere has been linked to both rainfall totals and freezing events (Fitzhugh *et al.*, 2003). However, freezing explained less of the variation in Mg and Ca concentration in the Hubbard Brook study than it did for K and NO₃. In the Afon Hafren, Ca and Mg were both correlated with temperature, but the correlation with rainfall was stronger and so the major control probably was the amount of rainfall.

The January–March Ca and Mg concentrations in the Upper Hafren showed the strongest correlation with the December–February NAO index (Spearman's Rank correlation, $r^2 = -0.804$, -0.783 , $p < 0.01$, respectively). Therefore, the changes in Ca and Mg concentration due to the NAO are visible both immediately and for a while thereafter. The lag time seen in the longer term indices may be caused by different processes, such as freezing, but there is no clear evidence for this.

Conclusions

The long-term stream water quality record from the Afon Hafren in mid-Wales shows K, Y, Al and DOC concentrations increasing during high NAO index winter months, while Ca, Mg and NO₃ concentrations increase during low NAO index winter months, with the increased concentrations lasting into the next month. The mechanisms through which the NAO influences the concentrations of these different constituents in streamwater vary. Stream NO₃ concentrations in the Afon Hafren are influenced by the NAO through temperature regulated processes in the soil but changes in catchment management, such as tree felling, can increase the importance of rain driven flushing. Potassium concentrations are affected by the rainfall volume aspect of the NAO and, again, the controls can be altered by catchment management changes. Little evidence is seen in the Afon Hafren data that freezing processes significantly increase stream K concentrations although this has been observed in the eastern USA. Yttrium, Al and DOC stream concentrations are all related to the NAO through the amount of rainfall occurring and the volume of runoff generated. Links between the NAO and stream Y concentrations have not been observed before and only the use of monthly, rather than longer-term, NAO indices allowed them to be seen in this study. Calcium and Mg are also indirectly influenced by the NAO through control of the stream's flow regime. Again, no evidence of concentration increases linked to a temperature mechanism is seen at Plynlimon although this has been seen in the USA.

The value of long-term stream chemistry data sets such as those used here is great as the changes in chemical concentration caused by alterations to the catchment and catchment stores can be distinguished from those caused

by climate variability, a near impossibility with short data sets given the time-spans of the different types of drivers and responses. By establishing the relationships between aspects of long-term stream chemistry data sets and climate indices such as the NAO, the projection of future changes in some aspects of stream chemistry may be possible using this climate index. However, the exact value of the analysis will vary from constituent to constituent, depending on the method and time scale over which the NAO affects its concentration. Further, it will also be important to consider the different aspects of the NAO which may affect stream concentrations in different situations as the chemical response in streams flowing through different ecosystems has been shown to be dissimilar.

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