

Transpiration and CO₂ fluxes of a pine forest: modelling the undergrowth effect

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Abstract. A modelling study is performed in order to quantify the relative effect of allowing for the physiological properties of an undergrowth grass sward on total canopy water and carbon fluxes of the Le-Bray forest (Les-Landes, South-western France). The Le-Bray forest consists of maritime pine and an herbaceous undergrowth (purple moor-grass), which is characterised by a low stomatal control of transpiration, in contrast to maritime pine. A CO₂-responsive land surface model is used that includes responses of woody and herbaceous species to water stress. An attempt is made to represent the properties of the undergrowth vegetation in the land surface model Interactions between Soil, Biosphere, and Atmosphere, CO₂-responsive, ISBA-A-g_s. The new adjustment allows for a fairly different environmental response between the forest canopy and the understory in a simple manner. The model's simulations are compared with long term (1997 and 1998) micro-meteorological measurements over the Le-Bray site. The fluxes of energy, water and CO₂, are simulated with and without the improved representation of the undergrowth vegetation, and the two simulations are compared with the observations. Accounting for the undergrowth permits one to improve the model's scores. A simple sensitivity experiment shows the behaviour of the model in response to climate change conditions, and the understory effect on the water balance and carbon storage of the forest. Accounting for the distinct characteristics of the undergrowth has a substantial and positive effect on the model accuracy and leads to a different response to climate change scenarios.

Key words. Hydrology (Surface processes; Soil moisture; Evapotranspiration)

1 Introduction

In the past few years, an increasing number of long-term studies on the energy and mass exchange between forest canopies and the atmosphere have been conducted

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(Tenhunen et al., 1998; Valentini et al., 1999; Wilson and Baldocchi, 2000; Wilson et al., 2001; Anthoni et al., 2002). These data represented an opportunity to test and modify Soil-Vegetation-Atmosphere Transfer (SVAT) schemes, and, in particular, to study the vegetation response to various climatic conditions in terms of the heat, water and carbon budget (Baldocchi and Wilson, 2001). These long-term measurements also allowed for the investigation of the modelling of biological processes like photosynthesis, transpiration, plant growth, and vegetation sensitivity to climatic events (drought in particular). The SVAT scheme ISBA-A-g_s (Calvet et al., 1998) is a version of the ISBA land surface model (Noilhan and Planton, 1989; Noilhan and Mahfouf, 1996) of Météo-France, coupled with an A-g_s (Appendix A) photosynthesis model (Jacobs et al., 1996). In ISBA-A-g_s the net photosynthetic assimilation (A_n) response to environmental factors is modelled in a rather mechanistic way and the stomatal conductance (g_s) is derived from A_n (Collatz et al., 1991). ISBA-A-g_s is a research SVAT scheme used in the atmospheric model Meso-NH (Lafore et al., 1998).

In a previous study, the ISBA-A-g_s model was improved in terms of water balance and plant response to water stress: in Calvet et al. (2004) the water stress parameterisation proposed for herbaceous species by Calvet (2000) was adapted to the case of woody plants. A large number of data were used to determine the response of woody species to soil moisture stress. In particular, the leaf-scale measurements performed by Picon et al. (1996) over two tree species, presenting contrasting responses to drought (maritime pine – *Pinus pinaster* – and sessile oak – *Quercus petraea*), were employed. In response to water stress, maritime pine presented higher values of the water use efficiency (WUE), a drought-avoiding strategy, while sessile oak presented little change in WUE, a drought-tolerant strategy.

In this study, we performed, for the first time, a detailed validation of the model by using long time series of measured water vapour and CO₂ fluxes. In particular, we attempted to determine to what extent the hypothesis of homogeneous leaf properties could be used in ISBA-A-g_s. The response to soil moisture stress presented in Calvet et al. (2004) was used in ISBA-A-g_s and the simulations were compared to measured

Table 1. The ISBA scheme soil and vegetation structure parameters over the Le-Bray forest. The values are either prescribed, measured, or calibrated.

Symbol	Definition	Value	Source
Soil parameters			
d_2	Soil root-zone depth	0.8 m	Measured – Ogée et al., 2003
<i>CLAY</i>	Clay fraction	2%	Measured – Ogée et al., 2003
<i>SAND</i>	Sand fraction	98%	Measured – Ogée et al., 2003
w_{wilt}	Wilting point	0.03 m ³ m ⁻³	Measured – Loustau et al., 1990
w_{fc}	Field Capacity	0.17 m ³ m ⁻³	Measured – Loustau et al., 1990
w_{sat}	Saturation	0.43 m ³ m ⁻³	Measured – Berbigier et al., 2001
C_3	Drainage coefficient	0.012	Calibrated – Ogée et al., 2001
Vegetation parameters			
h	Vegetation height	18 m	Measured – Berbigier et al., 2001
<i>veg</i>	Vegetation coverage	100%	Prescribed
α	Albedo	0.11	Measured – Berbigier et al., 2001
ε_s	Emissivity	0.98	Prescribed
C_v	Thermal coefficient	$1.2 \times 10^{-5} \text{ K m}^{-2} \text{ J}^{-1}$	Calibrated – Sect. 2.6
z_0	Roughness length	2.0 m	Measured – Ogée et al., 2003
$\ln(z_0/z_{0h})$	kB^{-1} term	3	Measured – Lagouarde et al., 1994
d	Displacement height	16.6 m	Measured – Ogée et al., 2003

fluxes at one FLUXNET (Valentini et al., 1999) long-term site in France: the Le-Bray forest (maritime pine – *Pinus pinaster*). The Le-Bray forest data set permitted one to assess different parameterisations of the biological processes, and to simulate the canopy energy, water and carbon balance. In particular, the Le-Bray forest presents a separate grass understory, and an intercomparison of two versions of ISBA-A-g_s, with and without a representation of distinct leaf characteristics of trees and grass, was performed. Finally, a sensitivity study was carried out with the Le-Bray forest site data in order to investigate the effect of climate change on water budget, net assimilation and net ecosystem exchange of CO₂.

2 Material and methods

In this section, the data are presented, as well as the methods we used in order to:

- characterise the canopy response to water stress (either modelled or observed),
- model the specific characteristics of the undergrowth vegetation,
- model the specific soil drainage properties of Le-Bray,
- characterise the energy transfer within the vegetation canopy,
- model the ecosystem respiration,
- assess the model accuracy.

The simulations performed by ISBA-A-g_s are presented and analysed in Sect. 3.

2.1 The Le-Bray site

In this study, the maritime pine forest of Le-Bray was studied for two annual cycles (1997 and 1998). The measurements were performed by INRA-Bordeaux. Table 1 summarizes the soil and vegetation parameters of the forest. The Le-Bray site (Berbigier et al., 2001) is located in the Les-Landes forest at about 20 km from Bordeaux (44°42N, 0°46W, altitude 62 m). The Le-Bray forest consists of a maritime pine stand of about 30 years and 18–19 m tall. The forest canopy exhibits two well-separated foliage layers. The upper one, between 12 and 18 m, consists of tree crowns with a maximum leaf area index (LAI) of 2.8 m² m⁻². The undergrowth forms the lowest layer and consists of purple moor-grass (*Molinia coerulea* L. Moench). The purple moor-grass is perennial but the leaves are green only from April to September, with a maximum LAI and height of 1.45 m² m⁻² and 0.7 m, respectively (Loustau and Cochar, 1991). The resulting total green LAI varies between 1.8 m² m⁻² in winter and 4.2 m² m⁻² in summer (Porté, 1999; Ogée et al., 2003). It reaches its maximum value in August and, during this period, the purple moor-grass LAI represents one-third of the total LAI (Fig. 1). The soil is a sandy hydromorphic podzol and a compact sand-layer called Alios is found at a depth of about 0.8 m. The hardness and chemical composition of this layer makes it barely penetrable by the roots, so that the extractable water content is low (about 115 mm). The soil moisture variation (Fig. 1) was measured above the Alios layer. Because of the low hydraulic conductivity of the Alios layer, the vertical

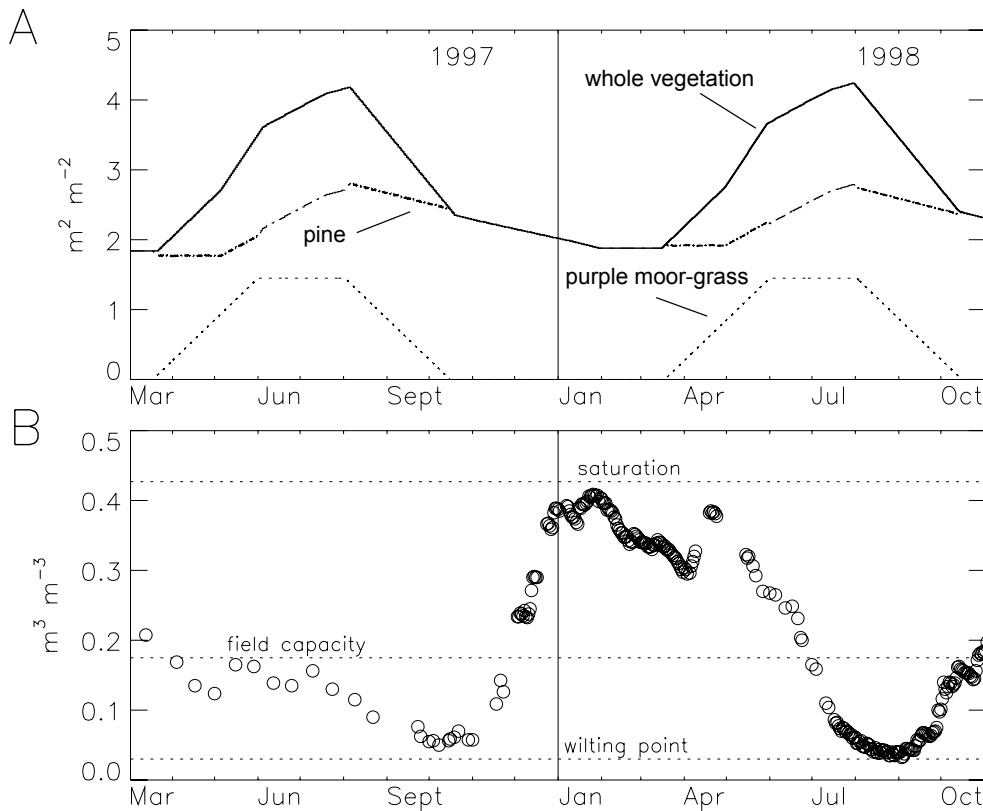


Fig. 1. Observed (A) leaf area index (LAI) and (B) root-zone soil moisture for Le-Bray. The Le-Bray LAI is the sum of the pine and purple moor-grass LAI (long- and short- dashed lines, respectively). Soil moisture measurements concern the 0.8 m soil column.

drainage is limited (Ogé and Brunet, 2002), and during winter and spring, the soil water content is generally above field capacity. During the two winters of the study, the water table level reached the soil surface. However, 1998 was drier than 1997 (the cumulative annual precipitation was 846 and 998 mm, respectively). The average annual precipitation and air temperature for this site are 930 mm and 12.5°C, respectively.

2.2 Micro-meteorological observations

Eddy fluxes of momentum, sensible heat (H), water vapour (E) and CO₂ (F_c) were measured above the tree crown at 25 and 18 m above ground (Berbigier et al., 2001), as well as net radiation (R_n). Heat storage terms (air, vegetation and soil) were estimated from temperature profiles measured by thermocouples placed at different levels in the soil and in the canopy. The meteorological variables (precipitation, solar radiation R_g , air temperature T_a , relative humidity q_a , and wind speed) were monitored on a 30-min basis. Surface infrared temperature, atmospheric pressure and atmospheric downward long-wave radiation (R_a) were measured.

2.3 Comparing leaf and canopy responses to water stress

The parameterisation of soil moisture stress (Appendix B) obtained by Calvet et al. (2004) from leaf-level gas-exchange

measurements performed by Picon et al. (1996) was employed to analyse the micrometeorological FLUXNET data,

- by implementing the improved leaf-level A- g_s parameterisation in the ISBA-A- g_s model of the canopy,
- by comparing the response of the leaf- to canopy-level stomatal conductances (g_s and g'_s , respectively), either observed or modelled, to leaf- and canopy-to-air saturation deficits (D_s and D'_s , respectively), and to the soil moisture availability.

In order to assess the effect of the purple moor-grass undergrowth on the Le-Bray forest, the leaf-level A- g_s module (producing $g_s - D_s$ relationships) was compared with the canopy-level observations of Le-Bray and with the canopy-level simulations of ISBA-A- g_s (producing $g'_s - D'_s$ relationships). The technique employed to obtain $g'_s - D'_s$ relationships at the canopy scale, for various soil moisture conditions, is described in Appendix C. All the micrometeorological measurements required to perform a $g'_s - D'_s$ analysis at the canopy scale (heat and water vapour fluxes, surface temperature, air temperature and humidity, LAI) were available for Le-Bray. The parameters of the A- g_s module are the unstressed mesophyll conductance at 25°C g_m^* , the maximum ratio between the intercellular and the atmospheric CO₂ concentration f_o^* , the maximum leaf-to-air

Table 2. Typical values of the variables of the A-g_s model for C3 plants, and temperature response. Note that ε_0 is the maximum quantum use efficiency, Γ the compensation point, $A_{m,max}$ the maximum net assimilation of the leaf, and g_m the mesophyll conductance (* denotes unstressed soil moisture conditions). The temperature dependence of Γ is described by means of a Q_{10} response function $X(T_s)=X(@25)Q_{10}^{(T_s-25)/10}$. T_1 and T_2 denote temperatures that enable a description of temperature inhibition characteristics: $X(T_s)=X(@25)Q_{10}^{(T_s-25)/10} \{[1+\exp\{0.3(T_1-T_s)\}][1+\exp\{0.3(T_s-T_2)\}]\}^{-1}$, where $X(T_s)$ and $X(@25)$ are the values of the variables corresponding to the leaf temperatures T_s and 25°C, respectively.

Parameter (X)	X_{25}	Q_{10}	T_1 (°C)	T_2 (°C)
ε_0 (mg J ⁻¹ PAR)	0.017	–	–	–
Γ (μmol mol ⁻¹)	45	1.5	–	–
$A_{m,max}$ (mg m ⁻² s ⁻¹)	2.2	2.0	8	38
g_m or g_{m^*} (mm s ⁻¹)	X_{25}	2.0	5	36

saturation deficit D_{max}^* , the cuticular (or minimum) conductance g_c and the critical extractable soil moisture θ_C . The corresponding parameter values for maritime pine are (Calvet et al., 2004) 1.58 mm s⁻¹, 0.61, 100 g kg⁻¹, 0 mm s⁻¹, and 0.10, respectively. These values are listed in Table 3.

2.4 Modelling the undergrowth effect

Previous studies have shown that the Le-Bray purple moor-grass undergrowth represents about 1/3 of the summertime total LAI and contributes 1/3 of the total stand evapotranspiration (Loustau and Cochard, 1991), whereas only 20% of the incident PAR reaches the undergrowth (Hassika et al., 1997).

In ISBA-A-g_s, the effect of PAR extinction by the vegetation canopy is accounted for by using a simple radiative transfer model and by integrating A_n and g_s over the vertical. A homogeneous leaf vertical distribution is assumed and the integration is performed by using a Gauss three-point quadrature (Calvet et al., 1998). In order to account for the specific behaviour of the undergrowth, different leaf properties were prescribed:

- at the two highest integration Gauss levels (Appendix D) of the radiative transfer module of ISBA-A-g_s, the parameters of maritime pine were prescribed in the A-g_s module, using the woody drought-avoiding strategy described in Calvet et al. (2004),
- at the lowest Gauss level both the herbaceous drought-tolerant and drought-avoiding strategies were tested, with specific parameters corresponding to purple moor-grass.

This new version of ISBA-A-g_s is described in Appendix D. This simple and pragmatic method allowed us to account

for undergrowth vegetation while using a single-layer SVAT model. These changes only affected the computation of net assimilation and stomatal conductance; other variables of the ISBA-A-g_s model, like surface temperature, CO₂ concentration or saturation deficit, remained representative for the whole canopy.

Purple moor-grass is characterised by a large value of the minimum conductance g_c . The g_c value given by Kerstiens (1996) for this species is more than 1 mm s⁻¹ (Table 3). This value is consistent with the high values of leaf conductance observed by Loustau and Cochard (1991) for the Le-Bray forest undergrowth, during dry periods characterised by high D_s values. Several simulations were made with the modified ISBA-A-g_s, in order to estimate the other A-g_s parameters of the undergrowth (g_m^* , D_{max}^* , θ_C). The effect of soil moisture on the g_m and D_{max}^* variables was described by using the parameterisation proposed by Calvet (2000) for C₃ herbaceous plants. The drought-tolerant response of g_s and A_n to soil water stress was found to provide better results than the drought-avoiding strategy. The following unstressed parameters were used (Table 3): $g_m^*=1.79$ mm s⁻¹ (at a temperature of 25°C), $D_{max}^*=19$ g kg⁻¹, with a value of θ_C of 0.3. The low value of D_{max}^* obtained for purple moor-grass (Table 3) is characteristic of a high sensitivity of stomatal aperture to leaf-to-air saturation deficit, as opposed to the low sensitivity of the water use efficiency to the soil water stress, characteristic of a drought-tolerant strategy. The other parameters of the A-g_s model listed in Table 2, were prescribed according to typical values given in Calvet et al. (1998) for C₃ plants.

2.5 Modelling the Alios effect

The Le-Bray soil moisture evolution is driven, to a large extent, by the presence of a near-surface water table in winter, due to the Alios layer (Sect. 2.1) acting as an obstacle to drainage at the bottom of the root zone. Consequently, the effective soil hydraulic conductivity for Le-Bray (Delpech, 1996) is much smaller (0.28 m d⁻¹) than for a usual sandy soil (15 m d⁻¹). In such conditions, the texture-driven value of the drainage coefficient C_3 proposed by Noilhan and Planton (1989) and used in the ISBA-A-g_s model was not applicable and had to be calibrated. Drainage was not measured at Le-Bray but it was the only unknown term of the water balance equation. The available measurements of precipitation, evapotranspiration, and soil water content, were used by Ogée et al. (2001) to estimate the C_3 parameter: $C_3=0.012$ (Table 1), a value about 215 times lower than the ISBA value for a sandy soil.

2.6 Forest heat transfer parameters

The vegetation heat transfer coefficient C_v used in the ISBA model and the logarithmic ratio between momentum and heat roughness lengths (z_0 and z_{0h} , respectively), also called kB^{-1} , had to be estimated. For open canopies, an empirical value of $kB^{-1}=2.3$ is often assumed, but significantly higher values were obtained by calibration (Calvet et al., 1998)

Table 3. Values of the parameters for the trees and undergrowth of Le-Bray.

Note that g_m is the mesophyll conductance, f_0 the maximum potential value of the ratio between internal and external leaf concentration of CO₂, D_{max} the maximum value of air saturation deficit (* denotes unstressed soil moisture conditions), g_c the leaf cuticular conductance and θ_c the critical extractable soil moisture.

Plant species	g_m^* (mm s ⁻¹)	f_0^*	D_{max}^* (g kg ⁻¹)	g_c	θ_c (mm s ⁻¹)	Response to soil moisture stress	Source
Maritime pine	1.58	0.61	100	0.0	0.10	Drought-avoiding ^a	Calvet et al., 2004
Purple moor-grass	1.79	0.95	19	1.07	0.30	Drought-tolerant ^b	Kerstiens (1996) for g_c Optimisation of the model's score on fluxes for the other parameters

^a Stress response for woody species by Calvet et al., 2004

^b Stress response for herbaceous species by Calvet, 2000

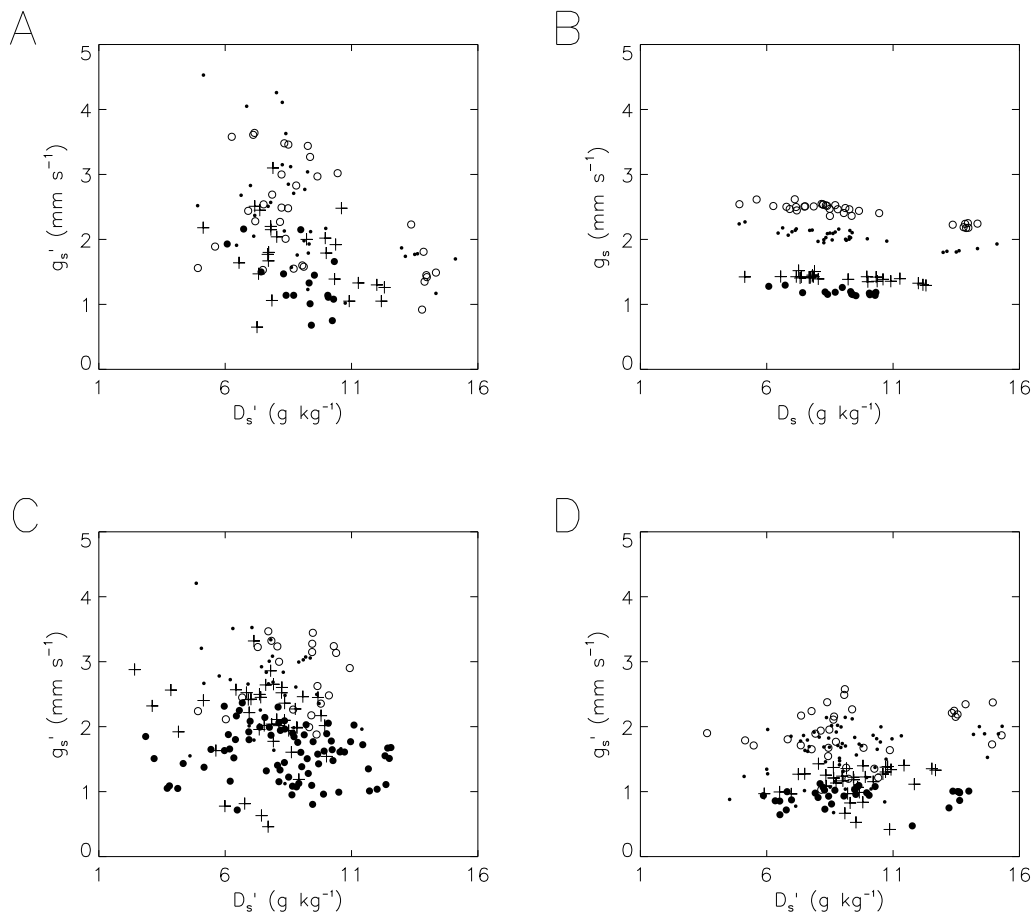


Fig. 2. Observed and simulated stomatal conductance of the Le-Bray forest versus leaf-to-air saturation deficit, for various soil moisture conditions: values of the normalised extractable soil moisture content, comprised in the intervals 0.10–0.25, 0.30–0.45, 0.70–0.80, 0.9–1.0 (solid circles, pluses, points, and open circles, respectively). **(A, C, D)** Canopy-level conductance g'_s vs. effective leaf-to-air saturation deficit D'_s , **(B)** leaf-scale g_s vs. D_s . The observed and simulated estimates were obtained from **(A)** micrometeorological measurements of Berbigier et al. (2001), **(B)** leaf-scale simulations by the A- g_s model using the Picon drought-avoiding parameterisation ($g_m^* = 1.58$ mm s⁻¹, $f_0^* = 0.61$, $D_{max} = 100$ g kg⁻¹, $\theta_c = 0.10$), and from ISBA-A- g_s simulations using the Picon drought-avoiding parameterisation, considering **(C)** a representation of the purple moor-grass undergrowth in the pine forest, or **(D)** a single layer of maritime pine. The g'_s and D'_s values were obtained by using the method described in Appendix C, for both the model (C, D) and the observations (A).

or experimental measurements (Verhoef et al., 1997). A value of 3, consistent with measurements of Lagouarde et al. (1994) over Le-Bray, was used in our simulations.

The parameter C_v was estimated by minimizing the rms error between simulated and observed half-hourly averaged total storage flux (in the ground and in different compartments of the canopy) given by Lamaud et al. (2001). The optimal value $C_v=1.2 \times 10^{-5} \text{ K m}^{-2} \text{ J}^{-1}$, is similar to the value, found by Delire et al. (1997) for an Amazonian forest site.

2.7 Estimating the ecosystem respiration

In-situ eddy correlation measurements provided half-hourly estimates of the net CO₂ flux, F_c (in this study, positive values of F_c correspond to a downward flux), representing the difference between ecosystem gross primary production A_g (i.e. the sum of net assimilation A_n and autotrophic dark respiration R_d) and the ecosystem respiration R_e :

$$F_c = A_g - R_e. \quad (1)$$

The R_e values were estimated with nighttime net CO₂ fluxes on turbulent nights (friction velocity $u^* > 0.4$) by Berbigier et al. (2001). The obtained expression is given by Eq. (2), in units of $\text{mg CO}_2 \text{ m}^{-2} \text{ s}^{-1}$:

$$R_e = 0.0422 \exp(0.105 T_{\text{Soil-LB}}), \quad (2)$$

where $T_{\text{Soil-LB}}$ represents the in-situ surface soil temperature (°C), at the soil-litter interface (Berbigier et al., 2001). In order to account for the difference between $T_{\text{Soil-LB}}$ and the soil temperature values (T_2) produced by ISBA-A-g_s, a Q_{10} temperature-response function was used to represent R_e in ISBA-A-g_s:

$$R_e = R_{e25} \cdot Q_{10}^{(T_2-25)/10}. \quad (3)$$

Best-fit R_{e25} and Q_{10} parameters of Eq. (3) were obtained by minimizing the rms error between the R_e estimates given by Eq. (2), based on the measured $T_{\text{Soil-LB}}$, and the values of R_e given by Eq. (3), based on the modelled T_2 . Equation (3) was found to be a good approximation of Eq. (2), with rms errors less than $0.03 \text{ mg CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, and a square correlation coefficient value higher than 0.85. The R_{e25} value of Eq. (3) obtained for Le-Bray was $0.39 \text{ mg CO}_2 \text{ m}^{-2} \text{ s}^{-1}$, with a Q_{10} value of 2.22.

2.8 Model statistics

The employed statistics to assess the model performance were the rms difference, the mean bias, and the skill score. The model skill score Ef (or Nash criterium) is defined as:

$$Ef = 1 - \frac{\sum_i (x_{iOBS} - x_{iMOD})^2}{\sum_i (x_{iOBS} - \bar{x}_{OBS})^2}, \quad (4)$$

where x are the values of the considered variable, either measured or simulated (*OBS* and *MOD* subscripts, respectively). A value of Ef of 1 corresponds to a perfect simulation and

values of Ef close to 0 correspond to a basic representation of x by its constant average value. Values of Ef lower than about 0.2 show that the model is not adequate.

3 Results

The ISBA-A-g_s model simulations for the Le-Bray forest were performed for the two years 1997 and 1998, by using the parameters given in Tables 1–3. The model was run by prescribing the LAI series presented in Fig. 1. An adaptation of ISBA-A-g_s to the presence of a contrasting vegetation undergrowth (Sect. 2.4) was used, and the stomatal behaviour of the pines was described by using the results of Calvet et al. (2004) at the leaf level.

3.1 Undergrowth contribution for Le-Bray

The purple moor-grass undergrowth had a pronounced effect on the canopy conductance g'_s . This conclusion was derived from the analysis of the relationship between g'_s and D'_s , either observed or modelled. For optimal PAR and temperature conditions, g'_s depends mainly on water stress, i.e. on atmospheric humidity and on soil water availability. The water stress is represented by two variables: the canopy-level saturation deficit D'_s and the soil moisture index θ (Eqs. (C2) and (B1), respectively). Analysing $g'_s - D'_s$ relationships at contrasting values of θ , is a way to characterise the canopy response to water stress.

In order to assess the stomatal sensitivity to air humidity of the Le-Bray forest for various soil moisture conditions, and the impact of the undergrowth vegetation, the canopy-level g'_s and D'_s derived from the measurements (and obtained by the method described in Sect. 2.3 and Appendix C) were compared with the g'_s and D'_s values produced by ISBA-A-g_s, with and without the improved representation of the undergrowth vegetation. In Fig. 2a, the observed Le-Bray $g'_s - D'_s$ relationship, based on the micrometeorological measurements, is shown for four classes of soil moisture. The different levels of soil water stress had an effect on g'_s and D'_s : while the ordinary decrease of g'_s , in response to increasing D'_s , was observed, a slight but noticeable difference appeared in the relative position of stressed and unstressed conditions. However, the obtained g'_s and D'_s values disagreed with the behaviour modelled in Calvet et al. (2004), using the leaf-level measurements of Picon et al. (1996) over maritime pine (Fig. 2b). The leaf-level A-g_s simulations obtained by using the Picon leaf-level parameters given in Table 3 predicted a less marked response of g_s to saturation deficit, and a clear, independent soil moisture effect. Figure 2c and d show that the difference between the model and the observations could be explained by including the undergrowth contribution in ISBA-A-g_s: while the model response at the canopy level, considering a single layer of pine (Fig. 2d), was very similar to leaf-level simulations, accounting for the purple moor-grass layer in the model as described in Sect. 2.4 (Fig. 2c) permitted one to simulate the

Table 4. Errors affecting the simulated Le-Bray site half-hourly outputs in 1997 and 1998 for different ISBA parameterisations in terms of rms error, mean bias (simulated minus observed), and skill score on evapotranspiration (E), heat flux (H), net radiation (R_n), heat storage flux (G), surface temperature (T_s), soil water storage ($w_2 \times d_2$) and net ecosystem CO₂ flux (F_c).

ISBA Version	Statistics 1997/1998	E	H	R_n	G	T_s	$w_2 \times d_2$	F_c
Standard ^a	rms error	38/38 W m ⁻²	43/50 W m ⁻²	13/15 W m ⁻²	35/23 W m ⁻²	1.8/1.0°C	49/31 mm	–
	mean bias	3/2 W m ⁻²	22/23 W m ⁻²	-8/-11 W m ⁻²	-17/1 W m ⁻²	1.2/0.4°C	39/3 mm	–
	skill score	0.81/0.77	0.75/0.66	1/0.99	0.33/0.49	0.93/0.98	0.67/0.92	–
A-g _s ^b	rms error	43/48 W m ⁻²	53/57 W m ⁻²	12/14 W m ⁻²	34/24 W m ⁻²	1.8/1.2°C	57/27 mm	0.215/0.212 mg m ⁻² s ⁻¹
	mean bias	-2/-2 W m ⁻²	25/26 W m ⁻²	-8/-10 W m ⁻²	-15/2 W m ⁻²	1.2/0.6°C	50/9 mm	0.010/0.029 mg m ⁻² s ⁻¹
	skill score	0.76/0.63	0.62/0.55	1/0.99	0.36/0.44	0.93/0.97	0.54/0.94	0.46/0.47
A-g _s ^c	rms error	40/40 W m ⁻²	43/46 W m ⁻²	12/14 W m ⁻²	34/22 W m ⁻²	1.7/1.0°C	30/25 mm	0.204/0.198 mg m ⁻² s ⁻¹
	mean bias	8/8 W m ⁻²	16/17 W m ⁻²	-7/-9 W m ⁻²	-16/2 W m ⁻²	1.1/0.4°C	10/-18 mm	-0.015/0.003 mg m ⁻² s ⁻¹
	skill score	0.79/0.75	0.75/0.71	1/1	0.38/0.52	0.94/0.98	0.87/0.95	0.51/0.54

^a ISBA (Noilhan and Mahfouf, 1996), no representation of photosynthesis, Minimum stomatal resistance $R_{s,min}=200 \text{ s m}^{-1}$

^b ISBA-A-g_s, one layer of maritime pine, drought-avoiding parameterisation

($g_m^*=1.58 \text{ mm s}^{-1}$, $f_0^*=0.6$, $D_{max}=100 \text{ g kg}^{-1}$, $\theta_c=0.1$)

^c ISBA-A-g_s, including the characteristics of the purple moor-grass undergrowth

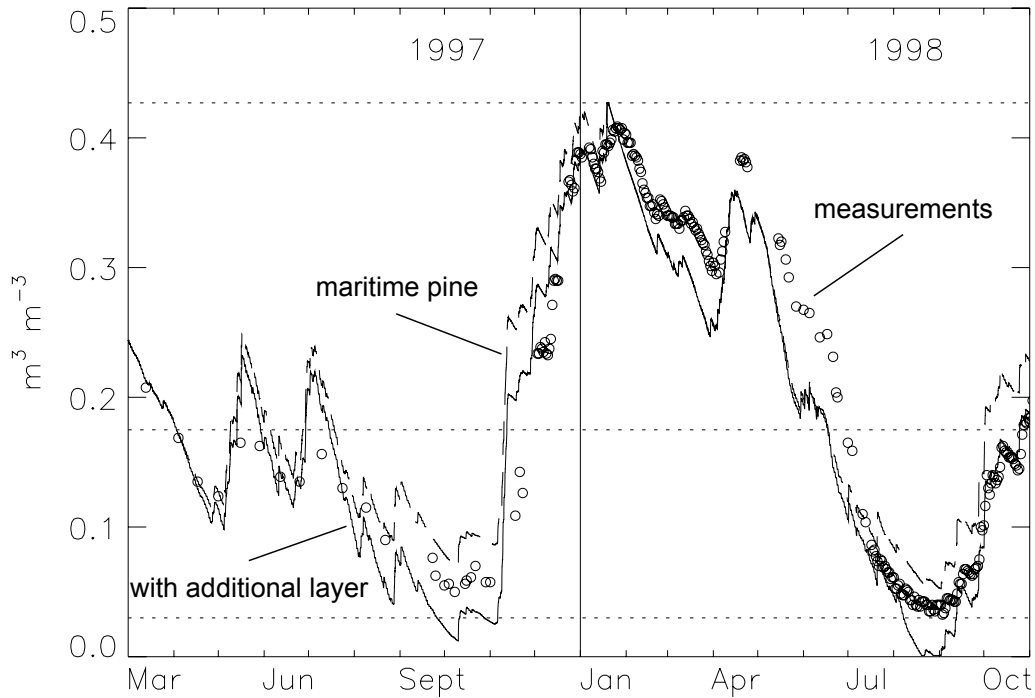


Fig. 3. Comparison of the ISBA-A-g_s simulations of soil moisture with the measurements (lines and circles, respectively), in 1997 and 1998, for Le-Bray. Two simulations are compared: considering a single layer of maritime pine, or a layer of maritime pine and an additional layer representing the purple moor-grass undergrowth (long-dashed and solid lines, respectively). The dashed lines indicate (from top to bottom) different soil water contents: saturation, field capacity, wilting point. Soil moisture measurements concern the 0.8 m soil column.

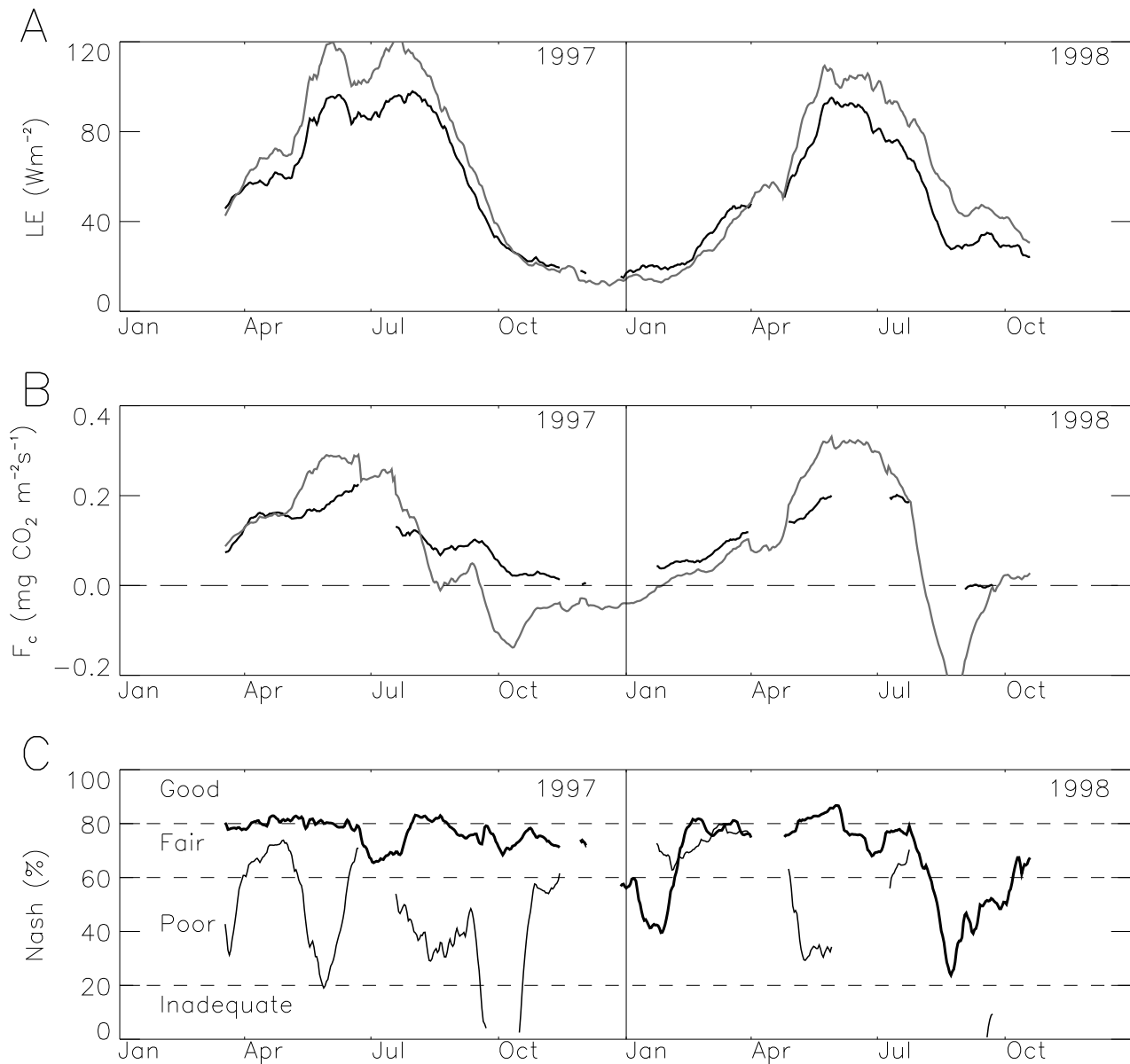


Fig. 4. Seasonal performance of the ISBA-A-g_s model over the Le-Bray forest, in terms of (A, C) evapotranspiration and (B, C) net CO₂ flux: (A, B) comparison of simulations with eddy covariance measurements (grey and dark lines, respectively), (C) skill score (Nash criterium) for evapotranspiration and net CO₂ flux (thick and fine lines, respectively). The simulated CO₂ flux was obtained by using the ecosystem respiration of Berbigier et al., 2001. The flux and skill score values are smoothed using a 30-day moving average, in order to show the general trend.

increased sensitivity to saturation deficit and the lower soil moisture control which was observed over Le-Bray (Fig. 2a). Therefore, the low apparent sensitivity of the Le-Bray forest to soil moisture stress, and its apparent high sensitivity to saturation deficit, could be explained by the presence of undergrowth species presenting a large stomatal sensitivity to saturation deficit and a high minimum conductance (Table 3). Note also that the parameters of Table 3 describing the purple moor-grass denoted a high photosynthesis capacity of the undergrowth: the $g_m * f_o *$ variable, which controls

the gross photosynthetic rate in the absence of stress (Calvet et al., 2004) reached 1.37 mm s^{-1} for the purple moor-grass, against 0.96 mm s^{-1} for maritime pine. Hence, accounting for the specific characteristics of the undergrowth is likely to have an impact on the simulation of the CO₂ flux (see below).

3.2 Surface water and energy budget

In order to test the suitability of the two versions of ISBA-A-g_s (with and without the properties of undergrowth vegetation), the model outputs (water vapour and heat fluxes, soil

Table 5. Observed and simulated terms of the carbon and water budgets of the Le-Bray forest: Gross Primary Productivity (GPP), ecosystem respiration (R_e), Net Ecosystem Exchange (NEE), evapotranspiration (E), and occurrence of normalized extractable soil moisture (θ) values lower than 0.3 and 0.6. Values are presented for 1997 (from March to December) and 1998 (from January to November).

Source	Climate ^a	GPP (g C m ⁻² yr ⁻¹)	R_e (g C m ⁻² yr ⁻¹)	NEE (g C m ⁻² yr ⁻¹)	E (mm yr ⁻¹)	Occurrence of $\theta < 0.3$ (%)	Occurrence of $\theta < 0.6$ (%)
Observations (Berbigier et al., 2001)	350	2300/-	1680/-	623/536	546/469	14/22	34/27
ISBA-A-gs ^b	350	2079/1718	1390/1017	689/701	472/449	0/10	18/20
	700	3551/2922	1401/1023	2150/1899	467/354	0/1	2/12
	700+	3503/2933	1691/1225	1812/1708	544/413	3/13	24/25
	700+pp	3444/2858	1696/1227	1748/1630	429/406	15/17	31/26
ISBA-A-gs ^c	350	1900/1558	1372/1002	528/556	573/435	26/24	38/31
	700	3455/2663	1379/1010	2076/1654	625/495	13/17	25/25
	700+	3025/2661	1653/1194	1371/1467	627/561	32/26	45/33
	700+pp	2855/2545	1662/1196	1193/1349	606/549	34/27	48/41

^a Climate scenario: 350 stands for the observed atmospheric forcing

700 stands for a CO₂-doubling scenario

700+ stands for a CO₂-doubling scenario combined with an air temperature increase of 3°C

700+pp stands for a CO₂-doubling scenario combined with an air temperature increase of 3°C, a precipitation wintertime (DJF) increase of 40% and summertime (JJA) and autumn (SON) decrease of 10 and 20%, respectively, adapted from the Southwestern Europe scenario of Déqué et al. (1998).

^b ISBA-A-gs, one layer of maritime pine, drought-avoiding parameterisation ($g_m^* = 1.58 \text{ mm s}^{-1}$, $f_0^* = 0.6$, $D_{max} = 100 \text{ g kg}^{-1}$, $\theta_c = 0.1$)

^c ISBA-A-gs, including the characteristics of the purple moor-grass undergrowth

moisture, and CO₂ fluxes) were compared with the 2-year field measurements (Table 4). The results obtained by using the standard version of ISBA (without a description of photosynthesis nor the ability to simulate the CO₂ fluxes) are presented, as well.

Accounting for the undergrowth permitted one to improve all the scores of ISBA-A-gs slightly. In particular, the best simulation of soil moisture was obtained (rms error of 30 mm in 1997, against 49 mm for the standard version of ISBA), while also performing well in terms of fluxes and surface temperature. However, the model predicted soil moisture values lower than the wilting-point at the end of the summer (Fig. 3), in contradiction with the measurements performed above the compact Alios layer. This behaviour was probably caused by the simplified modelling of the undergrowth contribution which did not permit one to simulate precisely the senescence of the purple moor-grass in September (i.e. the fraction of LAI attributed to the undergrowth was overestimated by the model). The non-stomatal evaporation of the undergrowth (due to the high value of g_c) then triggered below wilting-point evaporation. This problem was particularly acute at the end of the summer of 1998, as shown by the fall in seasonal skill score of the model on E and F_c (Fig. 4c). Note that this poorly simulated period was part of the general trend of the model to overestimate E at summertime (Fig. 4a).

3.3 Net ecosystem CO₂ exchange

Figure 4 shows that the model skill score on F_c was generally lower than the model skill score on E . In particular, F_c was poorly simulated in some situations because the employed respiration model (Eq. 3) was probably too simple to account for the complex interaction between climate and the biological phenomena involved in the respiration. Also, owing to the nonlinear nature of Eq. (3), a small deviation of R_{e25} or Q_{10} may have caused a significant bias in the estimation of the ecosystem respiration.

Accounting for the undergrowth for Le-Bray permitted one to improve the model skill score on F_c : 0.51 for 1997, against 0.46 without undergrowth description (Table 5). Also, the improvement was noticeable for cumulative values. Figure 5 compares the observed and the simulated cumulative carbon storage or net ecosystem exchange (NEE). The simulated NEE at Le-Bray corresponds to a large net sink of carbon (Fig. 5 and Table 5), consistent with the observations. However, the cumulative NEE, at the end of the two-year period, was overestimated by the single layer model. In spite of the high photosynthesis capacity of the undergrowth in unstressed conditions, accounting for the distinct characteristics of the undergrowth in the model simulations, triggered a decrease of about 20% in the modelled NEE (Fig. 5 and Table 5). This was caused by the higher sensitivity of the purple

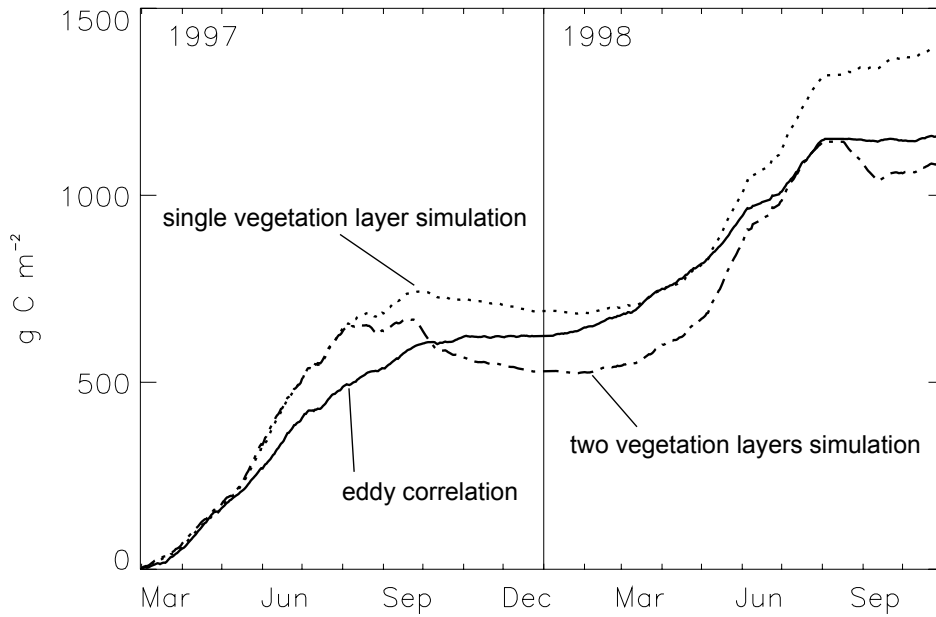


Fig. 5. The 2-year evolution of the cumulative daily net carbon storage (NEE) for the Le-Bray forest, as estimated from eddy correlation measurements and ISBA-A-g_s simulations (solid and dotted or dashed lines, respectively). The missing half-hourly data of the observed time series (9.4%) and the corresponding simulated values were replaced by zero values. Two simulations are displayed: assuming that the canopy consists of either a single layer of maritime pine or accounting for the characteristics of the purple moor-grass undergrowth (dotted and dash-dotted lines, respectively).

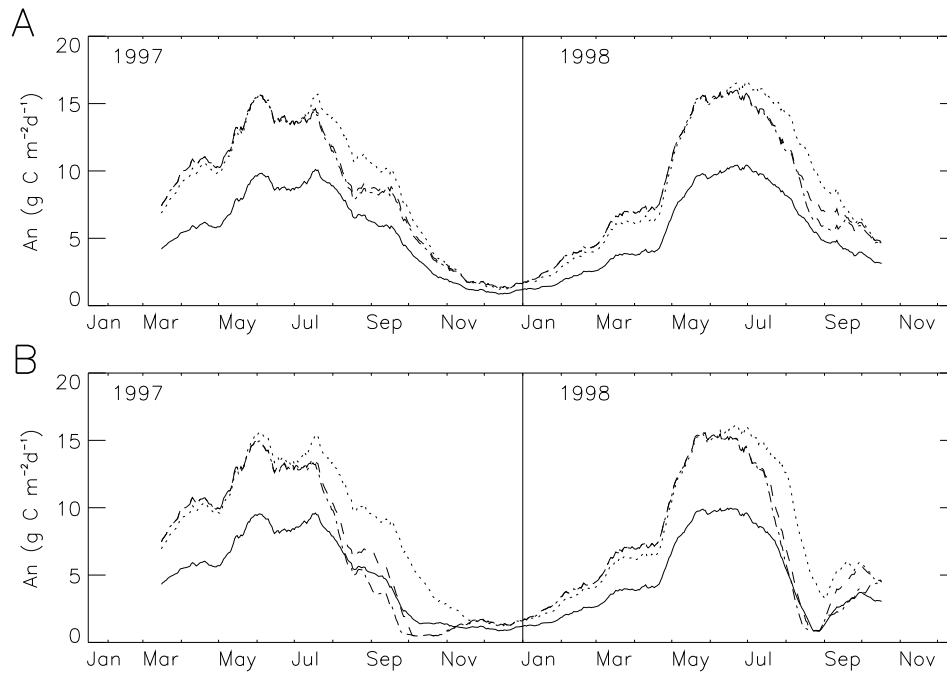


Fig. 6. Modelled leaf-net assimilation (A_n) of Le-Bray, for different climate scenarios, 350, 700, 700+ and 700+pp (solid, dotted, dashed, and dash-dotted lines, respectively). The ISBA-A-g_s model was run using the Picon drought-avoiding parameterisation ($g_m^*=1.58 \text{ mm s}^{-1}$, $f_0^*=0.6$, $D_{max}=100 \text{ g kg}^{-1}$, $\theta_c=0.01$), considering (A) a single layer of maritime pine, or (B) a layer of maritime pine and an additional layer representing the purple moor-grass undergrowth. The A_n values are smoothed using a 30-day moving average, in order to show the general trend.

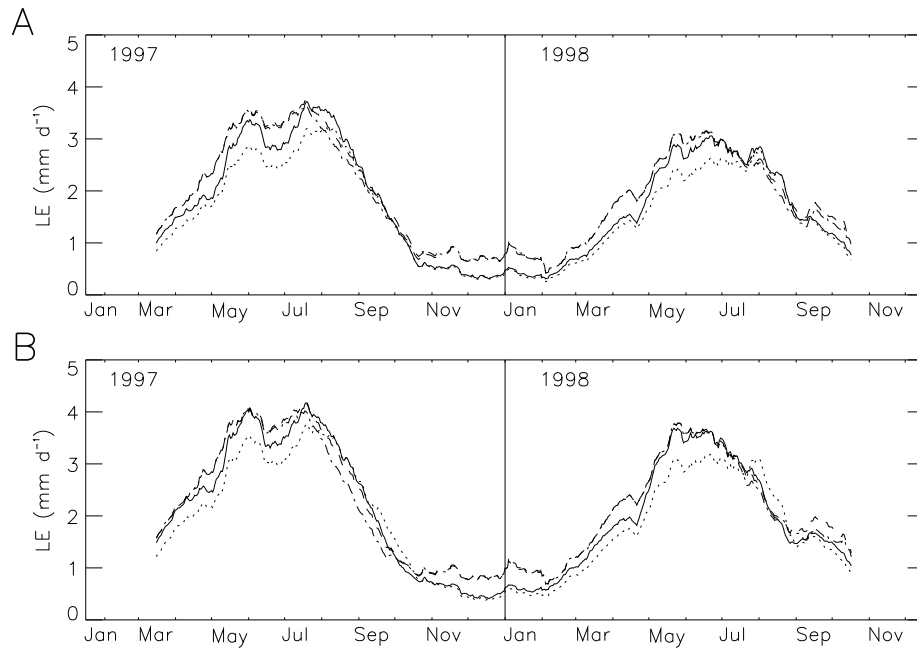


Fig. 7. Modelled ecosystem evapotranspiration (E) of Le-Bray, for different climate scenarios, 350, 700, 700+ and 700+pp (solid, dotted, dashed, and dash-dotted lines, respectively). Model parameterization used for these runs are identical to those of Fig. 6, considering: (A) a single layer of maritime pine, or (B) a layer of maritime pine and an additional layer representing the purple moor-grass undergrowth. The E values are smoothed using a 30-day moving average, in order to show the general trend.

moor-grass to the air humidity deficit and by the acceleration of the soil water depletion, related to the non-stomatal evaporation of the grass. The daytime F_c flux was overestimated by the model during the growing period and underestimated during autumn (Fig. 4). Attempts were made (not shown) to reduce the summertime overestimation of daytime F_c by prescribing lower values of g_m^* for pines: net assimilation diminished and the skill score (Eq. (1)) on F_c was improved. On the other hand, this correction tended to degrade the model performance in terms of mean bias and of cumulative values. In particular, the pre-existing underestimation of the annual gross primary productivity GPP (Table 5) was increased, causing a significant underestimation of NEE.

In spite of the improved description of the undergrowth, water vapour fluxes are better simulated than CO₂ ones. The simplicity of the respiration model is at least partly to blame.

4 Forest sensitivity to climate change

A sensitivity study was performed (Table 5 and Fig. 6), in order to highlight first-order responses to climate change of the Le-Bray forest, on both water and CO₂ fluxes, and the impact of the undergrowth vegetation. Three climate change scenarios (700, 700+, and 700+pp) were considered:

- CO₂-doubling;
- CO₂-doubling combined with an air temperature increase of 3°C;

- CO₂-doubling scenario combined with an air temperature increase of 3°C, and a seasonal redistribution of the precipitation regime (wintertime (DJF) increase of 40%, summertime (JJA) decrease of 10% and autumn (SON) decrease of 20%), respectively.

The latter scenario (700+pp) was adapted from the Southwestern Europe scenario of Déqué et al. (1998).

Irrespective of contrasting undergrowth characteristics, the simulations predicted a large effect of CO₂-doubling on GPP, which increased by 50 to 70%. The effect of CO₂-doubling on NEE was even more dramatic (700 scenario in Table 5) but was limited by the 3°C warming and by the redistribution of precipitation (700+ and 700+pp scenarios). Figure 6 shows that the first-order response of leaf-net assimilation A_n was governed by CO₂-doubling (the peak value of A_n was increased by 60 to 70%). The 3°C warming tended to accelerate the growing phase of A_n during spring.

On the other hand, the advantage of a reduction in evapotranspiration caused by CO₂-doubling (Fig. 7) was lost in the 700+ and 700+pp scenarios, and the occurrence of dry periods was increased (Table 5). This phenomenon triggered the acceleration of the decreasing phase of A_n during autumn (Fig. 6). The comparison between Figs. 6a and b shows that the purple moor-grass undergrowth was largely responsible.

Finally, prescribing in ISBA-A-g_s a climate change scenario, including CO₂-doubling, temperature increase, and redistribution of the precipitation regime, led to the following conclusions for Le-Bray:

- An increase of GPP of about 50 to 70%,

- A marked increase of NEE of about 130%,
- An increase (about 20%) of ecosystem respiration,
- A significant increase of the annual evapotranspiration and of the occurrence of drought periods.

5 Conclusion

The ISBA-A-g_s SVAT model was parameterised for the Le-Bray forest, consisting of a coniferous species – maritime pine – and a herbaceous undergrowth. Most model parameters could be estimated from in-situ measurements rather than derived from model calibration. In particular, the physiological A-g_s module of the model, assessed using independent leaf-level measurements and a functional parameterisation of the woody plant responses to water stress (Calvet et al., 2004), was used. This stress model was successfully applied to Le-Bray. This application study demonstrated the usefulness of introducing more complexity in the models of plant stomata response to atmospheric humidity and soil water availability: accounting for the undergrowth permitted one to improve the simulation of the surface fluxes, and root-zone soil moisture.

In a multi-layer stand like the Le-Bray forest, comprising a dense herbaceous undergrowth, the response to soil and atmospheric water stress of each canopy layer was different and had to be characterized. Note, however, that the purple moor-grass undergrowth of Le-Bray is a particularly extreme case since this species presents unusually high values of the leaf-minimum conductance.

The cumulated NEE was correctly described, although the simulated half-hourly net CO₂ flux was not as good as the simulated evapotranspiration. Finally, the sensitivity of the model simulations to climate change was assessed and it was shown that accounting (or not) for the undergrowth may affect the conclusions of a climate change impact study.

Appendix A

The A-g_s model

The A-g_s approach employed to describe the leaf-scale physiological processes in ISBA-A-g_s (Calvet et al., 1998) was the model proposed by Jacobs et al. (1996).

The photosynthesis rate in light-saturating conditions is expressed as

$$A_m = A_{m,\max} \left[1 - \exp \left\{ \frac{-g_m^* (C_i - \Gamma)}{A_{m,\max}} \right\} \right]. \quad (\text{A1})$$

The g_m^* variable (the unstressed mesophyll conductance) is corrected for leaf temperature using a Q₁₀-type function, together with the maximum photosynthesis $A_{m,\max}$ and the compensation point Γ . Typical values of $A_{m,\max}$ and Γ at a temperature of 25°C, for C₃ plants, are given in Table 2. The internal CO₂ concentration C_i is obtained by combining

the CO₂ concentration at the surface of the canopy C_s and Γ through the following closure equation

$$C_i = f C_s + (1 - f) \Gamma, \quad (\text{A2})$$

where the coupling factor f is sensitive to air humidity and depends on the cuticular conductance g_c and on g_m^* , f_o^* , and D_{\max}^* by

$$f = f_o^* \left(1 - D_s / D_{\max}^* \right) + \left(g_c / [g_c + g_m^*] \right) \left(D_s / D_{\max}^* \right), \quad (\text{A3})$$

where f_o^* is the value of f for $D_s=0$ g kg⁻¹ in unstressed conditions. The net assimilation is limited by a light deficit according to a saturation equation applied to the photosynthetically active radiation I_a :

$$A_n = (A_m + R_d) \left[1 - \exp \left\{ -\varepsilon I_a / (A_m + R_d) \right\} \right] - R_d, \quad (\text{A4})$$

where leaf respiration is given by $R_d = A_m / 9$, and the light conversion efficiency by $\varepsilon = \varepsilon_0 (C_i - \Gamma) / (C_i + 2\Gamma)$, where ε_0 is the maximum quantum use efficiency (Table 2). Finally,

$$g_s = g_c + 1.6 \left(A_n - A_{\min} \frac{D_s}{D_{\max}^*} \left(\frac{A_n + R_d}{A_m + R_d} \right) + R_d \left(1 - \frac{A_n + R_d}{A_m + R_d} \right) \right) / (C_s - C_i), \quad (\text{A5})$$

where A_{\min} represents the residual photosynthesis rate (at full light intensity) associated with cuticular transfers when the stomata are closed because of a high saturation deficit:

$$A_{\min} = g_m^* g_c (C_s - \Gamma) / (g_m^* + g_c). \quad (\text{A6})$$

Appendix B

Forest response to soil moisture stress

The results presented in Calvet et al. (2004) for maritime pine and sessile oak were used to build a general schematic parameterisation of the effects of the soil moisture stress in the A-g_s model, for trees. The stress function is expressed as a relative extractable soil water content, defined as

$$\theta = \frac{w - w_{wilt}}{w_{fc} - w_{wilt}}, \quad (\text{B1})$$

where w is the soil volumetric moisture in the root-zone, and w_{fc} and w_{wilt} (Table 1) are the root-zone moisture content at field capacity and wilting point, respectively.

The parameterisation was the following:

- In the drought-avoiding strategy, for values of θ higher than the critical extractable soil moisture, θ_C (that is for moderate soil moisture stress), f_0 (Appendix A) deviated from its unstressed value f_0^* through

$$f_0 = f_0^N + \left(f_0^* - f_0^N \right) (\theta - \theta_C) / (1 - \theta_C), \quad (\text{B2})$$

where f_0^N is the stressed value of f_0 corresponding to the constant unstressed mesophyll conductance, g_m^* .

- In the drought-tolerant strategy, g_m deviated from its unstressed value g_m^* through

$$g_m = g_m^N + (g_m^* - g_m^N) (\theta - \theta_C) / (1 - \theta_C), \quad (\text{B3})$$

where g_m^N is the stressed value of g_m corresponding to the constant unstressed coupling factor f_0^* .

- For values of θ below θ_C , (i.e. for more pronounced water stress)

$$g_m = g_m^N \theta / \theta_C \quad (\text{B4})$$

in the drought-tolerant strategy, while

$$g_m = g_m^* \theta / \theta_C \quad (\text{B5})$$

in the drought-avoiding strategy, and f_0 was derived from g_m by inverting the stressed regression line relating the values of g_m to f_0 in stressed conditions:

$$\ln(g_m) = 2.8 - 7.0 f_0 \quad (n = 29, R^2 = 76\%). \quad (\text{B6})$$

This regression equation was obtained by Calvet et al. (2004) by pooling the g_m and f_0 estimates of both maritime pine and sessile oak for low values of extractable soil moisture ($\theta < 0.13$).

- For the lowest values of θ the value of f_0 was forced to be lower than 0.99 in order to avoid non-physical values of g_s (in Eq. (A5) C_s must be higher than C_i). In stressed conditions, the values of g_m^* and f_0^* in Eqs. (A1), (A3), and (A6) were replaced by g_m and f_0 . The value of D_{max}^* remained unchanged.

Appendix C

Estimating canopy conductance from flux measurements

All the micrometeorological measurements required to perform a $g_s - D_s$ analysis at the canopy scale (heat and water vapour fluxes, surface temperature, air temperature and humidity, LAI) were available for the Le-Bray experimental stand. Following Calvet (2000), it was possible to estimate the equivalent stomatal conductance at the canopy scale (g_s') using estimates of the effective saturation deficit of the canopy (D_s'), and measurements of LAI and evaporative fluxes:

$$g_s' = \frac{E}{\rho_a L D_s'}, \quad (\text{C1})$$

where E , L and ρ_a represent transpiration, LAI, and air density, respectively. The value of the canopy transpiration E was not measured directly but rainless days were considered and it was assumed that the evapotranspiration measurements could be used to estimate the forest transpiration. Indeed, for rainless days, there is no evaporation from the water intercepted by the leaves, and direct soil evaporation is generally

low for forest canopies because the litter layer reduces the water vapour exchanges between the soil and the atmosphere.

The effective saturation deficit of the canopy was obtained from meteorological variables (air temperature T_a , and specific humidity q_a) measured on the flux tower, together with the infrared temperature T_s , evaporation E and heat fluxes H :

$$D_s' = \{q_{sat}(T_s) - q_a\} - c_p \left\{ \frac{E}{H} \right\} \{T_s - T_a\}, \quad (\text{C2})$$

where c_p , and q_{sat} are the specific heat of air, and the specific air humidity at saturation, respectively. In this equation, vertical temperature gradients in the canopy are assumed to be small.

The values of D_s' and g_s' were calculated for all soil moisture situations, in conditions which were favourable to the expression of a significant physiological response of the canopy to soil and atmospheric water stress: no rain, values of incoming solar radiation (R_g), latent flux (E) and sensible flux (H) higher than 400, 40 and 10 W m⁻², respectively.

Appendix D

Accounting for undergrowth vegetation in ISBA-A-g_s

In ISBA-A-g_s (Calvet et al., 1998), photosynthesis and leaf conductance are calculated at several levels in the canopy, and the obtained values are combined to obtain the canopy integrated net assimilation and conductance (A_{nI} and g_{sI} , respectively). The integration is performed through a 3-point Gauss quadrature method, following Jacobs (1995):

$$A_{nI} = L \times \sum_{i=1}^3 W_i A_n(z_i) \quad (\text{D1})$$

$$g_{sI} = L \times \sum_{i=1}^3 W_i g_s(z_i), \quad (\text{D2})$$

where z_i and W_i are the Gauss levels and weights, respectively, and L represents total LAI.

In the case of Le-Bray, a simple way of including the distinct behaviour of the purple moor-grass undergrowth is to run the A-g_s model with the purple moor-grass parameters for the near-floor Gauss level, corresponding to the lowest incident radiation. The proportion of LAI corresponding to the purple moor-grass undergrowth is 1/3, and measurements indicate that about 20% of the incident solar radiation reached the undergrowth. Since the lowest-level Gauss weight is about 1/3 ($W_3=0.28$), and since the employed radiative transfer equations predict a residual solar radiation at the lowest Gauss level of about 20%, Eqs. (D1–2) should be sufficient to account for the effect of the undergrowth vegetation on the fluxes. Note that the total LAI value is the same whether the properties of the undergrowth are accounted for or not.

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