

Plasma drift estimates from the Dynasonde: comparison with EISCAT measurements

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Abstract. Modern ionosondes make almost simultaneous measurements of the time rate of change of phase path in different directions and at different heights. By combining these ‘Doppler’ measurements and angles of arrival of many such radar echoes it is possible to derive reliable estimates of plasma drift velocity for a defined scattering volume. Results from both multifrequency and kinesonde-mode soundings at 3-min resolution show that the Dynasonde-derived *F*-region drift velocity is in good agreement with EISCAT, despite data loss during intervals of ‘blanketing’ by intense *E*-region ionisation. It is clear that the Tromsø Dynasonde, employing standard operating modes, gives a reliable indication of overall convection patterns during quiet to moderately active conditions.

Key words. Auroral ionosphere · Plasma convection · Instruments and techniques

1 Introduction

Ionosondes, operating near vertical incidence, have the potential to provide good temporal and spatial resolution at all latitudes. They are relatively inexpensive when compared with other ionospheric radar systems and are used extensively to monitor both long-term temporal and medium- to large-scale spatial variations in the atmosphere.

The use of modern ionosondes to estimate plasma drift is not new; the Doppler drift technique, using fixed-frequency soundings, was developed for the Digisonde (DGS) community in order to measure *F*-region plasma

flow in the polar cap (Buchau *et al.*, 1988; Cannon *et al.*, 1991). Verification of the Doppler drift technique by comparison with flow vectors derived from a co-located, beam-swinging incoherent-scatter radar (ISR) near Kangerlussuaq in Greenland shows mean DGS velocities to be within 50 ms⁻¹ of ISR velocities, within the uncertainties of each instrument (Scali *et al.*, 1995). Although the mean values correlate well there are a number of periods where marked discrepancies can be seen, especially in the magnetic north-south component. The authors attribute these discrepancies to the effects of horizontal gradients and temporal variations within the DGS sampling volume, aliasing in the frequency spectrum and beam-swinging errors in the ISR-derived velocities.

A similar technique, the VFIT technique, has been developed using an NOAA HF Radar operating as a Dynasonde (Grubb, 1979; Wright and Pitteway, 1979), the main difference being that multifrequency ionogram data are used rather than the two fixed frequencies of the Doppler drift technique. The VFIT (Wright and Pitteway, 1994) and Doppler drift techniques share nearly the same equations defining vector velocities from echolocation and line-of-sight Doppler measurements; the main differences are to be found at the level of data acquisition (e.g. Dynasonde ‘echoes’ and DGS ‘sources’). Instead of using $\Delta\phi/\Delta t$ directly, VFIT scales the phase measurements into a velocity:

$$V^* = \frac{c}{4\pi f} \frac{\Delta\phi}{\Delta t}, \quad (1)$$

where c is the speed of light in vacuo, f the radio frequency, ϕ the phase and t time. This tends to weight the Doppler measurements in favour of moving targets and against $\Delta\phi/\Delta t$ arising from changes in underlying ionisation. Other and more subtle differences appear in data weighting, windowing, summarising and qualifying the results.

Initial results have been published by Jarvis (1995) of Dynasonde-derived velocities over Halley, Antarctica

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using a multifrequency technique similar to VFIT; these compare favourably with simultaneous plasma flow vectors measured poleward of Halley by the PACE coherent HF radar. A direct comparison of Dynasonde-derived drift measurements with EISCAT-derived velocity vectors by Sedgemore *et al.* (1996) showed good agreement for measurements taken during quiet to moderately active conditions, with the results averaged over time-scales of around 30 min. In this paper we present results from an EISCAT comparison using an experimental Dynasonde pulse scheme which show good agreement on much shorter time-scales.

2 Dynasonde velocity measurements

Echoes are discriminated from impulsive noise in real time by a coincidence detection algorithm in the Dynasonde operating system (Wright and Pitteway, 1979; Rose, 1993); this is integrated with a multipulse scheme by which echo phases are measured in two receivers from a small six-antenna receiving array, at closely spaced times (0.01 s) and at radio frequencies differing by a few kHz. Physical parameters are then derived by a UNIPHASE algorithm (Wright, 1988). A detailed discussion of multipulse schemes can be found in Pitteway and Wright (1992) and Tsai *et al.* (1993).

With the four-pulse schemes used previously, eight phases were available to determine six parameters: stationary-phase group range R' , two components of echolocation XL (east) and YL (north), echo 'Doppler velocity' V^* , polarisation, and an average phase angle. An estimation-error parameter EP , of two degrees of freedom, was also obtained. For the data presented here an experimental eight-pulse scheme was used instead. This doubles the time-span of the pulse set with immediate benefit to the Doppler resolution (Jarvis, 1995); with its revised Δf pattern the Doppler aliasing interval is increased from 180° to 360° , and EP is estimated with ten degrees of freedom.

Individually for each echo, the line-of-sight velocity V^* is related to the velocity components V_x , V_y , V_z of a plane, totally reflecting surface located at XL , YL , ZL , where $ZL^2 = R'^2 - XL^2 - YL^2$, as:

$$V^* = \frac{XL \cdot V_x + YL \cdot V_y + ZL \cdot V_z}{R'} \quad (2)$$

With a group of $N > 3$ echoes, and the assumption that they share the same velocity vector, an estimate of the velocity vector is obtained by minimising:

$$S = \sum_{i=1}^N w_i^2 \cdot \left[V_i^* - \left(\frac{XL}{R'} \cdot V_x + \frac{YL}{R'} \cdot V_y + \frac{ZL}{R'} \cdot V_z \right)_i \right]^2 \quad (3)$$

A weight w , varying inversely with EP , is assigned to each observation. The weighted, minimised variance S is used together with diagonal elements of the covariance matrix to obtain confidence limits on the velocity components. S is used also with the weighted V^* variance $SQ = \sum_i (w_i \cdot V_i^*)^2$ to define $PV = 1 - S/SQ$:

the fraction of Doppler variance expressed by the solution. With PV we can quantify the validity of the assumption that the N echoes share the same velocity.

These steps do not in themselves lead us to favour any particular physical interpretation of the vector velocities obtained. Different interpretations, not many of which have been worked out in satisfactory detail, will apply in different situations. For example, in conditions where V^* , XL and YL behave in a manner organised by a medium-scale Atmospheric Gravity Wave (AGW), the resulting vector velocities may describe phase-velocity components of the AGW (Wright and Pitteway, 1982). For the present work a simple model is implicit for the reflecting surface already mentioned. Specular reflection is assumed to occur from the plasma-density surface appropriate to the Dynasonde measurement frequency, and this surface is assumed to advect with the plasma itself. In common with the interpretation of EISCAT F -region velocities, velocity components normal to the magnetic field are assumed to be independent of height (above 160 km) along the field line and imposed by a perpendicular electric field.

3 Instrumentation

The EISCAT experiment used in the following comparison is CP-1-K: a fixed, field-aligned beam with the remote sites intersecting at 278 km altitude. The velocity is actually determined for a single intersection height, but as the magnetic field lines are assumed to be at the same potential, each field-perpendicular component of plasma velocity should be almost constant over the whole of the F region so that the comparison with Dynasonde velocities is valid at all heights.

The Dynasonde transmits 60- μ s pulses at any frequency in the range 0.1–30 MHz and two receivers are tuned electronically to the transmitted frequency. For the measurements described in this paper, signals were received using an array of six dipole antennas arranged in the shape of a cross (Pitteway and Wright, 1992). For this experiment, B -mode soundings, each lasting around 150 s, were run every 3 min with a frequency range of 1–10 MHz, and the pulse scheme used was one with eight pulses per set in place of the usual four. Along with the 3-min B -mode sounding schedule, single kinesonde-mode soundings (K mode) with four frequencies transmitted at 3.5, 4.0, 4.5 and 5.4 MHz, using the same eight-pulse scheme, were run every 30 min. K mode resembles in some ways the DGS drift mode and is intended for high-time-resolution studies (Wright and Pitteway, 1979).

4 Velocity comparison

The EISCAT data in the following comparison were taken from the Common Programme CP-1-K of 17–19 June 1996. In this experiment data dumps occur every 5 s, although in this case the long-pulse raw data were

post-integrated over a 3-min period in order to reduce the noise errors to a reasonable level, and also to match the cycle time for the Dynasonde soundings. The integrated data were then used to derive *F*-region vector velocities at 278 km altitude with the tristatic method. Dynasonde *B*-mode soundings were separated by 3-min intervals, and all echoes with reflection frequencies between 2.8 and 8 MHz, virtual ranges between 200 and 500 km, and angles of arrival within 20° of the zenith were included in the analysis.

The eastward component of plasma velocity is shown in Fig. 1 at 3-min resolution. Looking first at the EISCAT velocities, a number of sudden westward surges can be seen at 1300, 1430 and 1600 UT, 17 June 1996, reaching a maximum of 600 m s⁻¹. At 2045 UT the plasma begins to flow eastward during another active period seen as hard-precipitation events in the EISCAT electron-density data. After 2300 UT the flow rate decreases and remains stable at around zero for 14 h, after which it turns westward again following the same pattern as on the previous day. Following the next flow reversal (Harang discontinuity) the eastward velocity increases to a maximum of 800 m s⁻¹ during a highly active period. EISCAT noise errors in this component range from 60 to 90 m s⁻¹ except during the final, disturbed period where they rise to around 120 m s⁻¹. In comparison, the Dynasonde errors, derived as a measure of the goodness of the fit, are more variable ranging from 15 to 200 m s⁻¹.

Throughout the whole experiment, lasting 45 h in all, there are sufficient Dynasonde echoes for the VFIT analysis, apart from two periods of ‘blanketing’ corresponding to the time of the major precipitation events seen in the magnetometer and EISCAT electron-density data. Blanketing occurs when an intense *E* layer blocks out all, or nearly all *F*-region echoes. The magnetic *H* component for 17 June indicated a number of disturbances, the first three of which correspond to the precipitation events seen during the first 6 h of the

experiment. The fourth matches with the hard-precipitation event that occurred shortly before 2100 UT. The magnetogram for 18 June showed very little activity until around 2315 UT when a clear substorm signature could be seen with a 500-nT bay in the *H* component.

After the first 3 h of small-amplitude velocities there is good agreement between EISCAT and Dynasonde *B*-mode velocities as the plasma flows westward, although the Dynasonde underestimates the peak magnitude at around 1600 UT. There is then an anomalous feature in the Dynasonde results seen as a 500 m s⁻¹ westward excursion at 1815 UT during a period where EISCAT sees the plasma moving westwards at 300 m s⁻¹. The ionograms from this time show a number of structured *E*-region echoes from within the Dynasonde sampling volume which may be the cause of this discrepancy.

At around 2100 UT there is a hard-precipitation event, and here the velocity fitting becomes impossible owing to insufficient echoes. The *F*-region signal returns at 2300 UT but with few echoes the velocity fits oscillate wildly for about 1 h. From 0000 until 0200 UT, 18 June, there is a significant discrepancy between EISCAT and the Dynasonde which again may be due to the influence of *E*-region echoes in the sampling volume. From 0200 to 1400 UT, the ionosphere is quiet and inactive and the east-west velocity is effectively zero throughout this period. At 1400 UT there is a small, 150 m s⁻¹ westward excursion lasting for 2 h. The velocity then reverses for a short while before turning westward again reaching a maximum of 500 m s⁻¹ at 1815 UT (Fig. 2). Throughout this time EISCAT and the Dynasonde are in close agreement, and this continues up to the Harang discontinuity and the onset of a major substorm at 2315 UT when blanketing occurs.

The *K*-mode velocities, shown also in Fig. 1, tend to follow the *B*-mode values. There is one interesting difference between 0000 and 0200 UT, 18 June. This corresponds to the period where there is a sustained discrepancy between EISCAT and the Dynasonde. It is

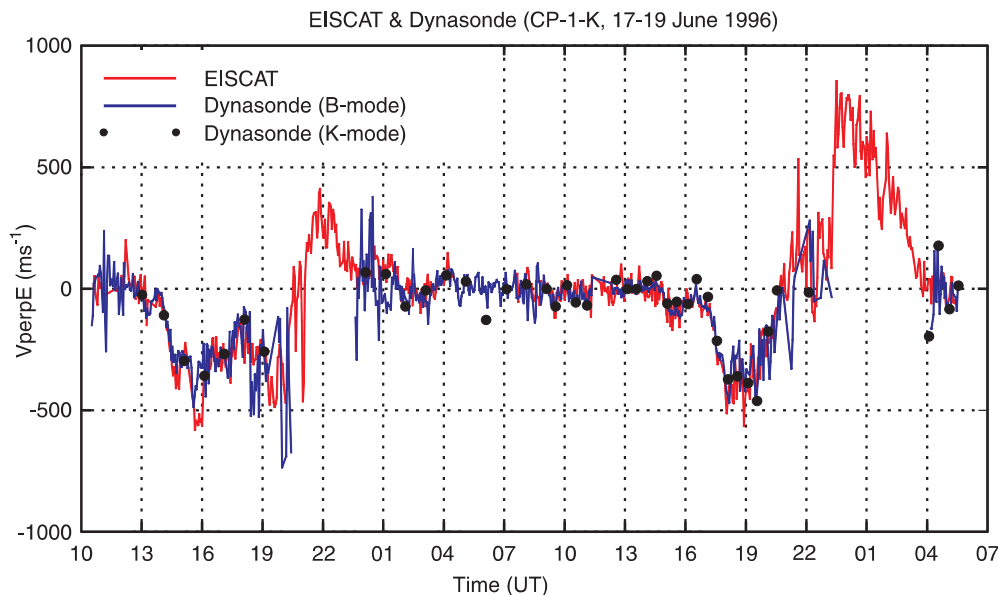


Fig. 1. EISCAT and Dynasonde *F*-region velocities (3-min resolution) from the experiment CP-1-K, 17–19 June 1996. The eastward component perpendicular to the magnetic field is shown. EISCAT noise errors range from 60 to 120 m s⁻¹, and Dynasonde errors from 15 to 200 m s⁻¹. There is a gap in the EISCAT data between 0430 and 0730 UT, 18 June

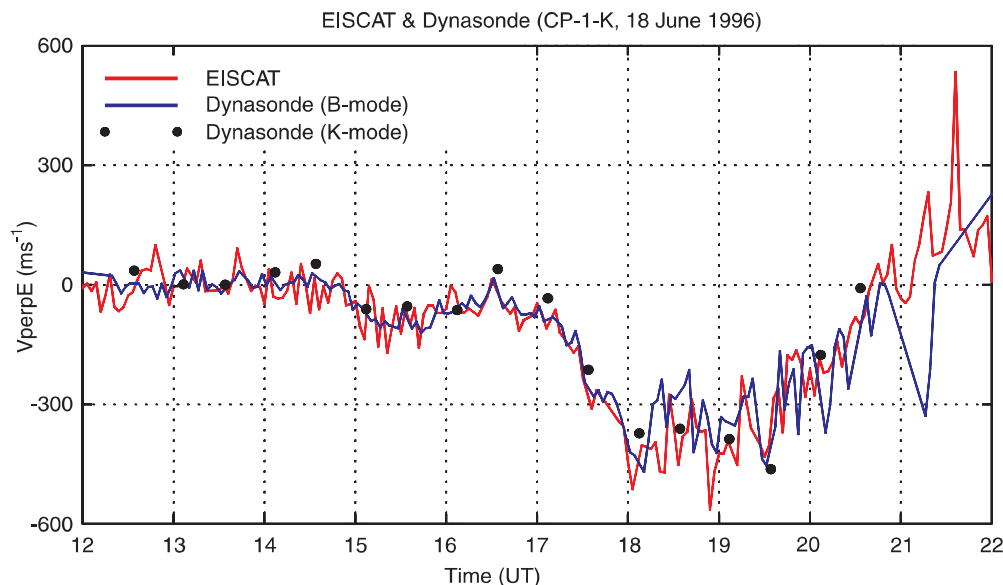


Fig. 2. EISCAT and Dynasonde *F*-region velocities (3-min resolution) from the experiment CP-1-K, 18 June 1996. The eastward component perpendicular to the magnetic field is shown, from 1200 to 2200 UT

difficult to interpret the discrepancy between *B*- and *K*-mode velocities at around 0630 UT (18 June) given the lack of EISCAT data at this time. Apart from these two instances the sounding modes give broadly similar results. A more rigorous test would involve an experiment which alternates *B*- and *K*-mode soundings at 6-min overall resolution. In general, however, the time-resolution advantage that *K* mode has over ionogram-type soundings is outweighed by the far greater amount of information contained within *B*-mode soundings.

The magnetic north-south component of velocity (Fig. 3) displays little variation; the magnitudes involved are rather small and the error bars proportionately large. There is, however, one (southward) feature of peak magnitude around 100 m s^{-1} at around 1900 UT, 18 June, that stands out as significant given the EISCAT

noise errors of $40\text{--}70 \text{ m s}^{-1}$. Despite the small magnitudes and the scatter in Dynasonde velocities it is still possible to see a positive correlation between the two distributions.

This paper has concentrated on horizontal plasma drift. VFIT, however, may be used to produce a three-dimensional vector fit. In this comparison the Dynasonde-derived field-aligned component is in good agreement with EISCAT, at around zero on average, but there are a small number of anomalously large values which could in future be used to identify ‘bad’ fits once an acceptable deviation from the mean was decided upon. It may be worth using such a restriction on the parallel component, although with the danger of rejecting genuinely large field-aligned flows. Another improvement to the technique would involve estimating

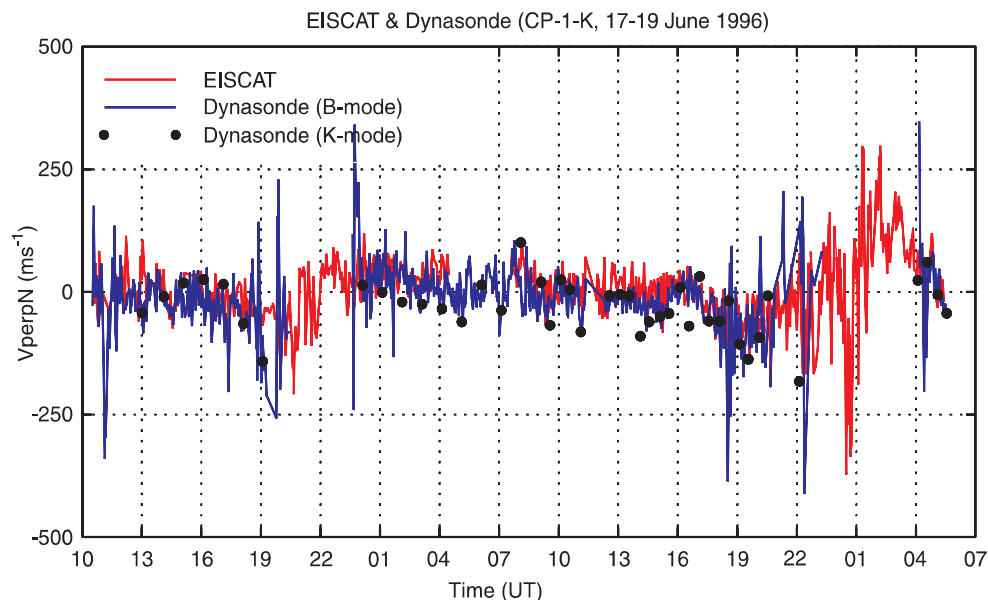


Fig. 3. EISCAT and Dynasonde *F*-region velocities (3-min resolution) from the experiment CP-1-K, 17–19 June 1996. The northward component perpendicular to the magnetic field is shown. EISCAT noise errors range from 40 to 70 m s^{-1} , and Dynasonde errors from 6 to 80 m s^{-1} . There is a gap in the EISCAT data between 0430 and 0730 UT, 18 June

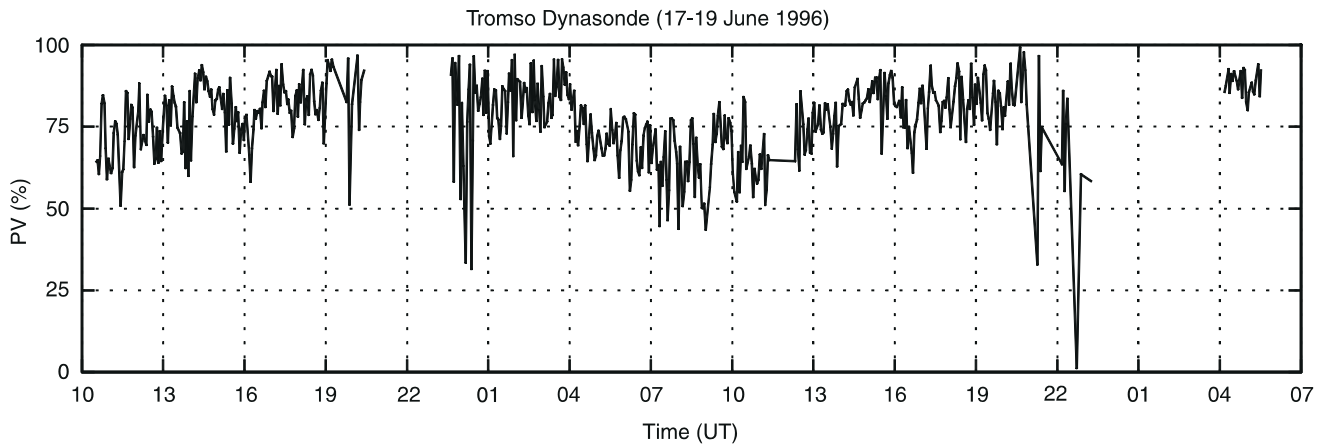


Fig. 4. Dynasonde percentage of Doppler variance (PV) expressed by the vector velocities shown in Figs. 1 and 3, during the experiment CP-1-K, 17–19 June 1996

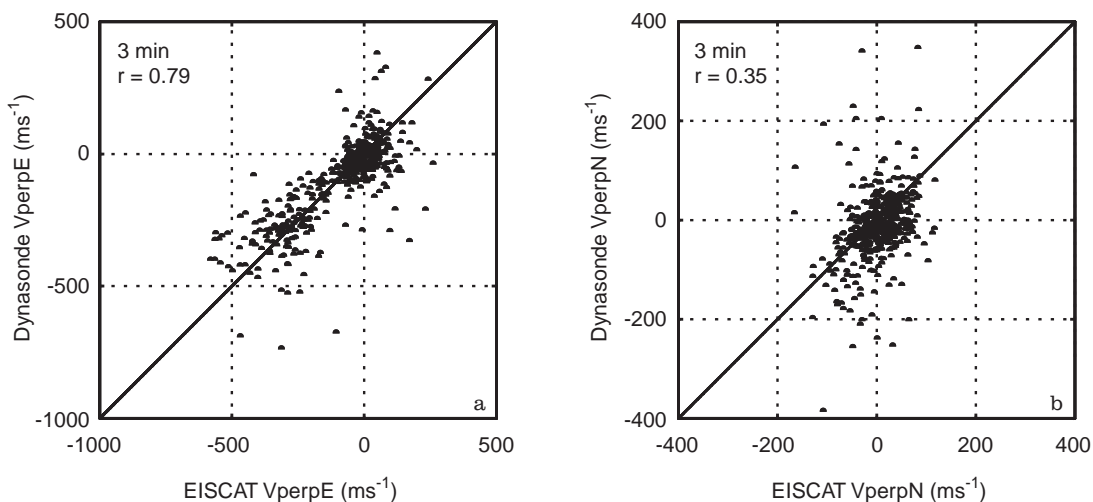


Fig. 5a, b. Correlation between EISCAT and Dynasonde F-region velocities (3-min resolution) during the experiment CP-1-K, 17–19 June 1996; **a** eastward and **b** northward components perpendicular to the magnetic field. In each case a line of unit slope is plotted for reference

horizontal gradients in velocity across the Dynasonde field of view by fitting velocity vectors to different quadrants.

Good agreement between EISCAT and Dynasonde B -mode velocities is reflected in the high PV throughout the course of the experiment (Fig. 4). This seldom falls below 60% and during the periods of close agreement it is around 80%. This is around 10% higher on average than for the four-pulse B -mode soundings used in the previous comparison. The few periods of low PV correspond to the periods leading up to blanketing. In those cases where there is no tristatic ISR to provide a true value of plasma drift velocity, it is important to have some criterion derived from Dynasonde data which indicates whether the derived velocities are real; the evidence so far is that PV provides just such a criterion.

When the correlation between the velocities from EISCAT and Dynasonde-derived data is calculated there is an approximately linear increase as the results are averaged over time-periods of 3, 6, 9 min and so on up to 54 min. In the previous comparison by Sedgemore

et al. (1996) it was necessary to average over periods of 30 min or more before a reasonably high correlation was obtained. With the present measurements of east-west plasma drift, data with a time-resolution of 3 min give a correlation coefficient of 0.79 and this increases to 0.85 when the data are averaged over 54 min. In the auroral zone north-south velocities are generally much smaller and hence the correlation is more seriously affected by random noise errors. Nevertheless, there is a 0.35 correlation between the north-south velocities measured with a 3-min resolution, and this increases to 0.50 for 54-min averages. Figure 5 shows the correlation at 3-min resolution, for both the eastward and northward vector components.

5 Conclusions

The critical test for the Dynasonde is whether it can provide reliable values of plasma drift velocity in a stand-alone situation. This is particularly important at

high latitudes where large and variable plasma drift occurs as a result of magnetospheric convection. For example, Dynasonde measurements of plasma drift above Longyearbyen could make a valuable contribution to the interpretation of data from the EISCAT Svalbard Radar (ESR), which at present is a monostatic system and hence cannot be used to measure velocities under conditions where rapid variations invalidate beam-swinging methods.

The first published comparison of EISCAT and Dynasonde velocities by Sedgemore *et al.* (1996) showed poor agreement for short-term variations but a high correlation between measurements of plasma velocity on time-scales of 30 min or more. The results presented here show that small-scale velocity variations can be resolved by the Dynasonde using standard operating modes at high time-resolution.

A major problem with using ionosondes to measure plasma velocity in the auroral zone is data gaps due to ratio absorption and intense *E* layers which blanket echo returns from the *F* region. In the polar cap these effects are far less common and velocities derived by Scali *et al.* (1995) using a DGS system based in Greenland show very few such gaps. This is a good argument for siting a Dynasonde next to the ESR, which for much of the time sits on the northern boundary of the auroral oval.

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