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Multistation digisonde observations of equatorial spread F in South America

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Abstract. Directional ionogram and F-region drift observations were conducted at seven digisonde stations in South America during the COPEX campaign from October to December 2002. Five stations in Brazil, one in Argentina, and one in Peru, monitored the ionosphere across the continent to study the onset and development of F-region density depletions that cause equatorial spread F (ESF). New ionosonde techniques quantitatively describe the prereversal uplifting of the F layer at the magnetic equator and the eastward motion of the depletions over the stations. Three of the Brazilian stations were located along a field line with a 350-km apex over the equator to investigate the relation of the occurrence of ESF and the presence of sporadic E-layers at the two E-region intersections of the field line. No simple correlation was found.

Key words. Ionosphere (equatorial ionosphere; ionospheric irregularities; instruments and techniques)

1 Introduction

To fully understand the processes that control the onset of equatorial spread F it is helpful to investigate time development and spatial distribution of this F-region phenomenon. In support of the Brazilian COPEX campaign (Abdu et al., 2003), digisondes collecting ionogram and F-region drift data at seven sites in South America (Fig. 1). The sites are located within about $\pm15°$ magnetic latitude and cover about 30° in longitude. The ionosondes typically recorded one multi-directional ionograms and a number of drift measurements every 5 min during the COPEX campaign period from October to December 2002. Three ionosondes were

installed along a magnetic meridian with one sounder close to the magnetic equator (Cachimbo, 0.4° N magnetic), and the two other sounders (Boa Vista, 12.9° N; Campo Grande, −10.6° N magnetic) close to the intersections with the E-region of the 350-km field line over Cachimbo (Fig. 2). The objective was to study the influence that sporadic E-layers at the field line footprints may have on the formation of ESF at Cachimbo. Major characteristics of ESF in the South American longitude sector, and especially over Brazil, have been studied during the last two decades (see, for example, Abdu et al., 1998; Batista et al., 1990).

2 Discussion

2.1 Signatures of low-latitude spread F in ionograms

The term "spread F" was coined to give a visual description of the broadened echo traces seen on ionograms at certain times (Fig. 3). The equatorial spread F (ESF) is usually associated with depletions in F-region densities, as shown by incoherent scatter radar (ISR) measurements at Jicamarca, Peru (e.g. Woodman and La Hoz, 1976; Anderson and Mendillo, 1983 and references therein). Sales et al. (1996) have used retracing to establish that the spread traces on equatorial region ionograms are the result of off-vertical echoes that return from field-aligned irregularities in the depleted region, as identified by 630 nm all-sky images. Modern ionosondes operate essentially as radar systems, i.e. they measure radar distances and angles of arrival of the received echoes (Reinisch, 1996). In contrast to a conventional radar system, the ionosonde transmits a wide radio beam that illuminates a large area of several hundred km in diameter in the F-region, resulting in echoes returning from many directions in the presence of irregularities. With the introduction of Doppler interferometry to ionospheric sounding (Bibl and

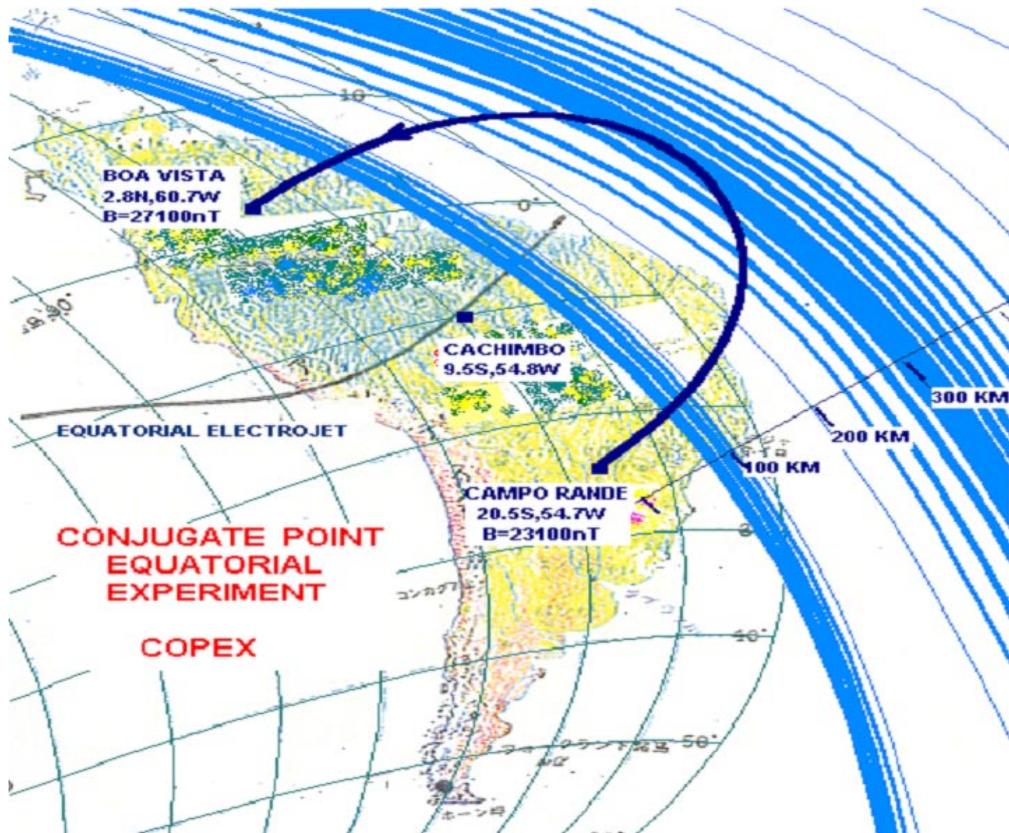**Fig. 1.** Location of the seven digisondes in South America.

Fig. 2. The Brazilian COPEX stations at the equator (Cachimbo) and the conjugates stations at Boa Vista and Campo Grande.
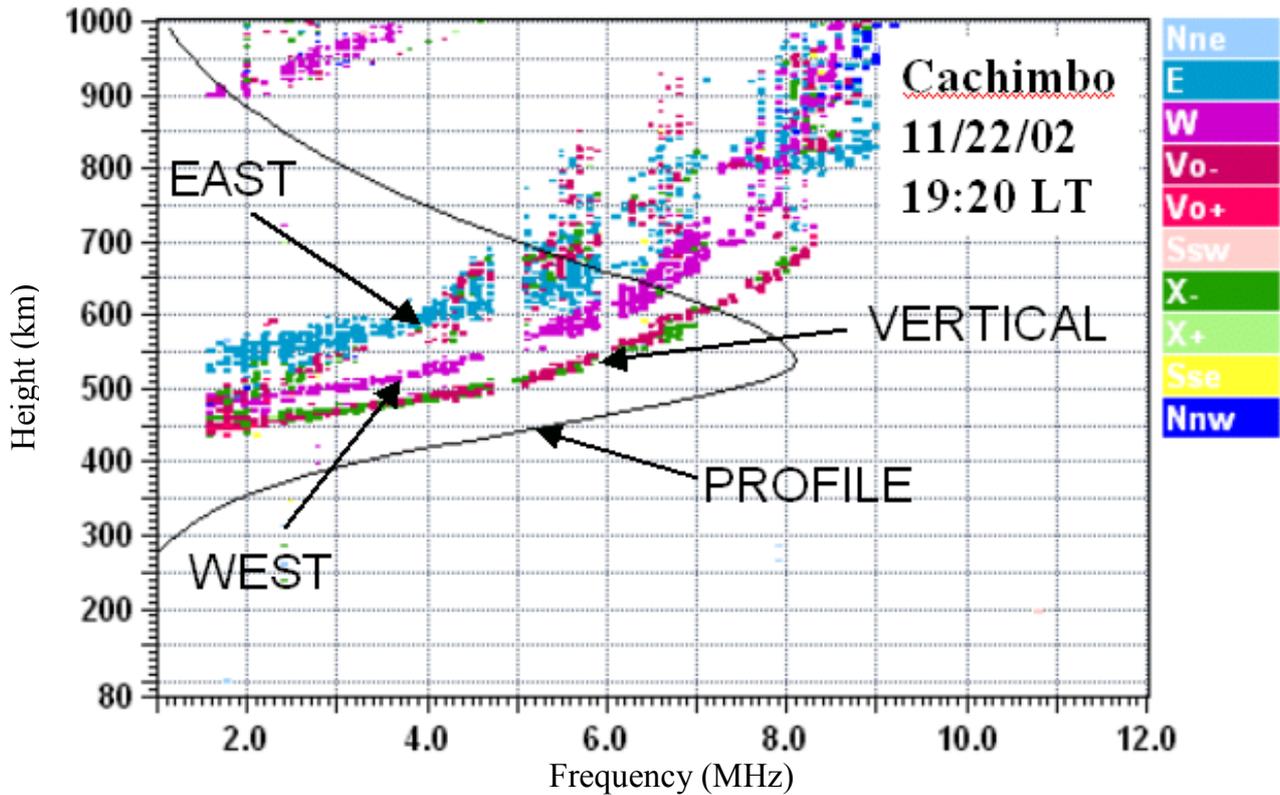
Fig. 3. Multi-directional ionogram at Cachimbo for 22 November at 19:20 LT showing echo traces from the west, east, and overhead, with an overlay of the vertical electron density (plasma frequency) profile.

Reinisch, 1978), it has become possible to identify the source regions of the spread echoes. The DPS digisondes use four small spaced antennas for signal reception arranged in a triangle with one antenna at the center. The signals from each antenna are Fourier analyzed to identify echoes with different Doppler frequencies.

Interferometry for each Doppler component then determines the source locations of the echoes, as illustrated by the high-resolution skymap in Fig. 4 observed at 19:13 LT, just a few minutes before the ionogram in Fig. 3 was recorded. The skymap, in geomagnetic coordinates, shows the locations of the F-region reflection points for frequencies between 3.8 and 5.3 MHz from within a cone of 30° around zenith; the colors indicate the Doppler shifts from -2 Hz (red) to $+2$ Hz (blue). The reflection points are distributed in the east-west direction, with Doppler frequencies of -0.4 Hz in the west to -2.0 Hz in the east. This Doppler frequency distribution suggests an upward and eastward drift velocity. The skymap measurements were repeated every 5 min to investigate the changing structure in the F-region and to determine the drift velocity components as a function of time. Figure 5 shows a subset of skymaps on 22 November 2002 for Cachimbo, spaced by 15 min, starting at 17:23 LT. Each skymap is the result of a 20-s measurement at 4 frequencies between 3 and 5 MHz. The F-layer becomes tilted by about 5° around 18:00 LT when the terminator moves over the station, and most reflection points are in the west.
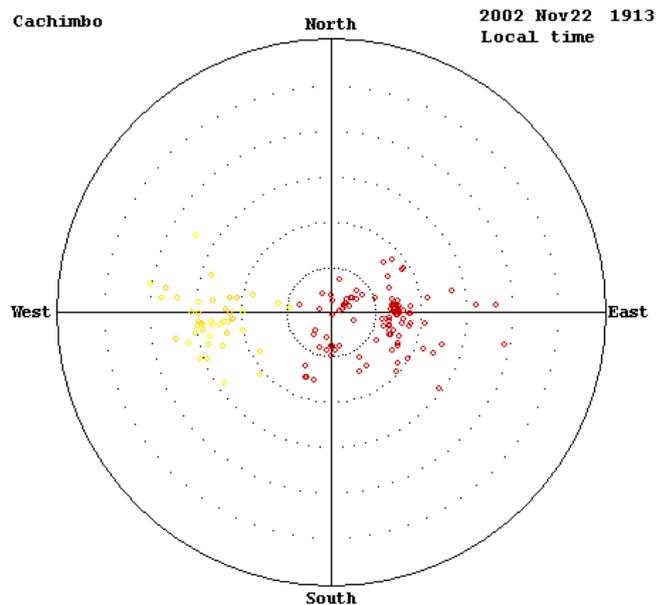

Fig. 4. Cachimbo skymap for 22 November at 19:13 LT showing echoes from the west with Doppler frequencies of -0.5 Hz, and from the east with Doppler frequencies of -2.0 Hz. The skymap displays zenith angles from 0 to 20°.
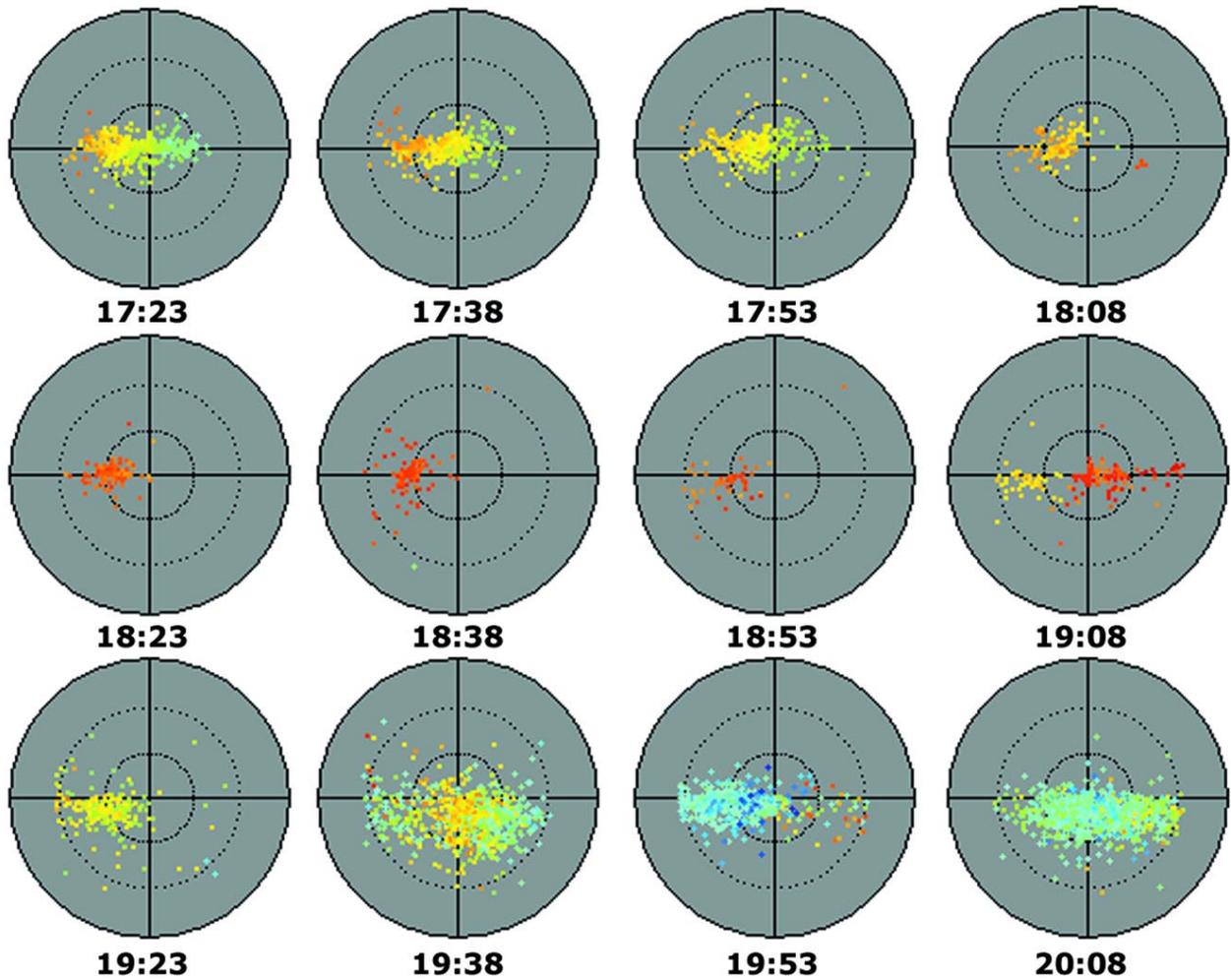
Cachimbo, 22 Nov 2002 (326), Local Time

17:23   17:38   17:53   18:08

18:23   18:38   18:53   19:08

19:23   19:38   19:53   20:08

Fig. 5. Skymaps for Cachimbo on 22 November from 17:23 LT to 20:08 LT, showing the effect of the terminator passing at ~18:00 LT.
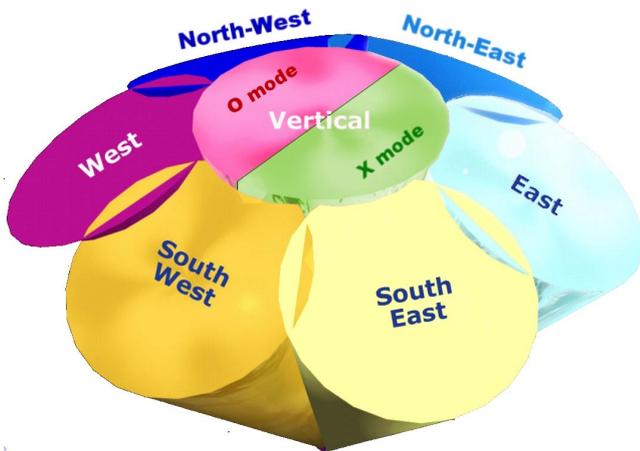

Fig. 6. Seven digitally synthesized beams for ionograms measurements.

For the ionogram display, the DPS synthesizes 7 beams pointing vertical and into six azimuthal directions at a selected fixed zenith angle, typically 20° (Fig. 6). At each sounding frequency, the 7 digital beams are calculated for each range (height bin), resulting in an ensemble of 7 ordinary (O) polarization ionograms and 7 extraordinary polarization (X) ionograms. From the Fourier spectrum at each range, only the maximum spectral component is selected and the beam direction with the maximum synthesized amplitude is assigned to the range bin in the ionogram. The Cachimbo ionogram in Fig. 3 uses different colors to indicate the directions from which the echoes are arriving, clearly showing the composite nature of the equatorial spread F ionograms, with F-region echoes arriving mainly from the west, east, and overhead in agreement with the skymap in Fig. 4. The overhead trace (O echoes red, X echoes green) is used to calculate the vertical electron density profile up to $hmF2$, superimposed on the ionogram as plasma frequency versus height (Huang and Reinisch, 2001). The topside part of the
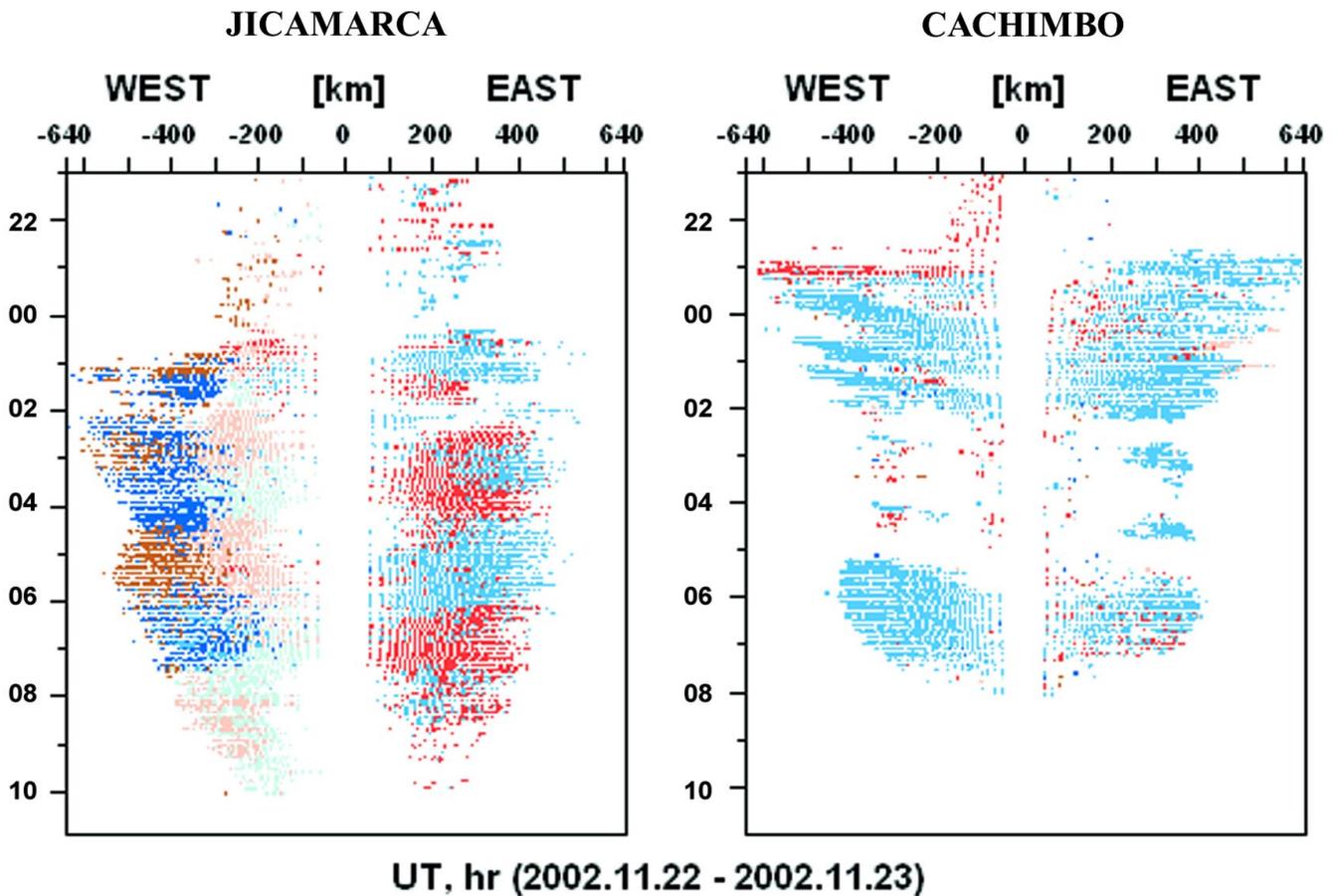
JICAMARCA   CACHIMBO

Fig. 7. Directograms for Cachimbo and Jicamarca on 22 November show the distances to the depletions observed in the ionograms.
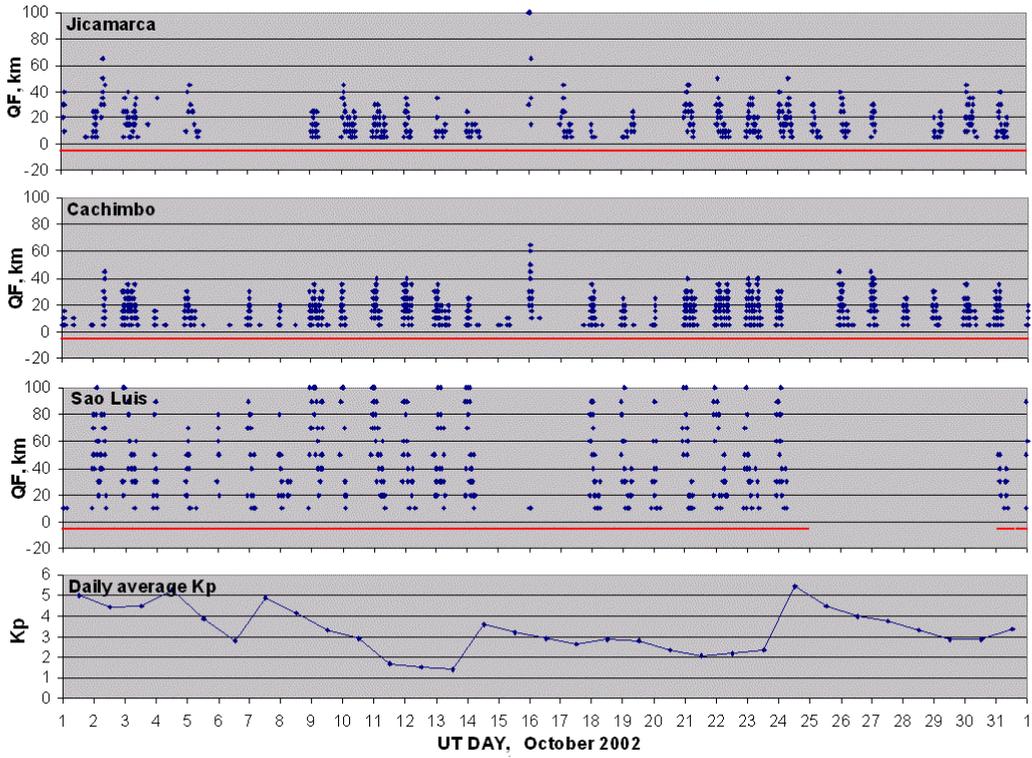
profile is derived by using the measured scale height at the F2 peak and assuming a matching α-Chapman density distribution (Reinisch and Huang, 2001).
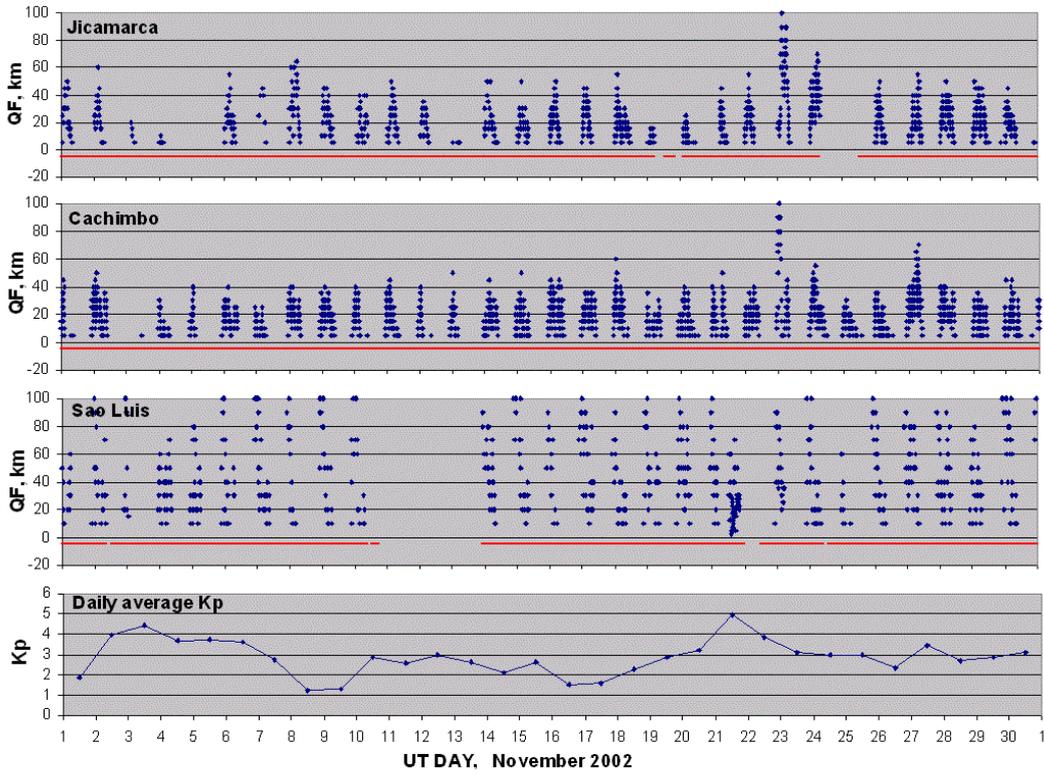
2.2 ESF development

The ionograms usually sees an approaching depletion in the west about 30 to 60 min before it arrives overhead, depending on the eastward drift velocity. Measuring the radar ranges to the vertical and oblique echoes and the direction from which the echoes arrive, it is possible to calculate the approximate horizontal distances to the depletions and plot them as a function of time in a "directogram". The directogram shows the distance to the depletions in the west and east, and suppresses the vertical echoes. Figure 7 shows the directograms for 22 November at Jicamarca and Cachimbo. As expected, Cachimbo sees the development of depletions one hour before Jicamarca detects any. These disturbances persist for about 9 h at both locations, although at a reduced level at Cachimbo from 03:00 to 05:00 UT. Eastward displacement of the ESF depletions with its irregularity is clearly seen over Cachimbo, with the displacement velocity decreasing towards post midnight hours (LT=UT−3 h). Over Jicamarca such displacement velocity is not clearly defined . This behavior might suggest the existence of a longitudinal variation

in the irregularity dynamics. It should be investigated if this longitudinal difference is the result of the moderate intensity magnetic disturbance (with ΣK_{ip}=31) that characterized this night.

We studied the occurrence of ESF at the three equatorial locations, Jicamarca in the west, Cachimbo, and Sao Luis in the east. Figure 8 shows the automatically generated display of observed range spread QF in km for a) October and b) November 2002. The ionogram cadence was 5 min at Jicamarca and Cachimbo, and 15 min at Sao Luis. A solid line underneath the data indicates that the sounder was operating. The averaged daily K_p values in the bottom panel do not reveal any apparent correlation with the ESF activity. On most of the 61 days, ESF occurrence is the same for the three stations, but on 13 days, listed in the group "some" in Table 1, one or two of the sites are quiet. Only one day, 14 October (none), had quiet conditions at all three sites. For comparison Table 1 shows one example, October 10 (all), with all sites active. The table lists the observed range spread QF, the maximum vertical drift V_z at the prereversal enhancement (Farley et al., 1986), and the maximum F2 peak height $hmF2$; the average K_p value is given for each of the days. High V_z and large $hmF2$ values are usually assumed to be preconditions for ESF development (Fejer et al., 1999) but our examples

(a)

(b)

Fig. 8. Automatically detected range spread at Jicamarca, Cachimbo, and Sao Luis for (**a**) October and (**b**) November 2003. Underlining shows the times of digisonde operation.

Table 1. Range spread, maximum vertical drift at prereversal enhancement, and the maximum F2 layer peak height for October and November 2002, grouped into all sites quiet (none), some but not all sites with ESF activity (some), and one example for all three sites active (all). A dash indicates "no data".

ESF	Date	K_p	spread F km Ji	Ca	max V_z m/s SL	Ji	Ca	max $hmF2$ km SL	Ji	Ca	SL
none	14 October	4−	0	0	0	11	32	30	504	518	538
	5 October	5+	0	0	80	16	42	29	480	567	588
	6 October	3−	0	30	90	55	45	48	581	612	557
	7 October	5−	0	20	80	56	50	37	552	614	638
	15 October	3+	100	65	0	55	50	43	545	631	620
	16 October	30	45	0	0	72	31	27	640	560	551
	19 October	3−	0	25	100	64	51	57	567	634	612
	24 October	5+	35	0	−	34	29	−	520	485	−
some	27 October	4−	0	25	−	68	37	−	510	552	−
	2 November	40	0	0	100	50	39	33	615	564	595
	3 November	4+	0	30	70	72	41	32	542	620	640
	4 November	4−	0	40	80	37	77	40	490	652	562
	12 November	30	0	50	−	50	64	−	560	673	−
	18 November	2+	0	35	100	49	39	47	715	590	627
all	10 October	30	30	40	100	70	59	35	638	648	557

Table 2. $foEs$ values at Boa Vista / Campo Grande (the larger value) at the time of the prereversal enhancement on the days without ESF at Cachimbo.

Date 2002	$foEs$ (22:00 UT) MHz	$foEs$ (22:15 UT) MHz	Time of V_z, max UT	$foEs$ (at time of V_z, max) MHz
5 October	2.4	2.3	22:00	2.4
14 October	4.7	7.9	22:10	7.7
16 October	3.9	2.6	22:05	2.6
24 October	3.4	3.8	21:55	2.4
2 November	>1.5	4.3	22:00	>1.5

show that these conditions are not sufficient. There is no systematic difference in V_z or $hmF2$ at the sites with or without ESF occurrence.

2.3 Effect of E-region conductivity

It is often speculated that the field-line integrated Pedersen conductivity is affected by sporadic E ionization at the E-region footprints of the magnetic field line through the F-region at the equator, and that Es may short circuit the electric field necessary to lift the F2 layer preventing the Rayleigh-Taylor instability to grow. The two digisondes at Boa Vista in the north and Campo Grande in the south have monitored the Es activity during the COPEX campaign close to the E-region intersection of the magnetic field line that crosses the equator at 350 km over Cachimbo. The $foEs$ critical frequency distribution for 68 nights with ESF activity peaks between 2 and 3 MHz. Figure 9 gives the distribution at 22:00 UT, and V_z at Cachimbo usually reaches its maximum values between 22:00 and 22:20 UT. The larger of the
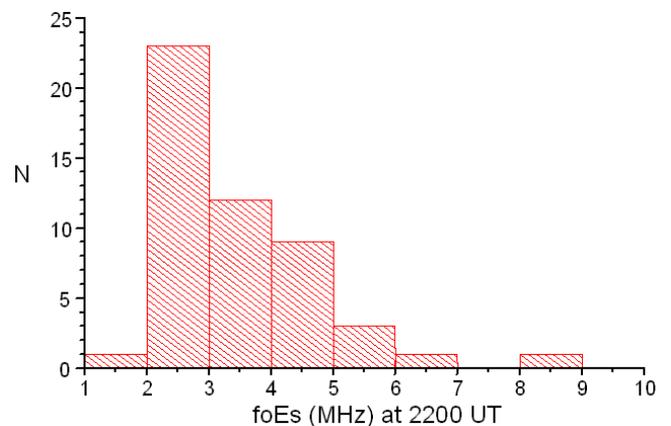

Fig. 9. Distribution of $foEs$ values at Boa Vista/Campo Grande at 22:00 UT for ESF days at Cachimbo.
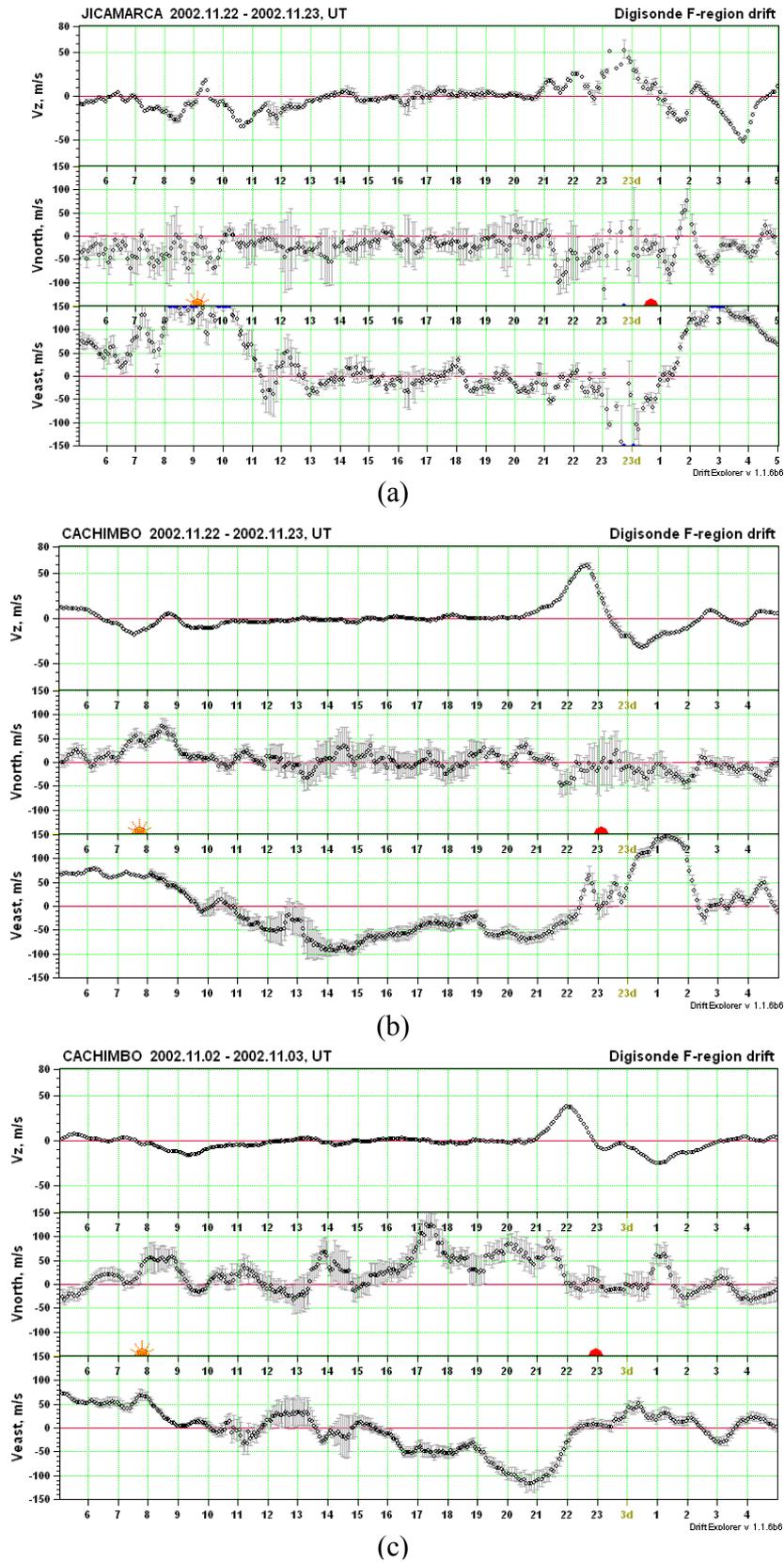
(a)

(b)

(c)

Fig. 10. Vertical, meridional, and zonal drift velocities on 22–23 November (strong ESF) at (**a**) Jicamarca and (**b**) Cachimbo, and (**c**) on 2–3 November (no ESF) at Cachimbo.

$foEs$ values observed at Boa Vista and Campo Grande was included in the statistics.

Table 2 summarizes the Es activity at Boa Vista and Campo Grande on the 5 quiet days at Cachimbo without ESF. Clearly the $foEs$ values are not any larger on these days than on the 68 disturbed days. A preliminary conclusion is that the presence or lack of sporadic E at the footprints of the magnetic field line plays no role in inhibiting or encouraging the development of equatorial spread F at the magnetic equator.

2.4 Equatorial F-region drift

For the first time, all three components of the F-region drift were simultaneously measured by the digisondes at several equatorial sites. The velocity components are calculated from the line of site velocities (Doppler frequency shifts) of the skymap source points (Reinisch et al., 1998). Figures 10a and 10b show the vertical, and the magnetic meridional and zonal velocities on 22 November, which had strong ESF activity at Jicamarca and Cachimbo. The error bars are shown in gray. For comparison, Fig. 10c shows the velocities at Cachimbo for 2 November when neither site observed ESF. On the quiet day, V_z at Cachimbo reached a value of 40 m/s compared to 65 m/s on 22 November, but this is not a particularly small value and many days with smaller values did have ESF. The maximum V_z at Jicamarca (not shown here) was about the same on these two days.

The horizontal components at Cachimbo show more significant differences on these two days. On 2 November just before sunset, the northward velocity reaches values above 100 m/s and westward velocities of 120 m/s, and very low eastward velocities prevailed after the prereversal enhancement. At this stage of our analysis we have not yet done a systematic evaluation of the horizontal plasma flow. It is likely that the gravity waves, which may be required to trigger the instabilities (Valladares et al., 1983) could be identified with the drift measurements.

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