



# The Sofia University Atmospheric Data Archive (SUADA)

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**Abstract.** Atmospheric sounding using the Global Navigation Satellite Systems (GNSS) is a well-established research field in Europe. At present, GNSS data from 1800 stations are available for model validation and assimilation in state-of-the-art models used for operational numerical weather prediction centres in Europe. Advances in GNSS data processing make it possible also to use the GNSS data for climatic trend analysis, an emerging new application. In Bulgaria and southeastern Europe, the use of GNSS for atmospheric sounding is currently under development.

As a first step, the Sofia University Atmospheric Data Archive (SUADA) is developed. SUADA is a user-friendly database, and includes GNSS tropospheric products like zenith total delay (ZTD) and derivatives like vertically integrated water vapour (IWV), as well as observations from radiosonde (RS) and surface atmospheric data. Archived in SUADA are (1) GNSS tropospheric products (over 12 000 000 individual observations) and derivatives (over 55 000) from five GNSS processing strategies and 37 stations for the period 1997–2013, with temporal resolutions from 5 min to 6 h, and (2) radiosonde IWV data (over 6000 observations) for station Sofia (1999–2012).

Presented are two applications of the SUADA data for the study of long- and short-term variations of IWV over Bulgaria during the 2007 heatwave and intense precipitation events in 2012.

## 1 Introduction

Atmospheric water vapour is the most abundant greenhouse gas involved in the climate feedback loop. As the temperature of the Earth's surface and atmosphere increases, so does the moisture-holding capacity of the atmosphere

and atmospheric water vapour is expected to increase in a warmer climate. The evidence is now indisputable (Dessler and Sherwood, 2009) that water vapour increase adds one degree Celsius to global warming for every one degree through greenhouse gas emissions. Traditionally, the long-term water vapour trends have been estimated using the global radiosonde (RS) data (Gaffen et al., 1992; Ross and Elliott, 1996, 2001). In the last decade, the use of GNSS tropospheric products has been employed to verify the RS trends, which are not homogeneous due to sensor changes. Gradinarsky et al. (2002) estimate the water vapour trend using GNSS data over Scandinavia and finds it to be positive, with  $0.1\text{--}0.2\text{ mm yr}^{-1}$  for the 1993–2000 period. They also report that winter trends are larger, than summer trends for the southern part of the region and opposite for the northern part. Nilsson and Elgered (2008) analysed data for Sweden and Finland for the period 1993–2006 and confirmed the previous findings. Ning (2012) estimates the linear IWV trends, using 15 year global GNSS data set. The trends are found to be in the range  $-1.65$  to  $+2.32\text{ kg m}^{-2}\text{ decade}^{-1}$  and the estimated trend uncertainties is of the same order, varying from  $0.21$  to  $1.52\text{ kg m}^{-2}\text{ decade}^{-1}$ .

In addition to climate monitoring, another application of the GNSS tropospheric products is to study development of convective clouds with intense precipitation. Recent studies such as Graham et al. (2012) in the Bernese Alps in Switzerland and van Baelen et al. (2011) in the Black Forest region in Germany used GNSS-derived water vapour to study isolated convection development. The first study shows that large transfers of air and water vapour occur from the Swiss plain to the mountains, with up to 50% increase in GNSS integrated water vapour (IWV) at the Alpine stations, coincident with strong airflow convergence at the same location (Graham et al., 2012). During the intense observation

campaign in the region of the Black Forest mountains in the summer of 2007 van Baelen et al. (2011) study the relationship between water vapour evolution and the life cycle of precipitation systems. They show that (1) frontal systems seem to develop preferentially where the largest amount of water vapour is available, and that (2) water vapour has a predominant role as a precursor for the initiation of local convection. Accumulation of water vapour on the crest of the orography leads to ridge convection and its passing over the orography triggers lee-side convection. De Haan (2008) shows the value of the real-time GNSS-IWV maps for nowcasting by examining two cases studies. The first case is a severe thunderstorm on 8 June 2007 and the second case are two thunderstorm events on 20 July 2007. In both cases the convergence of moist air contains information on the location of developing thunderstorms. For the analysis of the case studies several methods are used – radar observations, GPS IWV maps and surface wind observations. The conclusion is that the real-time GPS IWV maps are of good quality and can be helpful for nowcasting of severe thunderstorms. A nine-year study in northern Spain, conducted by Seco et al. (2012), reports that rain events are usually from atmospheric low pressure systems and water vapour entries are caused by Atlantic disturbances. They identify three precipitation patterns associated with different behaviour of water vapour during the year. Winter and summer months tend to have characteristic water vapour patterns, while spring and autumn are without clear patterns.

Atmospheric water vapour is also one of the most variable and important parameters for numerical weather prediction (NWP) and forecasting, but is under-sampled in current operational meteorological observing systems. Application of GNSS tropospheric products for NWP was the focus of EU projects WAVEFRONT, MAGIC, TOUGH and COST Action 716. The main achievements of those projects are (1) the set-up of near-real-time processing with data available with a 90 min time lag (Elgered et al., 2005), (2) quality control of the GNSS tropospheric products, and (3) set-up of a real-time GNSS archive. Following their success, the application of GNSS for NWP is now a well-established technique in Europe. Since 2005, E-GVAP (EIG EUMETNET GNSS Water Vapour Programme, EGVAP project, 2014) has been in charge of the collection and quality control of operational GNSS tropospheric products for NWP in Europe. More than 12 E-GVAP analysis centres (ACs) produce GNSS tropospheric products for over 1800 ground-based GNSS stations in Europe, with a target latency of 90 min, and make them available to national meteorologic services. The state of the art is data assimilation of hourly updated ZTD in NWP models. While the production, exploitation and evaluation of operational GNSS tropospheric products for NWP is well established in northern and western Europe, it is still an emerging research field in eastern and southeastern Europe.

This manuscript is a first step towards the application of ground-based GNSS tropospheric products in operational

meteorological and climate observing systems in Bulgaria/southeastern Europe. As a platform for archiving data on an ongoing basis, the Sofia University Atmospheric Data Archive (SUADA) was developed. Currently, SUADA includes GNSS tropospheric products and derivatives like IWV from 5 processing strategies and total of 37 stations for the period 1999–2013. IWV from the radiosonde station Sofia in Bulgaria is also archived in SUADA (About SUADA, 2014). The envisaged applications include (1) cross-validation of ground-based and satellite observations, and derivation of systematic biases, (2) validation of numerical models used for research and numerical weather prediction (NWP), (3) a study of water vapour distribution in Bulgaria/southeastern Europe, (4) detection of long-term water vapour trends in Bulgaria/southeastern Europe and links to heatwaves, droughts and changes in the pathway of the Atlantic cyclones, and (5) a study of how well state-of-the-art climate models, notably the one participating in the Intergovernmental Panel on Climate Change (IPCC) AR5 assessment, simulate the present climate of Bulgaria/southeastern Europe. SUADA was developed in close collaboration with the Institute of Applied Physics, University of Bern (IAP-UniBe). Since 2001, IAP-UniBe has operated the STARTWAVE (STudies in Atmospheric Radiative Transfer and Water Vapour Effects) database. The STARTWAVE database (Morland et al., 2006a) was funded within the framework of the NCCR Climate project 2001–2013 (NCCR, 2014). The STARTWAVE database was used for studies covering (1) validation of two operational NWP models used in MeteoSwiss (Guerova et al., 2003), (2) comparison with the 40 year reanalysis data (ERA40) of the European Centre for Medium Range Weather Forecasting (Morland et al., 2006b), and (3) evaluation of the ECHAM5 climate model. STARTWAVE data was used for instrumental intercomparisons, the major result being detection of daytime bias in the radiosonde observations (Guerova et al., 2005) as well as instrumental problems at the high-altitude Jungfrauoch station (Guerova et al., 2003). In Switzerland, a consistent positive IWV trend was found by Morland and Maetzler (2007). Analysing IWV during the 2003 heatwave summer, Guerova and Morland (2008) report a large positive IWV anomaly in June and a large negative anomaly during the heatwave in August.

This paper is organised as follows: Sect. 2 presents the SUADA structure and data sets; the GNSS meteorology method is presented in Sect. 3; and Sect. 4 presents two case studies of the application of GNSS meteorology for the short-term and long-term variation of the water vapour in Bulgaria. Conclusions are given in Sect. 5.

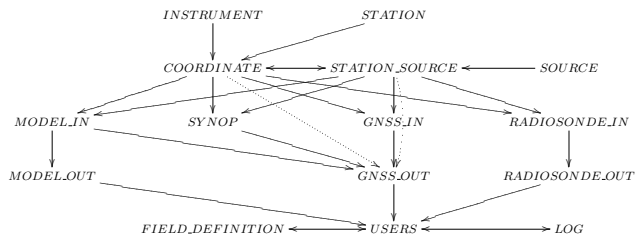


Figure 1. SUADA data structure and data flow.

## 2 Sofia University Atmospheric Data Archive (SUADA)

### 2.1 SUADA structure

SUADA is developed using Structured Query Language (SQL) for relational database management systems (Codd, 1970). The SUADA tables are structured as peers, with additional relations between them as shown in Fig. 1.

In Fig. 1, the SUADA tables are presented in groups. The first group of tables are “Information tables”: INSTRUMENT, STATION, COORDINATE, STATION\_SOURCE and SOURCE (rows 1 and 2 in Fig. 1). The second group of tables are the “Primary tables”: MODEL\_IN, SYNOP, GNSS\_IN and RADIOSONDE\_IN (row 3 in Fig. 1). The third group of tables are the “Secondary tables”: MODEL\_OUT, GNSS\_OUT and RADIOSONDE\_OUT (row 4 in Fig. 1). The last group of tables are the “Information tables for the web portal”: FIELD\_DEFINITION, USERS and LOG (row 5 in Fig. 1). A short description of the SUADA tables is given in Table 1. The tables are also accessible from the SUADA web portal.

### 2.2 SUADA data sets

#### 2.2.1 GNSS data sets

Currently, SUADA has five GNSS data sets processed with different software and strategies. As seen in Table 2, the GNSS data offer a high temporal resolution from 5 min to 6 h, and the IGS station in Sofia, Bulgaria (SOFI, marked by the red pointer in Fig. 2) has been available since 2001. The GNSS data sets are discussed below.

The first SUADA GNSS data set is IGS-repro1. In 2008, the International GNSS Service (IGS) initiated global GNSS data reprocessing campaign (IGS-repro1, Rebischung et al., 2012; Byun and Bar-Sever, 2009). Nine IGS analysis centres contributed to reanalysing the GPS data collected by the IGS global permanent network since 1994 in a fully consistent way using the latest models and methodology. The IGS-repro1 campaign started after adoption of a new set of antenna phase centre calibrations for 65 out of 232 sites of the global IGS network. Archived in SUADA are IGS-repro1 tropospheric products for station SOFI for the period 2001–2007. The ZTD and gradients are processed

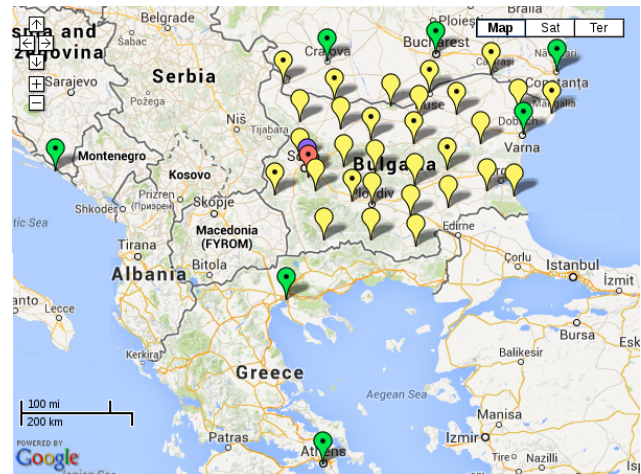


Figure 2. SUADA GNSS stations in Bulgaria/southeastern Europe.

with JPL GIPSY/OASIS software, and are available every 5 min for the period 1997–2007. The estimation approach is as follows. (1) Fixed orbits and clocks: IGS Final Re-Analyzed Combined (1995–2007) and IGS Final Combined 2008–Current. (2) Earth orientation: IGS Final Re-Analyzed Combined (1995–2007) and IGS Final Combined (2008–Current). (3) Transmit antenna phase centre map: IGS standards. (4) Receiver antenna phase centre map: IGS standards. (5) Elevation angle cut-off:  $7^\circ$ , (6) mapping function (hydrostatic and wet): GMF (Boehm et al., 2006a). (7) Data arc: 24 h. (8) Data rate: 5 min. (9) Estimated parameters: station clock (white noise), station position, wet zenith and (10) delay ( $3 \text{ cm h}^{-1}$  random walk), delay gradients ( $0.3 \text{ cm h}^{-1}$  random walk), phase biases (white noise).

The second SUADA GNSS data set is CODE-repro2 (Dach et al., 2009). This is the Center for Orbit Determination in Europe (CODE), at the Astronomical Institute of the University of Bern (CODE Analysis Strategy Summary, 2014), contribution for the second IGS reprocessing campaign (Meindl et al., 2011) initiated in 2013. In SUADA are archived CODE-repro2 tropospheric products with 2 h resolution for SOFI for the period 2001–2010. GNSS data (GPS and Glonass) is processed with Bernese GNSS Software v. 5.3 using (1) ITRF2008 reference frame, (2) elevation cut-off angle  $3^\circ$ , (3) ECMWF-based hydrostatic delay mapped with hydrostatic VMF1 (Boehm et al., 2006b). In addition to SOFI station, six European IGS stations are also archived: Zimmerwald (ZIMM), Switzerland; Onsala (ONSA), Sweden; Ondrejov (GOPE), Czech Republic; Medicina (MEDI), Italy; Matera (MATE), Italy; and Potsdam (POTS), Germany.

The third SUADA GNSS data set is produced by European Reference Frame (EUREF). EUREF is a European network operating since 1995 with the objective of providing a standard precise GNSS-based reference system for Europe. Since June 2001, tropospheric parameters have been estimated, by EUREF local analysis centres, on a weekly basis

**Table 1.** SUADA table names and short descriptions.

Table name	Short summary
INSTRUMENT	Instrument name and identification number
STATION	Station name
COORDINATE	Coordinates of the GNSS, synop and radiosonde stations
STATION_SOURCE	Station source information (either instrument or method)
SOURCE	Contact information of SUADA data providers (name, institution, telephones, etc.)
MODEL_IN	Numerical weather prediction (NWP) model data
SYNOP	Surface observations from the network of the National Institute of Meteorology and Hydrology
GNSS_IN	Tropospheric products from ground-based GNSS networks or individual stations
RADIOSONDE_IN	Data from the radiosonde network or individual stations
MODEL_OUT	Processed NWP model data (IWV or other)
GNSS_OUT	Processed GNSS data (IWV, ZHD or other)
RADIOSONDE_OUT	Processed radiosonde data (IWV)
FIELD_DEFINITION	List of abbreviations used in the SUADA tables
USERS	Contact information about SUADA data users (external and internal)
LOG	User log-in history

(post-processing mode EUREF-post), with a two-hourly sampling rate for more than 200 GNSS tracking stations of the permanent EUREF network (EUREF tropospheric delays, 2014). On the Balkan Peninsula, there are 15 EUREF stations: 5 stations each in Greece and Romania and 1 station in Turkey, Croatia, Macedonia, Slovenia and Bulgaria, totaling 15 stations. Bulgarian station SOFI has been part of the EUREF permanent network since 1997. In SUADA are uploaded SOFI tropospheric products from 2001 to 2004 processed by the BKG (Bundesamt für Kartographie und Geodäsie) analysis centre in Germany. BKG produces daily tropospheric solutions using fixed coordinates from weekly solution with Bernese software, 10° elevation cut-off angle and elevation dependent weighting. No a priori tropospheric model is used but the zenith total delay is estimated at 1 h intervals for each station and the mapping function is Dry Niell (BKG – EUREF Local Analysis Centre, 2014; Niell, 1996).

The fourth SUADA GNSS data set is provided by the ZenitGEO (Zenitgeo, 2014) private company. Since 2009, the company has operated a GNSS network with 30 GNSS stations, evenly distributed over Bulgaria (marked by the yellow pointers in Fig. 2). ZenitGEO processes the GNSS data and provides tropospheric products with a very high temporal resolution of 5 min (300 s). Currently, IWV is derived for 11 stations (marked by the yellow pointers with dots in Fig. 2), namely Vidin, Oryahovo, Lovech, Veliko Tarnovo, Ruse, Razgrad, Silistra, Shabla, Kyustendil, Pazardzhik and Sliven. It is to be noted that 8 of them are in Northern Bulgaria. The high temporal resolution of the GNSS product is degraded due to the low temporal resolution of the meteorological data set (see Sect. 2.2.3); therefore, in the near future, use of the NWP model data will be considered (see Sect. 2.2.4). This will also allow the spatial resolution for Bulgaria to be increased.

The fifth SUADA GNSS data set is a targeted processing performed by Keranka Vasilleva (Balkan) for the period 19–26 July 2007. GPS data from 19 GNSS permanent stations (AUT1, NOA1, BUCU, COST, DUBR, GLSV, GRAZ, MATE, ORID, PENC, POLV, ROZH, SOFI, Sulp, MIKL, WTZR, ZIMM, VARN, CRAI) from central and eastern Europe were processed with the Bernese software, version 5.0. Sixteen of them are IGS and EUREF stations. Seven sessions of 24 h have been created. For each session, hourly station coordinates and ZTD are estimated. The troposphere model used is Saastamoinen dry model (Saastamoinen, 1972) with Niell dry mapping (Niell, 1996) and tilting gradient model. Corrections to the introduced zenith values are estimated and the ZTD and gradients are obtained. Tropospheric products for the stations in southeastern Europe: Sofia (SOFI), Dubrovnik (DUBR), Athens (NOA1), Thessaloniki (AUT1), Craiova (CRAI), Constanta (CONS), Bucharest (BUCU) and Varna (VARN) (marked by the green pointers in Fig. 2) are uploaded in the SUADA.

### 2.2.2 Radiosonde data set

Atmospheric sounding using a radiosonde is a well-established method approved by the World Meteorological Organization (WMO). Radiosonde is widely adopted for measurements of vertical profiles of temperature, pressure, humidity, wind speed and direction. At station Sofia, Bulgaria, routine daily sounding are performed at 12:00 UTC. The station is operated by the Central Aerological Observatory at the National Institute of Meteorology and Hydrology (NIMH). Since 2005, a VAISALA RS92KL probe has been used. The relative humidity sensor has measurement range between 0 and 100 %, resolution 1 % and total uncertainty in sounding 5 %.

**Table 2.** GNSS data sets as of 1 October 2013.

Data set name	Tropos. product	Available	Number of stations	Observation frequency	Number of observations
IGS-repro1	ZTD	Jul 1997–Dec 2007	1	5 min	823 919
IGS-repro1	ZHD	Jan 2001–Dec 2007	1	3 h	16 619
IGS-repro1	ZWD	Jan 2001–Dec 2007	1	3 h	16 619
IGS-repro1	IWV	Jan 2001–Dec 2007	1	3 h	16 619
CODE-repro2	ZTD	Jan 2001–Dec 2010	7	2 h	411 306
CODE-repro2	ZHD	Jan 2001–Dec 2010	7	6 h	74 943
CODE-repro2	ZWD	Jan 2001–Dec 2010	7	6 h	74 943
CODE-repro2	IWV	Jan 2001–Dec 2010	7	6 h	74 943
EUREF-post	ZTD	Apr 2001–Nov 2004	1	1 h	23 880
EUREF-post	ZHD	Apr 2001–Nov 2004	1	3 h	6539
EUREF-post	ZWD	Apr 2001–Nov 2004	1	3 h	6539
EUREF-post	IWV	Apr 2001–Nov 2004	1	3 h	6539
ZenitGEO	ZTD	Nov 2011–May 2013	30	5 min	11 473 034
ZenitGEO	ZHD	Nov 2011–May 2013	11	3 h	23 233
ZenitGEO	ZWD	Nov 2011–May 2013	11	3 h	23 233
ZenitGEO	IWV	Nov 2011–May 2013	11	3 h	23 233
Balkan	ZTD	19–25 Jul 2007	8	1 h	1160
Balkan	ZHD	19–25 Jul 2007	8	3 h	763
Balkan	ZWD	19–25 Jul 2007	8	3 h	763
Balkan	IWV	19–25 Jul 2007	8	3 h	763

The radiosonde is widely used for intercomparison with GNSS-derived integrated water vapour (IWV). To compute the IWV from the radiosonde profiles (RS–IWV), the following equation is used:

$$\text{IWV} = \frac{1}{\rho_w} \int_{h_0}^{h_{\text{top}}} \rho_{\text{wv}}(h) dh, \quad (1)$$

where  $h_0$  is the altitude of the station, where the probe is released,  $h_{\text{top}}$  is the maximum achieved height by the probe during sounding,  $\rho_w$  is the density of water,  $\rho_{\text{wv}}$  is the density of water vapour. IWV is measured in millimetres (Guerova et al., 2003). A total of 6376 radiosonde IWV data for the period 1997–2012 is archived in SUADA.

The collection of the radiosonde data incurs substantial operational costs, which limits the temporal and spatial resolution of this observing system. Radiosonde data is a long-term observation, with time series of over 50 years, which is suitable for global climatic trend analysis, but which requires careful quantification of possible systematic biases. Studies in Japan (Ohtani and Naito, 2000), Switzerland (Guerova et al., 2005), and France (van Baelen et al., 2005) report bimodal distributions of the GPS and radiosonde residuals. Guerova et al. (2003) report an isolated case of IWV overestimation by the radiosonde related to the passage through low stratus clouds.

### 2.2.3 Surface observation data set

Surface observations of (1) pressure, (2) 2 m temperature, (3) 10 m wind speed and direction, (4) precipitation, (5) cloud cover and (6) current weather are archived in SUADA. The measurements are from the surface observation network (synop) of the National Institute of Meteorology and Hydrology (NIMH) in Bulgaria. The surface data is collected manually every 3 h starting at 00:00 UTC. The data is available from OGIMET weather information server (Ogimet Weather Information Service, 2014). In addition, surface observations from 3 stations in Romania (Constanta, Craiova and Bucharest) with hourly update and measurements by automatic weather stations are saved. The surface data is used for derivation of IWV from the GNSS tropospheric products as described in Sect. 3. The frequency of the surface observations is a limiting factor in obtaining high temporal resolution of water vapour. Often the surface data is not collocated with the GNSS station and altitude corrections are applied, which reduce the quality of the product.

### 2.2.4 Numerical weather prediction (NWP) model data set

As part of the SUADA project, the Weather Research and Forecasting (WRF) model version 3 (WRF, 2014) has been installed on the Sofia University PhysOn cluster (PHYSON, 2014). Since February 2013, the model has computed daily

forecasts of temperature and precipitation for Bulgaria (WRF Weather Forecast, 2014), with a horizontal resolution of 9.7 km. The initial and boundary conditions are from the Global Forecast System (GFS) model, with a horizontal resolution of  $0.5^\circ$ .

The WRF model description can be found in Skamarock et al. (2008). The model can be run with a spatial resolution between 1 and 10 km. Numerous specific models, such as the Hurricane Weather Research and Forecasting (HWRF) have been created upon WRF. From the first release in 1990 until now the model has evolved (Michalakes et al., 1998) and additional packages have been developed for interactive nesting, upgraded physics, three-dimensional data assimilation and simplified parallelisation (Michalakes et al., 2004). WRF has a large worldwide community, with over 20 000 users in over 130 countries.

The near-future use of the WRF model will be (1) to replace the synop observations for the derivation of GNSS–IWV, (2) to compare the model water vapour field with GNSS products, and (3) to assimilate GNSS–IWV.

### 3 GNSS meteorology

The concept of GNSS meteorology was suggested by Bevis et al. (1992). The propagation of the GNSS signal through the atmosphere is affected by the atmospheric gases. The magnitude of the atmospheric effects depends on several factors: on the composition of the atmosphere; on the elevation of the receiver (thus on the thickness of the atmosphere); on the elevation angle of the satellite; and finally on the amount of water vapour, which depends mainly on the current atmospheric conditions.

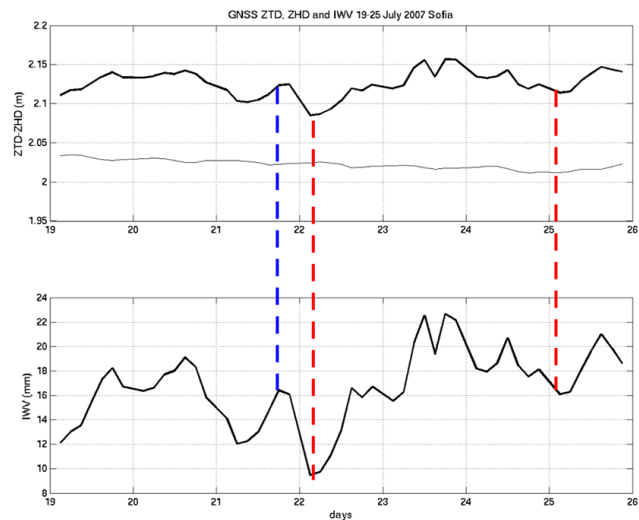
There are two contributing factors for the signal path delay in the lower atmosphere–troposphere: they are the hydrostatic delay and the wet delay. The hydrostatic delay is caused by all the gases in the atmosphere, except the water vapour. The hydrostatic delay in the zenith direction is called zenith hydrostatic delay (ZHD), and as seen from Fig. 3a, is relatively stable on a day timescale. It can be derived by using its dependency on the local atmospheric pressure (Bevis et al., 1992; Emaradson et al., 1998):

$$\text{ZHD} = (2.2768 \pm 0.0024) \frac{p_s}{f(h, \theta)}, \quad (2)$$

$$f(h, \theta) = 1 - 0.00266 \cos(2\theta) - 0.00028h, \quad (3)$$

where  $p_s$  is local surface pressure and  $f(h, \theta)$  is a factor dependent on height  $h$  and the latitude variation of gravitational acceleration  $\theta$ .

The second contributing factor is the zenith wet delay (ZWD). It is caused by the water vapour in the atmosphere. The ZWD has a large temporal variation in an hour timescale. This is the reason, why the GNSS-derived integrated water vapour (IWV) is so valuable with its high temporal resolution (Fig. 3b). The ZWD contributes less than 10 % of the



**Figure 3.** GNSS meteorology: ZTD, ZHD and IWV at station SOFI during the 2007 heatwave from 19 to 25 July. (a) Zenith total delay (ZTD), zenith hydrostatic delay (ZHD) in m, and (b) integrated water vapour (IWV) in mm. The diurnal cycle of IWV with an increase in the afternoon hours and a minimum in early morning is seen well.

zenith total delay (ZTD). ZWD and IWV can be calculated by using

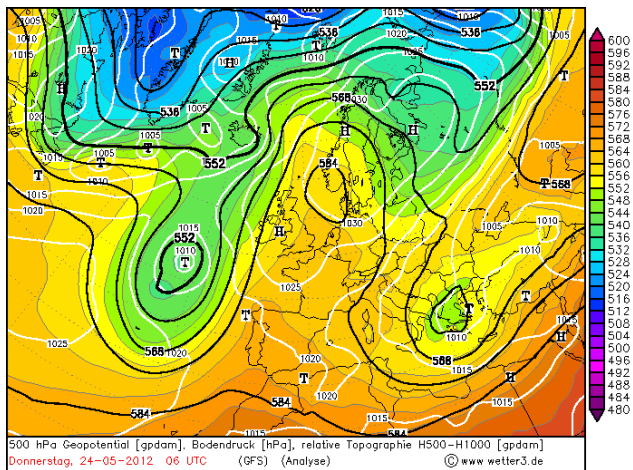
$$\text{ZWD} = \text{ZTD} - \text{ZHD}, \quad (4)$$

$$\text{IWV} = \frac{10^6}{(k_3/T_m + k_2)R_v} \text{ZWD}, \quad (5)$$

where  $k_2$ ,  $k_3$  and  $R_v$  are constant and  $T_m$  is the weighted mean atmospheric temperature. Uncertainties in the IWV derived from ground-based GNSS are caused by errors in (1) surface pressure (1 hPa corresponds to 0.33–0.37 mm in IWV) and (2) ZTD (1 mm corresponds to 0.15–0.16 mm in IWV). The accuracy of IWV is estimated to be in the range  $\pm 1$  mm.

At present, surface pressure and temperature from the synop stations are used to derive the IWV from GNSS (see Sect. 2.2.3). For example, for 11 stations from the ZenitGEO network appropriate surface synop stations from NIMH network (Sect. 2.2.3) are allocated. It is to be noted that the synop observations have a temporal resolution of 3 h, while GNSS tropospheric products have a temporal resolution of 5 min; thus, the derived ZHD, ZWD and IWV are degraded to being three-hourly. In addition, the altitude corrections for temperature and pressure are required for most of the synop stations. For the remaining 19 stations of the ZenitGEO network no appropriate surface stations are available. In near future use of surface pressure and temperature from the WRF model (Sect. 2.2.4) are envisaged.

It is to be noted that IWV decreases with the altitude with the bottom 5.5 km containing 97 % of total atmospheric water vapour content. Measuring an integrated quantity at a



**Figure 4.** Geopotential height at 500 hPa (black lines), surface pressure (white lines) and thickness (colour map) at 06:00 UTC on 24 May 2012 from the GFS analysis.

different altitude in the lower atmosphere will thus depend on location and, importantly, on the elevation above sea level. To represent IWV spatial distribution, two-dimensional maps are suggested by Morland and Maetzler (2007). They propose an altitude correction for 500 m a.s.l.:

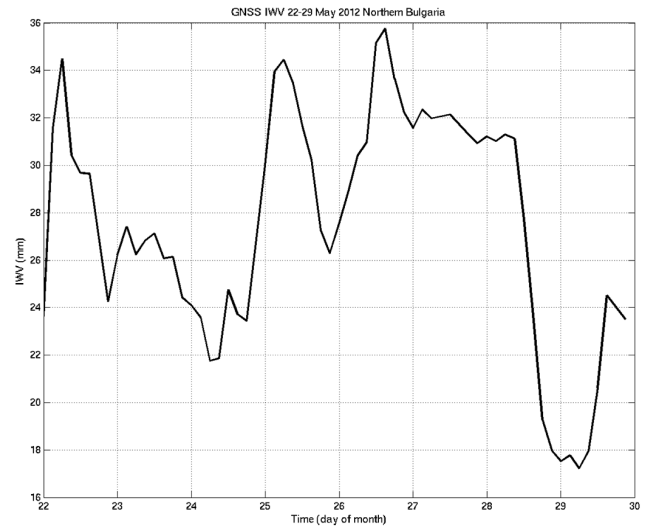
$$\text{IWV}(0.5) = a \times \text{IWV}(h) \times \exp\left[\frac{h - 0.5}{H}\right], \quad (6)$$

where  $\text{IWV}(0.5)$  is IWV at altitude 500 m,  $\text{IWV}(h)$  is the estimated IWV at altitude  $h$ ,  $a$  are empirically derived coefficients, and  $H$  is the scale height. This correction is applied to 11 stations from the ZenitGEO network at altitude between 36 and 542 m. The produced two-dimensional maps are used for convection case studies in Sect. 4.1.

## 4 Case studies

### 4.1 Short-term variation of IWV: intense precipitation events in 2012

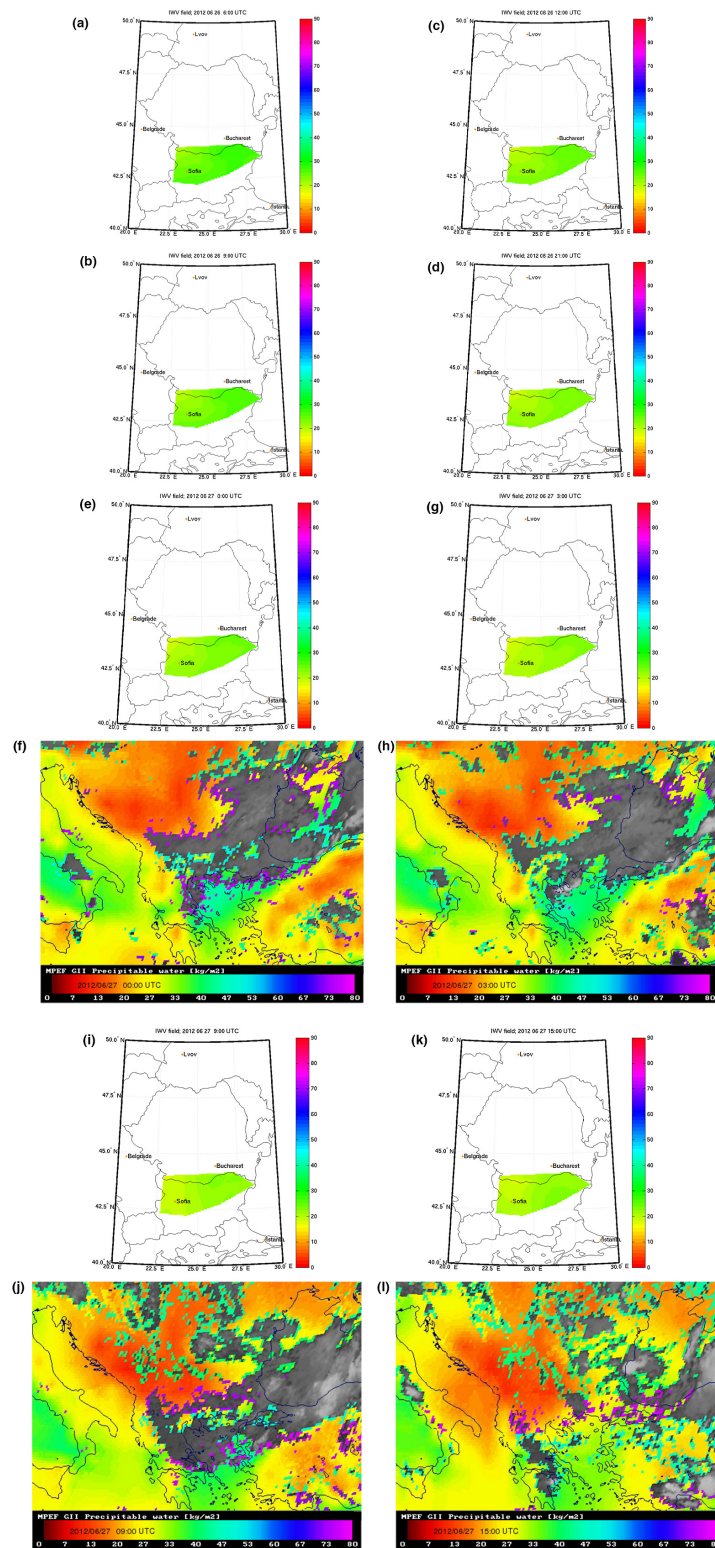
In this work, two events with intense precipitation in Bulgaria, one associated with frontal passage and one with the development of local convection, are presented. The first event is on 25 May 2012, when a cold front passes over Bulgaria. The period 22–29 May is characterised by a low-pressure field over the Mediterranean area. Prolonged precipitation and thunderstorms are a result of the air mass instability or frontal passage. From 24 to 25 May, a trough can be seen on the 500 hPa map, and a cyclonic field near the surface (Fig. 4). On 24 May, a cold front connected with a cyclonic centre over the Sea of Azov passes over Bulgaria. On the next day, 25 May 2012, another cold front passes Bulgaria from north to south. The intrusion of a cold air mass can be seen



**Figure 5.** GNSS-IWV for 22–29 May 2012 at Lovech, Bulgaria.

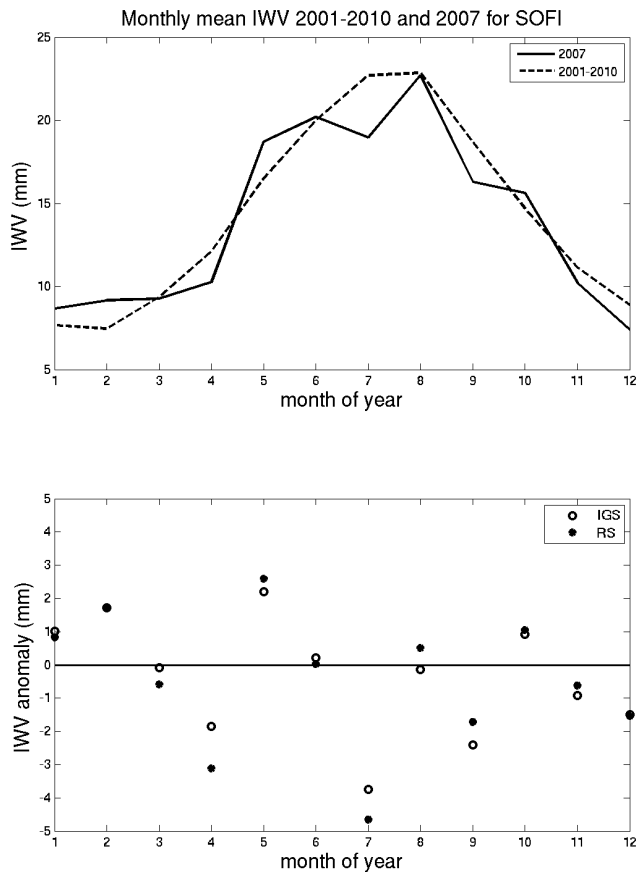
in Fig. 5 at 15:00 UTC over Northern Bulgaria. The water vapour values decrease by 10, to 25 mm (Fig. 4).

For the period 24–27 June, the weather is dynamic, changeable with unstable air mass, cumulonimbus cloud development and precipitation with different ranges and intensities. On 26 June, a cold front passes over Bulgaria, and the temperature at 850 hPa drops from 18 °C at 00:00 UTC to around 10 °C at 06:00 UTC on 27 June 2012. There is a considerable temperature gradient both at altitude and surface, and the maximum temperature decreased by 5–6 °C. At 12:00 UTC on 26 June, dry air mass was advected from north to west, spreading along the Balkan mountain range (Fig. 6). This pathway of cold and dry air is usually associated with intense precipitation of both rain and snow in the spring and autumn seasons. The consecutive two-dimensional IWV maps from 26 and 27 June capture well the advancement of the dry air mass. In less than 24 h the IWV in Northern Bulgaria decreased by half from above 35 mm at 06:00 UTC on 26 June to 15–20 mm at 03:00 UTC on 27 June. After the cold front passage on 26 June the air mass remains unstable with high relative humidity at level 700 hPa and warming at high altitudes (200 hPa) contributing for the convection development. On 27 June, the advected dry air catches the receding humid air mass, resulting in isolated convective cell development, with thunderstorms and intense precipitation. Intense rainfall of 74 L m<sup>-2</sup> for six hours is recorded between 09:00 and 15:00 UTC on 27 June in the Black Sea region of Kaliakra. The strong north–south gradient of IWV over the Balkan Peninsula is confirmed by the Meteosat-derived maps (Schroedter-Homscheidt et al., 2008). From the Meteosat maps, an isolated convective cell is clearly seen at 15:00 UTC in the infrared cloud cover image. The presented in this paper case studies demonstrate the synergy between GNSS and Meteosat water vapour maps. Future work will be a detailed



**Figure 6.** Two-dimensional IWV maps from GNSS on 26 June 2012 at (a) 06:00 UTC, (b) 09:00 UTC, (c) 12:00 UTC and (d) 21:00 UTC. Two-dimensional maps on 27 June 2012 at 00:00 UTC: (e) GNSS and (f) Meteosat, 06:00 UTC; (g) GNSS and (h) Meteosat, 09:00 UTC; (i) GNSS and (j) Meteosat, 15:00 UTC; (k) GNSS and (l) Meteosat. The colour maps of Meteosat and GNSS–IWV are reversed.



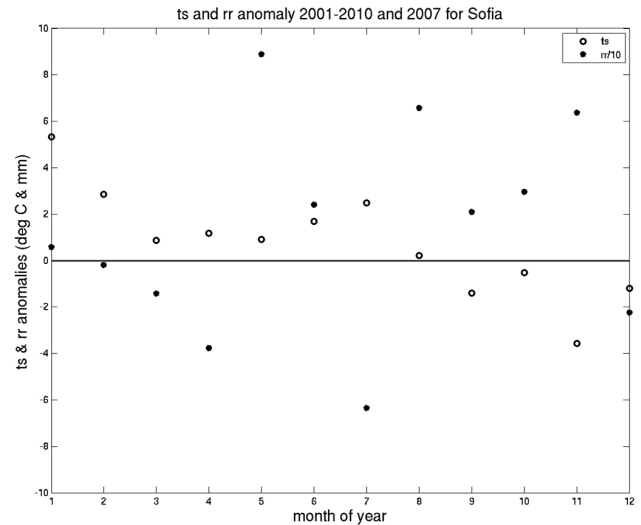


**Figure 7.** Top figure: monthly mean IWV for SOFI, Bulgaria (thick line 2007, dashed line 2001–2010). Bottom figure: monthly anomaly (difference 2007 mean and 2001–2010 mean) from GNSS (open circles) and radiosonde (filled circles).

analysis of 20 further convective situations for 2012. This work is a contribution to working group two of COST Action ES1206 “Advanced Global Navigation Satellite Systems tropospheric products for monitoring severe weather events and climate (GNSS4SWEC)”.

#### 4.2 Long-term variation of IWV: the 2007 heatwave

Heatwaves have become a common summer feature in south-eastern Europe (Matzarakis et al., 2007). The July 2007 heatwave has the largest geographical extension, reaching Bulgaria. The atmospheric circulation leading to the heatwave is characterised by a northerly displacement of the subtropical jet stream (flow at 200 hPa) that allowed subtropical African air to reach southeastern Europe as far as 50° N. The GNSS–IWV from the IGS repro2 and RS–IWV are used to study the 2007 heatwave. The annual and seasonal mean GNSS–IWV for the period 2001–2010 is compared to the 2007 period, and is presented in Table 3.



**Figure 8.** Monthly anomaly of temperature (open circles) and precipitation (filled circles) for Sofia, Bulgaria.

As seen in Table 3, the annual GNSS–IWV in 2007 is 14.0 mm, and is similar to the 2001–2010 mean. The seasonal values show that in 2007 the GNSS–IWV is larger for winter (+5 %) and smaller for summer and autumn (–5 and –6 % respectively). For comparison, the IWV from the radiosonde (RS–IWV) station in Sofia is also presented in Table 3. The RS–IWV annual, winter, summer and autumn seasonal means in 2007 have the same tendency as GNSS–IWV. It is to be noted that there is a difference between the sampling rates of GNSS and RS, the first being 3 h, while the second is once a day at 12:00 UTC. In addition, the radiosonde station is in Sofia (marked by the red pointer in Fig. 2) and at an altitude of 590 m, while the GNSS station is on the Plana mountain (marked by the blue pointer in Fig. 2) and at an altitude of 1120 m.

In addition, the monthly IWV anomalies are studied. In Fig. 7 are plotted GNSS–IWV and RS–IWV anomalies. When GNSS–IWV in 2007 (solid line in Fig. 7a) is compared to 2001–2010 (dashed line in Fig. 7a), the following features stand out: (i) an IWV decrease in April, (ii) an IWV increase in May, and (iii) a sharp IWV decrease in July. Clearly seen from Fig. 7b is that the largest negative IWV anomaly is in July, about –4 mm from GNSS–IWV and –5 mm from RS–IWV. There is very good correlation of the anomaly from the two techniques despite the different sampling rates and locations. The difference between the GNSS–IWV and RS–IWV anomaly is under 0.5 mm in 7 months, between 0.5 and 1 mm in 2 months, and about 1 mm in 3 months.

The 2007 winter was 2.4 °C warmer than the 2001–2010 (column 5 in Table 4). The 2007 spring and summer were with 1 and 1.4 °C warmer. In particular, July 2007 was +3.7 °C warmer (Fig. 8) and with less IWV than the 2001–2010 with –16 and –19 % respectively for the GNSS–IWV and RS–IWV. It is to be noted that the annual precipitation

**Table 3.** Columns 1–3 are the station name and annual mean for 2001–2010 and 2007 respectively; columns 4–5: winter DJF (December, January and February) mean for 2001–2010 and 2007; columns 6–7: spring MAM (March, April and May) mean for 2001–2010 and 2007; columns 8–9: summer JJA (June, July and August) mean for 2001–2010 and 2007; columns 10–11: autumn SON (September, October and November) mean for 2001–2010 and 2007. The 2007 departure from the 2001–2010 mean is given in % in the brackets.

Station	2001–2010	2007	2001–2010	2007	2001–2010	2007	2001–2010	2007	2001–2010	2007
IWV–IGS repro2	mm	mm	DJF	DJF	MAM	MAM	JJA	JJA	SON	SON
SOFI	14.3	14.0	8.0	8.4	12.7	12.7	21.8	20.6	14.8	14.0
Change		–2 %		+5 %		0 %		–5 %		–6 %
IWV–RS	mm	mm	DJF	DJF	MAM	MAM	JJA	JJA	SON	SON
Sofia	15.5	15.1	8.8	9.1	13.6	13.2	23.9	22.5	16.0	15.5
Change		–3 %		+3 %		–3 %		–6 %		–3 %

**Table 4.** Annual and seasonal means of IWV, temperature and precipitation for station Sofia, Bulgaria for winter (DJF), spring (MAM), summer (JJA) and autumn (SON) for 2001–2010 and 2007.

Met station	2001–2010	2007	2001–2010	2007	2001–2010	2007	2001–2010	2007	2001–2010	2007
Temperature	[°C]	[°C]	DJF	DJF	MAM	MAM	JJA	JJA	SON	SON
Sofia	8.4	8.8	0.2	2.6	10.7	11.7	20.4	21.8	11.3	9.5
Change		+5 %		+1200 %		+9 %		+7 %		–16 %
Precipitation	mm month <sup>–1</sup>	mm month <sup>–1</sup>	DJF	DJF	MAM	MAM	JJA	JJA	SON	SON
Sofia	55	69	40	34	55	66	73	86	50	91
Change		+25 %		–15 %		+20 %		+17 %		+82 %

amount in 2007 was 25 % higher than the 2001–2010 amount. However, the winter was 15 % drier (column 5 in Table 4), and from spring, only the month of May has a positive precipitation anomaly over 80 mm, as seen in Fig. 8. In the summer of 2007, the month of July was very dry (about 60 mm less than the 2001–2010 mean).

## 5 Conclusions

The Sofia University Atmospheric Data Archive (SUADA) is a regional database for Bulgaria and southeastern Europe. GNSS tropospheric products (over 12 000 000 individual observations) and derivatives (over 55 000) from five GNSS processing strategies and 37 stations for the period 1997–2013 are archived in SUADA. The temporal resolution of the GNSS data is from 5 min to 6 h. In addition, over 6000 individual IWV from the radiosonde station in Sofia for the period 1999–2012 are archived in SUADA.

The application of SUADA data is shown in case studies for intense precipitation events in 2012 and during the heatwave in 2007. At 12:00 UTC on 26 June dry air mass was advected from the northwest, spreading along the Balkan mountain range. Near the Black Sea coast, the advected dry air catches the receding humid air mass, resulting in the development of local convection and intense rainfall with 74 L m<sup>–2</sup>. The two-dimensional IWV maps capture well the advancement of the dry air mass. In less than 24 h, the IWV in Northern Bulgaria decreased by half, from above 35 mm

at 06:00 UTC on 26 June to 15–20 mm at 03:00 UTC on 27 June. The strong north–south gradient of IWV over the Balkan Peninsula is confirmed by the Meteosat-derived product. The second application of the SUADA data is a study of the IWV anomaly during the 2007 heatwave in Bulgaria. Despite the difference in the location and sampling rate, the two data sets give a negative IWV anomaly in July 2007 of about –4 mm from GNSS–IWV, and –5 mm from RS–IWV. July 2007 has less IWV compared to 2001–2010, with –16 and –19 % respectively for GNSS–IWV and RS–IWV.

This work is a first step in the application of GNSS tropospheric products for atmospheric sounding in Bulgaria and southeastern Europe. The work will continue with (1) an improvement in the temporal and spatial resolution of the GNSS–IWV, notably the ZenitGEO, (2) validation of NWP model WRF with the GNSS data set for Bulgaria, (3) analysis of additional convective cases in 2012, and (4) an analysis of the diurnal cycle of GNSS–IWV during 18 fog events at Sofia Airport in the period 2011–2012. The work is a contribution to working groups two and three of COST Action ES1206 “Advanced Global Navigation Satellite Systems tropospheric products for monitoring severe weather events and climate (GNSS4SWEC)” 2013–2017.

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