



# Using globally available soil moisture indicators for flood modelling in Mediterranean catchments

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**Abstract.** Floods are one of the most dangerous natural hazards in Mediterranean regions. Flood forecasting tools and early warning systems can be very beneficial to reducing flood risk. Event-based rainfall–runoff models are frequently employed for operational flood forecasting purposes because of their simplicity and the reduced number of parameters involved with respect to continuous models. However, the advantages related to the reduced parameterization oppose to the need of a correct initialization of the model, especially in areas characterized by strong climate seasonality. In this case, the use of continuous models could be desirable but it is very problematic in poorly gauged areas where continuous rainfall and temperature data are not available. This paper introduces a Simplified Continuous Rainfall–Runoff model (SCRMM), which uses globally available soil moisture retrievals to identify the initial wetness condition of the catchment, and, only event rainfall data to simulate discharge hydrographs. The model calibration involves only three parameters. For soil moisture, besides in situ data, satellite products from the Advanced SCATterometer (ASCAT) and the Advanced Microwave Scanning Radiometer for Earth observation (AMSR-E) sensors were employed. Additionally, the ERA-Land reanalysis soil moisture product of the European Centre for Medium-Range Weather Forecasting (ECMWF) was used.

SCRMM was tested in the small catchment of the Rafina River, 109 km<sup>2</sup>, located in the eastern Attica region, Greece. Specifically, sixteen recorded rainfall–runoff events were simulated by considering the different indicators for the estimation of the initial soil moisture conditions from in situ, satellite and reanalysis data. By comparing the performance of the different soil moisture products, we conclude

that: (i) all global indicators allow for a fairly good reproduction of the selected flood events, providing much better results than those obtained from setting constant initial conditions; (ii) the use of all the indicators yields similar results when compared with a standard continuous simulation approach that, however, is more data demanding; (iii) SCRMM is robust since it shows good performances in validation for a significant flood event that occurred on February 2013 (after calibrating the model for small to medium flood events). Due to the wide diffusion of globally available soil moisture retrievals and the limited number of parameters used, the proposed modelling approach is very suitable for runoff prediction in poorly gauged areas.

## 1 Introduction

In the context of climate change, in which runoff production mechanisms appear to be exacerbated by the modification of climatic variables, the flood frequency regime is altered and an increasing frequency of extreme events is to be expected. The Report of Intergovernmental Panel on Climate Change (IPCC, 2001) on potential effects of climate change highlights that “flood magnitude and frequency are likely (a 66–90 % probability) to increase in most regions”. Notwithstanding this issue, Europe seems to lack suitable and reliable procedures to promptly address the fundamental issues of flood-risk assessment and management. Even though several important laws and directives, both at national and at EU level, have been issued addressing this point, i.e. European Floods Directive 2007/60/CE, such legislative tools have not succeeded yet in effectively reducing the devastating and

catastrophic effects of extreme flood events (Barredo, 2006). To this end, flood forecasting and early warning systems have been identified as fundamental tools for prevention and protection from flood risk.

The development of an early warning system for flood forecasting is particularly difficult for small- to medium-sized basins (area < 400 km<sup>2</sup>) for which the hydrologic response is extremely fast and an hourly (or finer) temporal resolution is required (Younis et al., 2008). To accomplish this task, a rainfall–runoff (RR) model able to simulate the runoff formation process (i.e. estimation of losses) and also the hydrological routing along hillslopes and channels has to be set up. Besides the spatial discretization (lumped versus distributed models) and the process description (physically based versus conceptual models), RR models applied for operational flood forecasting can be subdivided in two main categories: continuous and event-based (Brocca et al., 2011; Paquet et al., 2013). On the one hand, continuous RR models simulate the temporal evolution of the soil moisture (SM) conditions of the catchment, thus being able to model the complex interaction between rainfall and SM which is necessary to properly predict flood hydrographs (Camici et al., 2011). However, the different processes (infiltration, percolation, evapotranspiration, interception) involved in the simulation of the SM temporal evolution may require a large number of parameters to be identified. This could easily introduce significant uncertainties into the model prediction and non-identifiable problems (Beven, 2006). Moreover, continuous models require long-term and uninterrupted time series for the input data (at least rainfall and temperature) and this could be a strong limitation in many regions worldwide, mainly if hourly observations are needed (Viviroli et al., 2009). On the other hand, event-based RR models need a reduced parameterization, and they are easy to be applied, even from users without extended hydrological expertise, and require low computational effort. For that, these types of models are very appealing, and frequently employed within operational flood forecasting systems (Berthet et al., 2009; Coustau et al., 2012). The major limitations of event-based models lie in the definition of the initial SM conditions that could be very different from one storm event to another (Tramblay et al., 2012; Coustau et al., 2012, Van Steenberghe and Willems, 2013). This issue is particularly challenging in regions characterized by strong seasonality of the climate as it occurs in Mediterranean basins (Aronica and Candela, 2004).

Nowadays, several SM data sources are available, also on a global scale. Specifically, SM information can be obtained from in situ and satellite sensors or from land surface models. From in situ observations, it is worth mentioning the International Soil Moisture Network (Dorigo et al., 2011), an international cooperation to establish and maintain a global in situ SM database that can be used for global analysis, e.g. the validation of the retrieval algorithms applied to remotely sensed observations. Moreover, several satellite SM

products are globally and freely available from active and passive microwave sensors, e.g. the Advanced SCATterometer (ASCAT) (Bartalis et al., 2007), the Advanced Microwave Scanning Radiometer for Earth observation (AMSR-E) (Owe et al., 2008), and the Microwave Imaging Radiometer with Aperture Synthesis (MIRAS) (Kerr et al., 2010). The accuracy and maturity of these satellite products have contributed to the implementation of a fully operational near-real-time (NRT) SM processing chain for ASCAT (Wagner et al., 2013) from the European Organisation for the Exploitation of Meteorological Satellites (EUMETSAT) since December 2008. Finally, modelled SM data obtained from numerical weather prediction systems operated by international meteorological centres (e.g. the ECMWF, European Centre for Medium-Range Weather Forecasting) are also available on a global scale. As an example, in the framework of the implementation of EUMETSAT's H-SAF project (Satellite Application Facility on Support to Operational Hydrology and Water Management), an ASCAT root zone SM profile product has been developed based on ASCAT surface SM data assimilation into the ECMWF Land Surface Data Assimilation System (De Rosnay et al., 2013). All these SM data sets, which are globally available, might be potentially used for the initialization of event-based RR models in different catchments and regions worldwide, even for poorly gauged areas.

Some studies attempted to relate the RR model initial conditions with different external indicators of SM estimated by in situ, satellite and modelled data (Brocca et al., 2009a, b, 2011a; Tramblay et al., 2010, 2011, 2012; Beck et al., 2009; Coustau et al., 2012; Graeff et al., 2012). In situ data were employed in many studies investigating the relationship between SM and runoff (e.g. Penna et al., 2011; Matgen et al., 2012; Graeff et al., 2012), thus indirectly determining their potential use for RR modelling. Brocca et al. (2009a, 2011b) and Beck et al. (2009) used the SM products from ASCAT and AMSR-E for RR modelling in Italy, Luxembourg and Australia. However, the main purpose of these studies was to investigate the relationship between modelled and observed antecedent wetness conditions, without explicitly building a model that incorporates external SM data.

In this study, a Simplified Continuous RR Model (SCRRM) is proposed by exploiting SM provided by ground, satellite and reanalysis data. This new approach offers the advantages of continuous models, with the difference that the temporal evolution of SM over a long-term period is assessed by using SM directly from external sources, thus avoiding simulating processes such as evapotranspiration, evaporation and groundwater flow. The “observed” SM time series (i.e. those provided by external sources) are used to set the initial saturation conditions of the catchment within an event-based model (Melone et al., 2001). In this way, the model is presented as “simplified continuous” since it links SM time series and a conceptual model for flood prediction. Doing so, the benefits of the event-based simulation, e.g. model

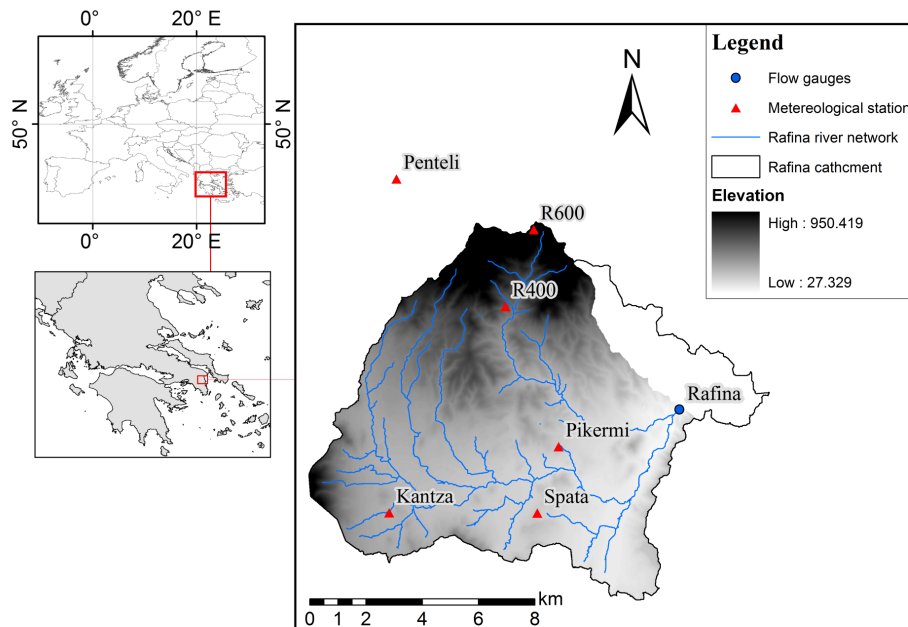


Fig. 1. Study area and locations of meteorological and runoff gauges stations used in the study.

parsimony, simplicity of use and no need for continuous rainfall and evapotranspiration data are achieved, while leaving the temporal evolution of the SM balance assessed by external sources. The results suggest that SCRRM can be particularly advantageous for operational purposes, especially in poorly gauged areas, where there is a lack of continuous meteorological data sets and a need to reduce the model parameterization, as well as to simplify the model structure in order to facilitate the model setup by end users (Montaldo et al., 2005; Coccia et al., 2009; Todini, 2009).

In the following, we first describe the study area (Rafina River basin, 109 km<sup>2</sup>, Greece), the meteorological data sets and the selected SM indicators to be used within SCRRM, then, besides the other models used (an event-based model and a standard continuous model) we describe the structure of SCRRM. In the results, we present both the comparison between in situ and the selected SM indicators used and the performance obtained from SCRRM compared to those obtained from the other models used in this study.

## 2 Study area and data sets

### 2.1 Rafina catchment

The study area is the Rafina river basin upstream from the Rafina gauged section (109 km<sup>2</sup>).

This is a periurban area in the greater southeast Mesogeia region in eastern Attica, Greece (Fig. 1). The area geographically extends from east of the Hymettus mountains to the coastline of Evoikos Gulf. The mean altitude of the region is 227 m a.s.l. (with the minimum altitude being 0 and

the maximum 909 m a.s.l.). Ground slope ranges from 0 to 37.8 % with a mean value of 7.5 %. Increased slopes and irregular terrain exist mainly at the upstream parts of the area. Attica has a typical subtropical Mediterranean climate, with prolonged hot and dry summers succeeded by considerably mild and wet winters. The mean annual precipitation is approximately 400 mm, while snowfall is rare. Drought periods usually begin in May and last until October. The daily mean temperature ranges between 27 °C during the summer months and 11 °C during the winter months (Papathanasiou et al., 2013).

Geologically, the study area is part of the Attico–Cycladic Massif. The Neogene and Quaternary deposits fill up the degradations and tectonic grabens of the east Attica Basin and consist of alluvial deposits and Mio-pliocene lacustrine and terrestrial deposits, conglomerates, sand, marls, clayey material, marly limestones and clays (Jacobshagen, 1986). Rafina catchment is covered by different and often conflicting land uses. More specifically, it includes forests (~30 %), arable soils and grasslands (~50 %) mainly located upstream, and urban cells (~20 %) located downstream (Alonistioti, 2011). The study area is under constantly increasing urbanization, its northern part is forested with flammable material and sediment load transfer and soil erodibility are intense (Papathanasiou et al., 2009, 2012).

### 2.2 Hydro-meteorological data

The Laboratory of Hydrology and Water Resources Management of the School of Civil Engineers of the National Technical University of Athens (NTUA) operates the Hydrological Observatory of Athens (HOA), a dense monitoring

**Table 1.** Main characteristics of the selected rainfall–runoff events.  $\theta_e$  represents the in situ relative soil moisture at the beginning of each selected event.

Dates	Maximum discharge [m <sup>3</sup> s <sup>-1</sup> ]	Lag time [h]	Total rainfall [mm]	Direct runoff [mm]	Runoff coefficient [–]	Duration [h]	$\theta_e$ (in situ) [–]
21 Mar 2009	3.5	2.9	14.6	0.3	0.02	22	0.43
25 Oct 2009	10.7	3.8	27.8	1.3	0.04	25	0.22
26 Oct 2009	4.0	3	12.3	0.3	0.03	19	0.33
3 Nov 2009	4.4	4.1	21.0	0.3	0.02	24	0.38
11 Dec 2009	12.4	4.0	34.8	1.8	0.05	24	0.52
2 Jan 2011	3.3	5.5	18.5	0.2	0.01	30	0.40
3 Feb 2011	39.5	5.4	87.5	10.7	0.12	40	0.49
18 Feb 2011	7.5	4.1	12.5	0.6	0.04	19	0.42
24 Feb 2011	18.8	5.2	47.2	5.8	0.12	37	0.58
7 Mar 2011	5.5	5.0	17.0	0.8	0.05	26	0.48
17 Apr 2011	2.9	4.1	12.7	0.2	0.02	23	0.22
26 Apr 2011	5.0	5.8	28.6	1.0	0.04	37	0.27
12 Jun 2011	9.3	4.4	28.2	1.1	0.04	19	0.31
19 Dec 2011	6.7	8.5	18.8	1.0	0.05	35	0.09
10 Jan 2012	2.5	9.2	15.0	0.7	0.05	39	0.34
21 Feb 2013	137.2	4.6	107.1	14.5	0.13	38	0.67

hydrometeorological network in the greater Athens area. HOA is the evolution from the METEONET network that has been operating since 2005 and consists of several active meteorological and active runoff measuring stations, properly located in the area. The stations are equipped with sensors that measure with 10-minute temporal resolution environmental parameters of hydrometeorological interest. Detailed information of the instruments can be found at <https://hoa.ntua.gr>. Parallel to that, the National Observatory of Athens (NOA) also operates a dense meteorological network in the greater Athens area also recording valuable meteorological information in 10-minute temporal resolution (NOA, 2012, [www.meteo.gr/meteosearch](http://www.meteo.gr/meteosearch)).

For this study, rainfall data were extracted from six rainfall stations of HOA and NOA networks (Fig. 1). As the area extends over only 109 km<sup>2</sup>, this number can be considered appropriate to take into account the spatial variability of rainfall in the catchment. Temperatures were retrieved from the thermometer located in Pikermi. Measured stages at the Rafina gauged site were used to develop updated rating curves and thus evaluate discharges at the locations of the gauge. The period of analysis ranges from 12 March 2009 to 28 February 2013. The main data analysis included a quality control to remove inconsistent values, and aggregation operations to produce hourly based time-step temporal resolution.

The rating curve developed for the Rafina site was based on velocity measurements carried out from the end of 2009 for low to medium water levels. It was assumed that (i) the instrument always worked under ideal conditions (as verified from the discharge data set), and, (ii) the geometry of gauging the cross-section was stable in time, even though some

changes may have occurred during high flood events due to erosion, sediment transport and deposition. Based on that, it is likely that the rating curve contains inherent uncertainties, which may affect the estimated runoff coefficients. However, this is the best data available. Moreover, this issue extends beyond the main goal of the paper and it is expected to not affect the results since the same errors are present for all the investigated configurations.

The mean areal rainfall was calculated by the Inverse Distance Weighting Method. Direct runoff was evaluated as in Melone et al. (2002) by using an appropriate baseflow separation technique.

For this study, rainfall events were extracted by selecting those with a continuous rainfall characterized by a total rainfall larger than 10 mm, and no rainfall in the preceding day. Eventually, 16 rainfall–runoff events were analysed with cumulated rainfall and runoff coefficients ranging from 12.3 to 107 mm and from 0.01 to 0.13 mm, respectively (Table 1, Fig. 2). For the most significant event – occurring on 21 February 2013 – the maximum recorded peak discharge was 137.2 m<sup>3</sup> s<sup>-1</sup>, while for the least intense event – recorded on 10 January 2012 – a peak discharge of 2.5 m<sup>3</sup> s<sup>-1</sup> was observed.

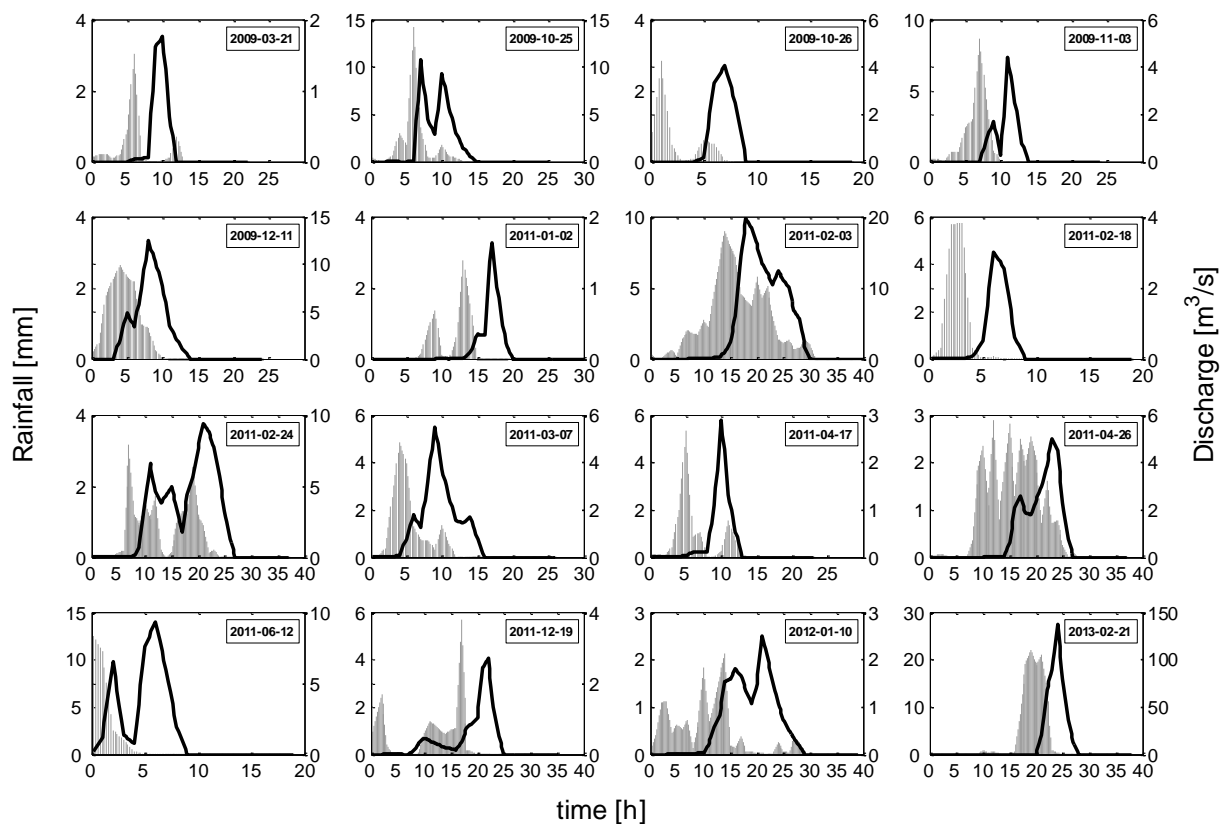
### 2.3 Soil moisture indicators

Four different SM indicators were selected covering the period 2009–2013. In particular, the selected indicators were (i) ground SM obtained from measurements carried out at a depth of 25 cm at Pikermi station (see Fig. 1); (ii) ASCAT-derived SM product (Wagner et al., 1999); (iii) AMSR-E-derived SM product (Owe et al., 2001); and (iv) ERA-Land

**Table 2.** Soil moisture indicators used in this study.

Soil moisture indicator	LAT <sup>1</sup>	LON <sup>2</sup>	Spatial resolution	Temporal resolution	Band [GHz]	Retrieval algorithm	Depth [cm]
Remote sensing soil moisture							
In situ	37.9919	23.9194	N/A	10 min	N/A	N/A	25
ERA-Land	38.2406	23.3417	80 km	1 day	N/A	N/A	0–7, 8–28, 28–100
ASCAT	37.9326	23.8872	25 km	≈ 1 day	5.25 (C band)	WARP 5.5 v1.1	0–3
AMSR-E	37.875	23.875	56 km	≈ 1 day	6.90 (C band)	LPRM	0–3

<sup>1</sup> LAT = latitude, <sup>2</sup> LON = longitude.

**Fig. 2.** Observed rainfall and discharge for the selected flood events.

SM product from the ECMWF (Balsamo et al., 2012). Given the catchment extension (about 100 km<sup>2</sup>) and the spatial resolution of the satellite (ASCAT and AMSR-E) and modelled (ERA-Land), products-only data of the pixel nearest to the centroid of the catchment were used and the corresponding relative surface SM time series were retrieved from the appropriate database, that is, no spatial aggregation was carried out. Moreover, pixels near the sea or in mountain areas were

not considered. Table 2 summarizes the main characteristics of the selected soil moisture indicators.

### 2.3.1 In situ soil moisture data

SM data selected for this study were collected at Pikermi station (see Fig. 1). The station was installed at the beginning of 2009 from HOA and can be considered representative of the catchment in terms of land cover characteristics and soil

properties. The sensor is a water content reflectometer Campbell Scientific CS616-L, installed in clay soil measuring volumetric SM at a depth of 25 cm through frequency domain reflectometry (FDR) technique. The measurements refer to the period from March 2009 to February 2013 and have a 10-min temporal resolution. Since this station is the only one measuring SM within the catchment boundaries, it was assumed that measurements reflect the temporal variability of the soil moisture in the catchment.

### 2.3.2 Satellite soil moisture: ASCAT and AMSR-E

Remote sensing soil moisture products were retrieved from one active (ASCAT) and one passive sensor (AMSR-E). Both of them allow the retrieval of surface SM (SM representative of the first centimetres of the soil). In this section, after a description of each product we present the method for calculating the root zone SM from the remotely sensed surface SM.

ASCAT is a real-aperture radar instrument on board the MetOp satellite. It measures radar backscatter at C-band (5.255 GHz) in VV polarization. Its spatial resolution is 25 km then re-sampled at 12.5 km. In the study area, measurements are available at least once a day (07:00–08:00 UTC in descending orbit and/or 18:00–20:00 UTC in ascending orbit). The surface SM product (equivalent to a depth of 2–3 cm of the soil) is calculated from the backscatter measurements through a time-series-based change detection approach previously used for the ERS-1/2 by Wagner et al. (1999). The SM is derived by selecting the historical lowest and highest backscatter measurement to which is respectively assigned 0% (dry), and 100% (wet) reference. The ASCAT surface SM product used for this study covers the period 2009–2013 considering both ascending and descending overpass.

The AMSR-E sensor is the instrument on board the NASA Aqua satellite. It is a passive microwave radiometer measuring at 6.9 GHz (C-Band) and five higher frequencies. The sensor provided measurements from May 2002 to October 2011 with daily ascending and descending overpasses and a spatial resolution of about 56 km (resampled at 0.25°). In this study the Land Parameter Retrieval Model (LPRM) was used as a retrieval algorithm (Owe et al., 2001, 2008) on data for the period 2009–2011 (AMSR-E sensor stopped working in October 2011). The LPRM was developed by the University of Amsterdam (VU) in collaboration with NASA, and was successfully tested over many sites in Europe (e.g. Brocca et al., 2011c).

Remotely sensed products provide knowledge of soil moisture for a very thin surface layer (ca. 0–5 cm), however this is not sufficient for hydrological applications concerning RR transformation. Indeed, root-zone SM data are the main control parameters on the catchment response to a given storm event (Brocca et al. 2010). To obtain the root-zone SM product (SWI; soil water index) from the satellite-based

surface observations, the semi-empirical approach developed by Wagner et al. (1999) was adopted. The recursive formulation of the method relies on (Albergel et al., 2012)

$$\text{SWI}(t_n) = \text{SWI}(t_{n-1}) + K_n [m_s(t_n) - \text{SWI}(t_n - 1)], \quad (1)$$

where  $m_s(t_n)$  is the surface SM observed by the satellite sensor,  $\text{SWI}_{t_n}$  is the soil wetness index representing the profile-averaged saturation degree and time  $t_n$  is the acquisition time of  $m_s(t_n)$ . The gain  $K_n$  at time  $t_n$  is given by (in a recursive form)

$$K_n = \frac{K_{n-1}}{K_{n-1} + e^{-\left(\frac{t_n - t_{n-1}}{T}\right)}}, \quad (2)$$

where  $T$  is the characteristic time length and represents the timescale of SM variation to obtain the SWI. For the initialization of this filter,  $K_1$  and  $\text{SWI}_1$  were set to 1 and  $m_s(t_1)$ , respectively (Albergel et al., 2012). The approach is also known as exponential filter. The reader can find a more detailed description of this approach in Wagner et al. (1999) and Albergel et al. (2012).

### 2.3.3 ERA-Land soil moisture product (ECMWF)

The ECMWF provides medium-range global forecasts for some environmental variables that include soil temperature, evaporation and SM. ERA-Land stems from ERA-Interim (Dee et al., 2011) for the atmosphere, however its land surface components were completely revised to match ECMWF most recent developments, e.g. the HTESSSEL land surface model (Balsamo et al., 2012, 2013; Albergel et al., 2012). HTESSSEL uses the dominant soil texture class from the FAO (FAO, 2003) data set for each grid point. Albergel et al. (2012) provide further details concerning the integration of the soil types into the ECMWF product.

The ERA-Land SM product used in this study ranges from 1 January 1979 to 31 December 2012 and is available for 00:00, 06:00, 12:00 and 18:00 UTC with a spatial resolution of about 80 km (T255) considering four layers of soil (0–7, 7–28, 28–100 and 100–289 cm). In this study, the SM values relative to the first two layers (0–28 cm) were compared with in situ data, while the values of the first three soil layers (0–100 cm) were used for the RR transformation.

## 3 Flood modelling

In the following, we present a description of all the models used in this paper: an event-based model “Modello Idrologico Semidistribuito” (MISD, Melone et al., 2001), a continuous model (MISDc model “Modello Idrologico Semidistribuito in Continuo”; Brocca et al., 2011b) and the “Simplified continuous rainfall–runoff model”, SCRRM, which might be viewed as an evolution of the previous two. It has to be pointed out that external sources of SM, i.e. from satellite,

from in situ or from ERA-Land products, are used here as a proxy of observations for the assessment of the wetness state of the catchment (Beck et al., 2009; Brocca et al., 2009a, b).

### 3.1 Event-based rainfall–runoff model: MISD

The event-based RR model considered in this study is the lumped version of MISD.

#### 3.1.1 Losses

MISD employs the Soil Conservation Service–Curve Number (SCS-CN) method for estimation of losses. The choice of the SCS-CN method is due to its wide use since 1980 (Kim and Lee, 2008) and its simplicity. In particular, for a storm event the partitioning of rainfall into runoff using the SCS-CN method is based on the following equation:

$$Q = \frac{(P - F_a)^2}{P - F_a + S} \quad P \geq F_a, \quad (3)$$

where  $F_a$  is the initial abstraction,  $S$  is the soil potential maximum retention estimated as a function of land use, hydrological soil group, and total precipitation of the previous five days;  $Q$  is the direct runoff depth; and  $P$  is the rainfall depth. The quantity  $F_a$  is considered linearly dependent on  $S$  by

$$F_a = \lambda S, \quad (4)$$

where  $\lambda$  is the initial abstraction coefficient (Soil Conservation Service, 1993). Equation (1) is extended for the time evolution of the effective rainfall rate,  $e(t)$ , within a given storm as (Melone et al., 2001)

$$e(t) = \frac{dQ}{dt} = \frac{p(t)(P(t) - F_a)(P(t) - F_a + 2S)}{(P(t) - F_a + S)^2} \quad P(t) \geq F_a, \quad (5)$$

where  $p$  is the rainfall rate and  $P(t) = \int_0^t p(\tau) d\tau$ .

#### 3.1.2 Routing

Once the time evolution of effective rainfall is computed, the routing to the outlet of the catchment is obtained from the convolution of the rainfall excess and the Geomorphological Instantaneous Unit Hydrograph (GIUH), such as proposed by Gupta and Waymire (1980). In the model, the lag time is evaluated through the relationship proposed by Melone et al. (2002)

$$L = \eta 1.19 A^{0.33} \quad (6)$$

with  $L$  being the lag time (h),  $A$  the area of the catchment ( $\text{km}^2$ ), and  $\eta$  a parameter to be calibrated (Moramarco et al., 2005).

To sum up, the parameters of MISD model are the lag-area relationship parameter,  $\eta$ , the initial abstraction coefficient,  $\lambda$ , and the soil potential maximum retention,  $S$ .

### 3.2 Continuous rainfall–runoff Model: MISDc

MISDc is a continuous rainfall–runoff model successfully applied to the Tiber River for flood prediction and operational purposes (<http://www.cfumbria.it/MISDc/>). The lumped version of MISDc model used in this study couples a soil water balance model (SWB, Brocca et al., 2008) to simulate the soil moisture temporal pattern, and MISD as an event-based RR model (MISD) for transferring the excess of the rainfall to the outlet section of the catchment. The two models are linked through an experimentally derived linear relationship between  $S$  of the SCS method, and the relative soil moisture at the beginning of the event. The MISDc structure was established by analysing the hydrologic response of a small experimental catchment equipped with an automatic soil moisture monitoring system (Brocca et al., 2009a).

The SWB model is a simple bucket model representing the main processes needed for SM simulation: infiltration, percolation and evapotranspiration. More specifically, the processes are represented for infiltration through the Green–Ampt equation, for drainage by a gravity-driven non-linear relationship and for actual evapotranspiration by a linear relationship with the potential evapotranspiration, calculated through a modified Blaney and Criddle method. The reader is referred to Brocca et al. (2008, 2013) for the full description of the model equations. The SM simulated by the SWB is used to calculate the parameter  $S$  of the SCS method by means of an experimentally derived relationship between  $S$  and SM (Brocca et al. 2009a, b):

$$S = a(1 - \theta_e), \quad (7)$$

where  $\theta_e$  is the modelled relative soil moisture at the beginning of the event and  $a$  is a parameter to be estimated. Once the  $S$  parameter is estimated, MISD is used for simulating the flood hydrograph (see Sect. 3.1.2).

MISDc requires as input rainfall and air temperature data; the model output is both the direct runoff hydrograph in correspondence to flood events and the catchment average relative soil moisture. In MISDc, seven parameters have to be estimated, namely the maximum water capacity of the soil layer  $W_{\max}$ , the saturated hydraulic conductivity  $K_s$ , the ratio between the wetting front soil suction head and the soil layer thickness  $\psi/L$ , the pore size distribution index  $m$ , the correction coefficient for potential evapotranspiration  $b$ , the lag-area relationship parameter  $\eta$ , the initial abstraction coefficient  $\lambda$  and the parameter  $a$  of the  $S - \theta_e$  relationship (see Brocca et al., 2011b for further details).

### 3.3 Simplified continuous rainfall–runoff model: SCRRM

Continuous RR models like MISDc simulate SM to take the variability of the wetness conditions prior to a rainfall event into account. They require as input, continuously measured temperature or evapotranspiration and rainfall measurements



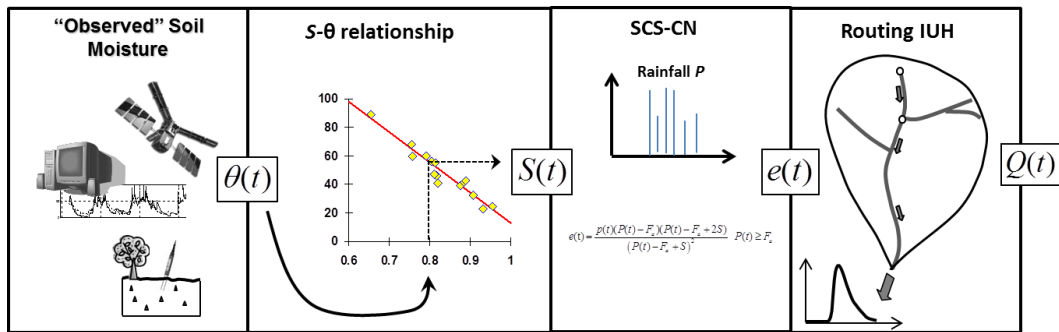


Fig. 3. Structure of the Simplified Continuous RR Model (SCRRM).

to model SM and produce the discharge hydrograph. If the SM at the beginning of an event is provided by an external indicator, e.g. in situ or globally available SM observations such as satellite and model-based reanalysis products, a new methodology approach can be derived, as schematized in Fig. 3. SCRRM is similar to a continuous model but with some significant differences. In particular, the temporal evolution of the soil wetness conditions of the catchment is not modelled from rainfall and temperature data like in MISDc but it is integrated directly into the model from SM observations (i.e. SWB is replaced in SCRRM by SM observations). Based on that, SCRRM uses SM indicators provided by external sources to infer the value of  $S$  parameter for runoff determination. Like in MISDc, the model exploits the observed linear behaviour between the wetness state of the soil and the parameter  $S$  of the SCS method (Brocca et al., 2009a, b) via Eq. (7). Once  $S$  is known, the rainfall excess is calculated and transferred to the outlet section as in MISD (see Sect. 3.1.2).

In synthesis, the SCRRM proposed in this paper uses the SM and the event rainfall data as sole inputs to simulate hourly flood hydrographs. Since the SM is provided by an external indicator, the  $S - \theta_e$  relationship becomes a model relation embedded in the model structure and it is used to estimate the value of  $S$  for the analysed events. The calibration of the model involves the following three parameters: the coefficient of initial abstractions  $\lambda$ , the parameter  $a$  of the  $S - \theta_e$  relationship, and the parameter  $\eta$  of the lag time–area relationship.

It has to be noted that the soil layer depth that controls the RR transformation is usually larger than few centimetres. As a result, the application of SCRRM with satellite products is taken into account by including as an additional parameter of SCRRM the characteristic time length  $T$  of the exponential filter of Wagner et al. (1999) described in Sect. 2.3.2, or by considering the appropriate soil layer depth for the ERA-Land reanalysis product (see Sect. 2.3.3).

For this study, a lumped model was employed, even though the same concept can be easily applied to spatially distributed models. Finally, it is highlighted that the different components of the model (i.e. SCS-CN and GIUH) can be

changed while keeping the general model structure shown in Fig. 3.

### 3.4 Performance scores and data preparation

#### 3.4.1 Performance in flood modelling

For assessing the performance of the proposed models, the Nash–Sutcliffe efficiency coefficient (NS) was used to evaluate the agreement between the simulated and observed hydrographs for each one of the selected flood events (Table 1):

$$NS = 1 - \frac{\sum_{t=1}^{T_{ev}} (Q_{obs} - Q_{sim})^2}{\sum_{t=1}^{T_{ev}} (Q_{obs} - \bar{Q}_{obs})^2}, \quad (8)$$

where  $Q_{obs}$  and  $Q_{sim}$  are the observed and simulated discharges at time  $t$  respectively,  $\bar{Q}_{obs}$  is the mean value of the observed discharge during the event and  $T_{ev}$  is the event duration. In particular, as an objective function the mean Nash–Sutcliffe,  $\overline{NS}$  was calculated as

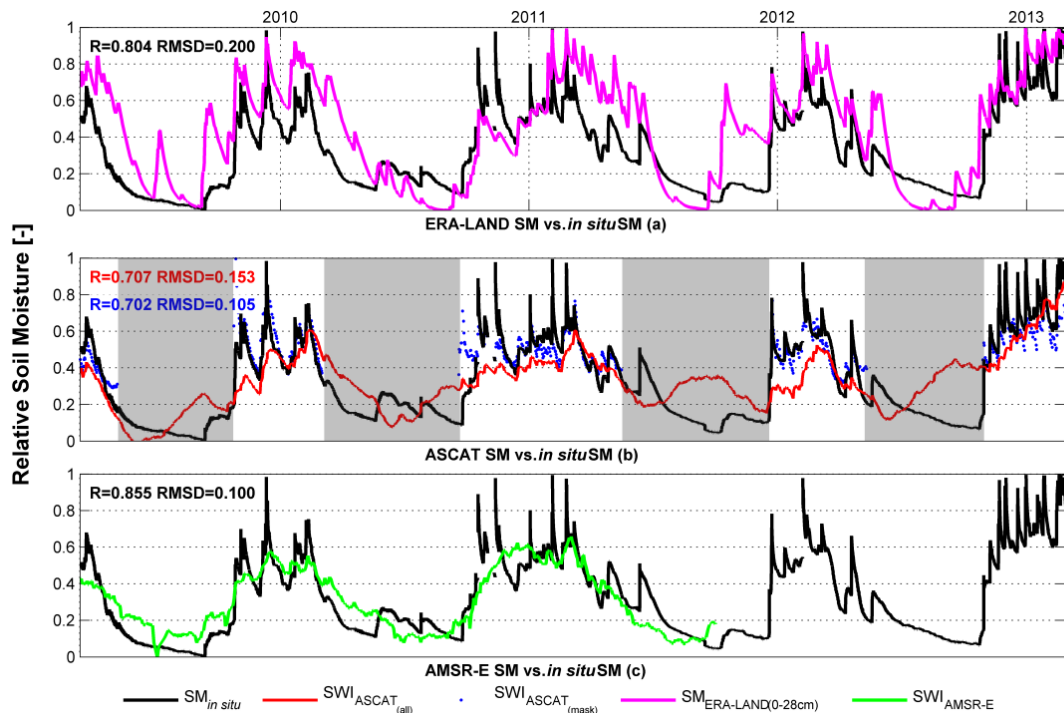
$$\overline{NS} = \frac{\sum_{j=1}^{N_{ev}} NS_j}{N_{ev}} \quad (9)$$

and an hourly time step was used in the simulations. In Eq. (9)  $N_{ev}$  is the number of the events considered, whereas index  $j$  refers to the event of Table 1. For model calibration, a standard gradient-based automatic optimization method (“fmincon” function in MATLAB<sup>®</sup>) was used. In addition, to evaluate the performance of the model in reproducing flood events, the percentage error on peak discharge

$$E_{Q_p} = 100 \frac{\max(Q_{obs}) - \max(Q_{sim})}{\max(Q_{obs})} \quad (10)$$

and the percentage error on direct runoff volume





**Fig. 4.** Relative soil moisture temporal pattern for the different soil moisture indicators selected for this study (period March 2009–February 2013). Comparison between the relative values of the ground soil moisture ( $SM_{in\ situ}$ ) obtained from measurements carried out at a depth of 25 cm (number of data,  $N = 34\,972$  hourly data) with (a) ERA-Land soil moisture product from ECMWF ( $SM_{ERA-LAND(0-28cm)}$ ) obtained from the weighted mean of the 0–7 and 7–28 cm products ( $N = 1450$ ); (b) SWI ASCAT-derived soil moisture product considering the whole data set ( $SWI_{ASCAT(all)}$ ,  $N = 1326$ ,  $T = 147.5$  days) and the data set with data in shaded windows removed ( $SWI_{ASCAT(mask)}$ ,  $N = 631$ ,  $T = 9.5$  days); (c) AMSR-E-derived soil moisture product ( $SWI_{AMSR-E}$ ,  $T = 7.5$  days;  $N = 584$ ).  $R$  refers to the correlation coefficient, RMSD are the root mean squared differences.

$$E_V = 100 \frac{\sum_t^{T_{ev}} Q_{obs} - \sum_t^{T_{ev}} Q_{sim}}{\sum_t^{T_{ev}} Q_{obs}} \quad (11)$$

were both evaluated for each single event and as a mean of all the selected events.

### 3.4.2 SM data preparation

ASCAT SM is given in percentage (index between 0 and 100) while AMSR-E, in situ and ERA-Land data are in  $m^3\ m^{-3}$ , therefore, to allow for a robust comparison and a homogeneous parameter calibration, all soil moisture data sets (in situ, remotely sensed and modelled) were rescaled between [0, 1] using their own maximum and minimum values over the period of analysis. The evaluation of satellite and re-analysis SM data against in situ observations was carried out using the correlation coefficient,  $R$ , and the root mean squared differences (RMSD) as in Brocca et al. (2011c).

## 4 Results and discussions

### 4.1 Assessment of the soil moisture indicators

Figure 4 compares the relative values of the ground SM ( $SM_{in\ situ}$ ) obtained from measurements carried out at a depth of 25 cm with (i) ERA-Land SM product from ECMWF, obtained from the weighted mean of the 0–7 and 7–28 cm soil layers, (Fig. 4a); (ii) ASCAT-derived SWI product (Fig. 4b); and (iii) AMSR-E-derived SWI product (Fig. 4c). Data refer to the period March 2009–February 2013.

ERA-Land (Fig. 4a), is satisfactorily correlated with in situ data ( $R = 0.804$ ) but it shows higher RMSD with respect to the other products, ( $RMSD = 0.200$ ) likely due to its coarse spatial resolution ( $\sim 80$  km) and to errors in the meteorological forcing (mainly precipitation) that are obtained directly from the global model and not from ground observations.

For remotely sensed SM indicators (Figs. 4b, c), the evaluation of SWI data was obtained from maximizing the correlation coefficient  $R$  by varying the  $T$  parameter of the exponential filter from 1 to 200 days at step of 0.5 days. For ASCAT, an anomalous behaviour was observed in dry periods which involved an anomalous increase without the occurrence of

rainfall (see shaded areas in Fig. 3b). Such behaviour has been observed in other arid and semi-arid regions as well, and it is assumed to be due to volume-scattering effects from dry sub-surface soil layers (Wagner et al., 2013). Given that to date there is no solution to this problem, the analysis was carried out twice: once by considering the whole data set ( $SWI_{ASCAT(all)}$ ) in red in Fig. 4b) and once with removed anomalous data ( $SWI_{ASCAT(mask)}$ ) in blue in Fig. 4b). In the first case the parameter  $T$  is 147.5 days with a correlation coefficient  $R=0.707$  and  $RMSD=0.153$ , whereas without anomalous data  $T=9.5$  days,  $R=0.702$  and  $RMSD=0.105$ . While the first result conflicts with the results of Wagner et al. (1999), Ceballos et al. (2005) and Brocca et al. (2011c), the second one is more similar to the studies of Wagner et al. (1999) which found  $T$  equal to 20 days for an area of about 600 000 km<sup>2</sup> in the Ukraine. Therefore, in what follows, only the ASCAT data set without anomalous data is considered. For AMSR-E product (Fig. 4c),  $T=7.5$  days with  $R=0.855$  and  $RMSD=0.100$ . In this case, the lower values of  $T$  are due to the smoother behaviour of the AMSR-E SM product with respect to ASCAT SM.

Overall, both ERA-Land and satellite products show good correlation ( $R > 0.7$ ) with in situ data. For satellite products, apart from the scattering problem described for ASCAT in dry periods, such good correlations exclude problems due to vegetation cover (Owe et al., 2008; Wagner et al., 2013)

## 4.2 Model performances

In the following, the performances of SCRRM and MISDc are presented for a simple split sample test procedure. The model parameters were calibrated on the first eight flood events (from 24 March 2009 to 18 February 2011), whereas the remainder of them (from 24 February 2011 to 21 February 2013) was used in validation. The results obtained with MISDc are used as a baseline to assess the consistency and the reliability of SCRRM. Table 3 summarizes the value of the parameters obtained for all the models used in this study.

### 4.2.1 MISD using “observed” initial conditions

As a first analysis, to assess the optimal antecedent wetness conditions for the selected events, the soil potential maximum retention,  $S$ , was expressed by using observed rainfall and direct runoff depth by inverting Eq. (3) as in Beck et al. (2009) and Brocca et al. (2009a). That is,  $S$  is the value of the potential maximum retention that reproduces the direct runoff volume, henceforth it is indicated as “observed  $S$ ”,  $S_{obs}$  (Brocca et al., 2011b).

The calibration of the model was carried out by maximizing  $\overline{NS}$  for estimating at the same time as  $\lambda$ ,  $\eta$  and  $S_{obs}$ . The value of the parameters obtained in calibration is  $\lambda=0.0015$  and  $\eta=0.669$  (Table 3), while  $S_{obs}$  ranges from 254 to 1188 mm (see Table 4), indicating a wide range of initial conditions at the beginning of the events. In calibration,

**Table 3.** Calibrated parameters of the models used in the study obtained from the maximization of the mean NS among all the events.  $MISD_{S_{obs}} = MISD$  using “observed” initial condition  $S_{obs}$  (obtained from inverting Eq. 3);  $MISD_{S_{cost}} = MISD$  model using constant observed initial conditions (equal to the mean of  $S_{obs}$  among all the selected events); in situ = model SCRRM using  $S$  estimated from in situ SM measurements; ASCAT = SCRRM using  $S$  estimated from ASCAT SM measurements; AMSR-E = SCRRM using  $S$  estimated from AMSR-E SM measurements; ERA-Land = SCRRM using  $S$  estimated from ERA-Land SM measurements; MISDc = MISDc model (only  $a$ ,  $\eta$  and  $\lambda$  are shown).

Model	$a$ [mm <sup>-1</sup> ]	$\eta$	$\lambda$	$T$
$MISD_{S_{obs}}$	–	0.669	0.0015	–
$MISD_{S_{cost}}$	–	0.349	0.0010	–
In situ	1254	0.643	0.0010	–
ASCAT	1250	0.645	0.0010	168.3
AMSR-E	1250	0.647	0.0021	120.0
ERA-Land	1057	0.634	0.0023	–
MISDc	1184	0.650	0.0010	–

NS values range from 0.35 to 0.84, with a mean value of 0.61, indicating an acceptable fit of the model. Due to the low reliability of the estimated rating curve as well as the weak response of the catchment to rainfall inputs in terms of runoff, for low-magnitude flood events, the peak discharges are usually overestimated. Indeed, a successful application of whatever rainfall–runoff model to this area may be very difficult since the contribution of the rainfall to runoff is very small (see the low runoff coefficients in Table 1). However, these have to be considered the best possible results that can be attained for this catchment since the initial conditions are derived from observed values of rainfall and discharge.

Theoretically, in validation,  $S_{obs}$  cannot be calibrated since discharge values are unknown. However, aiming only to have a baseline for comparing the performances obtained from the other models, the parameters obtained in calibration were used to predict the last eight events assuming  $S_{obs}$  as known (i.e. best initial condition estimation). The results are presented in Table 4 and show performances very similar to those obtained in calibration.

### 4.2.2 MISD using constant initial conditions

If an event-based RR model is run without the knowledge of the initial conditions, a constant value for all the events has to be selected. To highlight this situation, in the calibration period the initial condition for all the selected flood events was set equal to the mean of  $S_{obs}$  (659.1 mm in Table 4 for the flood events of the calibration period). The values of  $\lambda$  and  $\eta$  parameters obtained in calibration are given in Table 3. In this case, mean and median NS is equal to 0.04 and 0.13, respectively (Table 4). In validation, mean  $S_{obs}$  was set equal to 451.04 mm (see Table 4, mean of the  $S_{obs}$  of the flood

**Table 4.** Performance of the models used in the study in terms of Nash–Sutcliffe coefficient NS.  $S_{\text{obs}}$  refers to “observed” soil potential maximum retention determined by using observed rainfall and direct runoff depth;  $\text{MISD}_{S_{\text{obs}}} = \text{NS}$  obtained from MISD using  $S_{\text{obs}}$ ;  $\text{MISD}_{S_{\text{cost}}} = \text{NS}$  obtained from MISD using constant observed initial conditions (equal to the mean of  $S_{\text{obs}}$  among all the selected events); in situ = NS obtained from SCRRM using  $S$  estimated from in situ SM measurements; ASCAT = NS obtained from SCRRM using  $S$  estimated from ASCAT SM measurements; ERA-LAND = NS obtained from SCRRM using  $S$  estimated from ERA-Land SM measurements; AMSR-E = NS obtained from SCRRM using  $S$  estimated from AMSR-E SM measurements; MISDc = NS obtained from MISDc. Numbers in brackets in the validation period refer to the statistics calculated for the flood events for the period 24 February 2011 to 12 June 2011 (the only available for ASCAT).

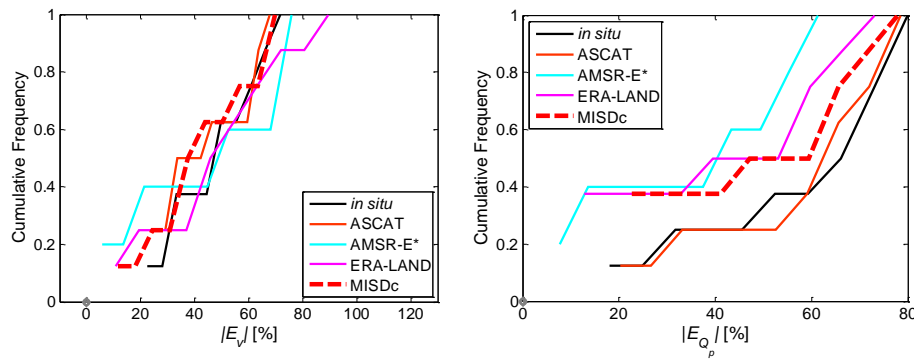
Dates	$S_{\text{obs}}$ (mm)	$\text{MISD}_{S_{\text{obs}}}$	$\text{MISD}_{S_{\text{cost}}}$	In situ	ASCAT	AMSR-E	ERA-Land	MISDc
Calibration								
21 Mar 2009	653.3	0.40	−0.04	0.40	0.40	0.43	0.43	0.28
25 Oct 2009	543.3	0.57	0.64	0.40	0.36	0.34	0.43	0.37
26 Oct 2009	372.0	0.35	0.20	0.12	0.06	0.06	0.10	0.08
3 Nov 2009	1051.2	0.68	−0.83	0.60	0.67	0.68	0.56	0.68
11 Dec 2009	599.1	0.84	0.84	0.87	0.86	0.73	0.87	0.76
2 Jan 2011	1187.5	0.64	−1.30	0.48	0.49	0.29	0.64	0.52
3 Feb 2011	612.5	0.80	0.79	0.80	0.79	0.82	0.78	0.79
18 Feb 2011	253.9	0.59	0.06	0.19	0.23	0.38	0.33	0.35
Mean	659.0	0.61	0.04	0.48	0.48	0.47	0.52	0.48
Median	605.8	0.62	0.13	0.44	0.45	0.41	0.50	0.44
Max	1187.5	0.84	0.84	0.87	0.86	0.82	0.87	0.79
Min	253.9	0.35	−1.30	0.12	0.06	0.06	0.10	0.08
Validation								
24 Feb 2011	330.4	0.84	0.81	0.69	0.68	0.82	0.82	0.77
7 Mar 2011	323.8	0.81	0.45	0.44	0.63	0.80	0.81	0.79
17 Apr 2011	646.5	0.23	−0.54	0.16	0.18	0.04	0.16	0.20
26 Apr 2011	718.9	0.77	0.49	0.70	0.64	0.42	0.67	0.74
12 Jun 2011	329.9	0.53	−0.04	0.18	0.08	0.32	0.34	0.27
19 Dec 2011	302.1	0.53	−0.08	0.09	0.23	–	0.18	0.11
10 Jan 2012	289.9	0.46	0.15	0.09	0.13	–	0.11	0.03
21 Feb 2013	670.0	0.85	0.39	0.82	0.83	–	0.49	0.85
Mean	451.04	0.63 (0.63)	0.20 (0.23)	0.40 (0.43)	0.43 (0.48)	(0.47)	0.45 (0.56)	0.47 (0.55)
Median	330.15	0.65 (0.76)	0.27 (0.45)	0.31 (0.44)	0.43 (0.62)	(0.41)	0.42 (0.67)	0.51 (0.74)
Max	718.9	0.85 (0.84)	0.81 (0.81)	0.82 (0.70)	0.83 (0.68)	(0.82)	0.82 (0.82)	0.85 (0.79)
Min	289.9	0.23 (0.23)	−0.54 (−0.54)	0.09 (0.16)	0.08 (0.08)	(0.04)	0.11 (0.16)	0.03 (0.20)

events of the validation period) obtaining results similar to those achieved in calibration with poor prediction skills of the model. Overall, these results highlight the importance of a correct selection of the initial condition for flood modelling in this catchment.

#### 4.2.3 SCRRM and MISDc models

In the following, the performances of SCRRM and MISDc models are presented and compared for the selected calibration and validation periods. In particular, SCRRM was run by considering in situ, ASCAT, AMSR-E and ERA-Land SM data sets with the value of the SM at the beginning of the event picked up directly from time series by selecting the SM value corresponding to the preceding date nearest to

the event date. For ASCAT, only the results with removed anomalous data are presented. For AMSR-E, the validation was carried out for only five events (from 24 February 2011 to 12 June 2011, the results of this set of events for the other SM indicators are presented in brackets) since for the last three events AMSR-E SM was not available (AMSR-E sensor ceased working on October 2011). The calibration of both models was carried out by maximizing  $\overline{\text{NS}}$ . The calibrated values of the parameters are presented in Table 3. Note that for the two satellite products the characteristic time length  $T$  was also calibrated as specified in Sect. 3.3. For MISDc, only the parameters in common with SCRRM are shown. Table 4 summarizes the results in terms of NS for calibration and validation.



**Fig. 5.** Cumulative frequency of absolute error in volume  $|E_V|$  (a), and absolute error in peak discharge,  $|E_{Q_p}|$  (b). (a) was obtained from the validation of the events from 24 February 2011 to 21 February 2013, and (b) using SCRRM with the initial condition calculated with in situ data (in situ), with ASCAT data (ASCAT), AMSR-E data (AMSR-E) and with ERA-Land data (ERA\_LAND). MISDc refers to the use of “Modello Idrologico Semidistribuito in continuo (Brocca et al., 2011b). \* Results for AMSR-E refer to only five events (AMSR-E sensor ceased to work on October 2011).

NS values obtained from the calibration of SCRRM using in situ SM data range from 0.12 to 0.80 with  $\overline{NS} = 0.48$ . The calibration was also carried out to take into account the soil layer depth influence on the RR transformation by using the parameter  $T$  as it was done for the two satellite products. In this case, the results only showed a little improvement and are not presented here for the sake of brevity. For ASCAT and AMSR-E,  $\overline{NS}$  is equal to 0.48 and 0.47 and median NS is equal to 0.45 and 0.41, respectively. Note that in this case, the calibrated values of  $T$  (168.3 and 120 days for ASCAT and AMSR-E, respectively, see Table 3) are much higher than the ones presented in Sect. 4.1. These results are consistent with those of (Brocca et al., 2010) which highlighted an influence of a deep soil layer in the RR transformation process. For ERA-Land SM used in SCRRM, NS ranges from 0.10 to 0.87 with a mean of 0.52. The SM in this case was calculated by considering the weighted mean of the three SM available layers (0–7, 0–28, 28–100 cm) because it yielded better results with respect to the first two layers. For MISDc the performance in calibration is similar to the ones obtained from SCRRM, with NS ranging from 0.08 to 0.79 and mean NS equal to 0.48.

Overall, in terms of mean NS, the performance scores of SCRRM are encompassed in the ones obtained by using MISD with observed initial conditions, and those considering a constant  $S$ . Moreover, they are similar to those of MISDc. More significant events (e.g. 11 December 2009, 3 February 2011, 2 January 2011), which are critical for flood forecasting, yield better results for all the selected indicators.

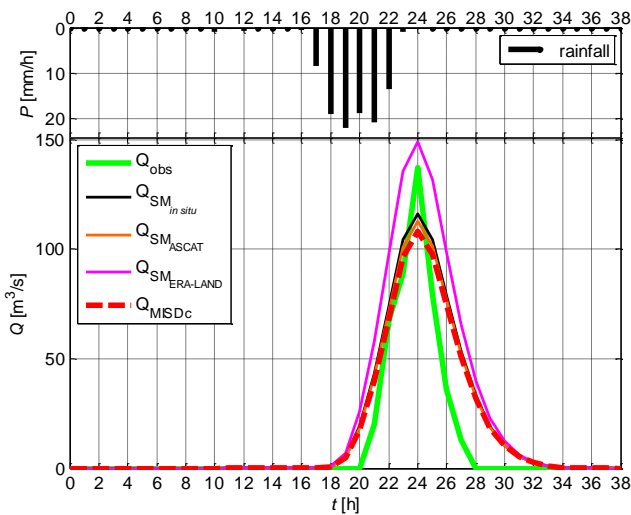
In validation, the results in terms of mean and median NS for SCRRM show a slight deterioration of the performances. The two satellite products perform in a quite similar way when choosing the same number of events, whereas the results of ERA-Land are comparable with those obtained through MISDc. Figure 5a, b plots the cumulative frequency of the absolute error in volume,  $E_V$ , and in peak discharge,

$E_{Q_p}$ , obtained in validation, respectively. For  $E_V$ , all products and MISDc show the same performances whereas, when considering  $E_{Q_p}$ , ERA-Land shows better scores with respect to ASCAT and in situ data.

To sum up, in validation SCRRM gives results similar to MISDc, with ERA-Land SM slightly overperforming the satellite products in terms of NS and  $E_{Q_p}$ . In situ data, which are based on soil moisture measurements at a depth of 25 cm, have a behaviour similar to the other indicators both in terms of mean NS and  $E_V$  (Fig. 5a), but worse results in terms of median  $E_{Q_p}$  (Fig. 5b).

#### 4.2.4 Event of 21 February 2013

The event of 21 February 2013 is particularly interesting for the area since it is one of the largest recorded at the Rafina gauged site since the installation of the gauge. As a result, it is particularly useful in assessing the robustness of the SCRRM model given that the model was calibrated for small to medium events. Unfortunately, for this event only SM from ASCAT, in situ, and ERA-Land was available (not from AMSR-E). Figure 6 shows the observed and predicted flood hydrograph for such an event obtained from SCRRM using SM from in situ, ASCAT and ERA-Land data (using the parameter calibrated on the first eight events and shown in Table 3) as well as results obtained from MISDc. As it can be seen, results are good using SCRRM with in situ and ASCAT data: NS is equal to 0.85 and 0.78,  $E_{Q_p}$  is equal to  $-14.5$  and  $-12.3$  % and  $E_V$  is equal to 52.2 and 48 %, respectively. ERA-Land slightly overestimates the peak discharge with  $E_{Q_p} = 9.5$  % but yields larger errors in NS = 0.49 and in volume  $E_V = 93.8$  % probably due to an overestimation of the initial SM conditions. Eventually, results are in line with MISDc which yields NS = 0.82 and  $E_{Q_p} = -20.5$  %.



**Fig. 6.** Validation of the event of 21 February 2013 at the Rafina gauged site. Comparison between observed  $Q_{\text{obs}}$  and predicted flood hydrographs using SCRRM with in situ data ( $Q_{\text{SM}_{\text{in situ}}}$ ), ASCAT data ( $Q_{\text{SM}_{\text{ASCAT}}}$ ), ERA-Land data ( $Q_{\text{SM}_{\text{ERA-LAND}}}$ ) and MISDc ( $Q_{\text{SM}_{\text{MISDc}}}$ ).

## 5 Conclusions

A Simplified Continuous RR Model (SCRRM) has been developed to simulate discharge hydrographs in a small Greek catchment located in the Attica region. The model uses globally available soil moisture as a proxy of the wetness state of the catchment and event rainfall data (i.e. continuous rainfall time series are not needed) to simulate hourly discharge hydrographs. In the proposed model, the soil moisture provided by an external indicator is used to infer the initial condition through a linear experimental relationship previously studied by Brocca et al. (2009a, b). This new modelling approach keeps the advantages of event-based RR models but also overcomes the issues related to the selection of a proper initial condition.

Different SM indicators were used within SCRRM and compared in terms of flood hydrographs prediction (in situ, ASCAT, AMSR-E, and ERA-Land data) and against the outcomes of a continuous model. The model robustness was successfully tested by a split sample test procedure. Despite the inherent uncertainties in the available rating curve, it was found that

1. SCRRM model satisfactorily reproduces the selected flood events with performances comparable with those obtained from a continuous simulation approach (i.e. MISDc);
2. all the selected indicators work well, with the ERA-Land product slightly overperforming (on average) in situ and satellite products;

3. satellite products perform similar to each other, yielding results comparable with those obtained from MISDc. Particularly for ASCAT, an anomalous behaviour (already observed in several arid and semi-arid regions) was observed in dry periods, which involved an anomalous increase of the soil moisture despite the absence of rainfall;
4. SCRRM is robust since it provides good performance in validation for one of the most significant events recorded at the Rafina gauged site, which occurred on February 2013 (based on calibrated parameters obtained for small to medium magnitude events).

Overall, the similarity of the performance between SCRRM and a standard continuous simulation approach recommend the use of global available products to overcome the lack of ground data for hydrological applications in many areas of Europe. Moreover, the limited number of parameters used by the model along with the ever-increasing availability of globally available soil moisture data sets necessitates a greater effort in research for the applicability of the model for operational flood forecasting purposes to poorly gauged sites of the Mediterranean areas. Future studies need to assess the performance of the model in other catchments to support these conclusions.

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