



Aerosol optical depth over the Arctic: a comparison of ECHAM-HAM and TM5 with ground-based, satellite and reanalysis data

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Abstract. We compare ground-based measurements of aerosol optical depth and Ångström parameter at six Arctic stations in the period 2001–2006 with the results from two global aerosol dynamics and transport models, ECHAM-HAM and TM5. Satellite measurements from MODIS and the MACC reanalysis product are used to examine the spatial distribution and the seasonality of these parameters and to compare them with model results. We find that both models provide a good reproduction of the Ångström parameter but significantly underestimate the observed AOD values. We also explore the effects of changes in emissions, model resolution and the parametrization of wet scavenging.

1 Introduction

The Arctic is extremely vulnerable to past and future climate change, through complex interactions which can lead to severe regional impacts on the local hydrology, cryosphere and ecosystems, and to feedbacks on the global climate system (IPCC AR4, 2007). Significant changes have already occurred during recent decades, involving loss of sea-ice and snow-cover (e.g., Serreze et al., 2007) and affecting other important components of the environment (e.g., Post et al., 2009; Richter-Menge and Overland, 2011, and references therein). Anthropogenic and natural atmospheric aerosols

play a crucial role in these processes. While concentrations in the Arctic are low on average, they reach a maximum in spring, forming the so-called Arctic haze, composed mainly of sulfates, particulate organic matter, black carbon and dust, and believed to be mainly anthropogenic (Shaw, 1995; Quinn et al., 2007). The direct climatic effect of aerosols on the radiative balance of the atmosphere, through absorption and scattering of incoming short-wave radiation, is particularly significant in the polar areas, because of the high surface albedo due to snow and ice, leading to atmospheric warming and reduction of the solar radiation reaching the surface. Indirect effects linked with cloud droplet formation are also expected to be important in the Arctic owing to the low aerosol number concentrations in these areas (Law and Stohl, 2007). The deposition of carbonaceous particles on snow reduces the surface albedo, leading to changes in snow temperature and structure, modifications in seasonal snow melting and aging, sea ice thinning and glacier reduction (Flanner et al., 2009). On a global scale these processes have been recognized to contribute significantly to the climatic forcing of these aerosols (Holland and Bitz, 2003; Hansen and Nazarenko, 2004; Flanner et al., 2007; Ramanathan and Carmichael, 2008).

The source areas for nearly all air pollutants in the Arctic are believed to be in mid latitudes (Law and Stohl, 2007), but the presence of transport barriers and complex removal

processes leads to a difficult identification of the transport pathways, which has advanced only recently (see, e.g., Koch and Hansen, 2005; Stohl, 2006; Shindell et al., 2008; Hirdman et al., 2010a,b; Huang et al., 2010; Bourgeois and Bey, 2011).

Global atmospheric and aerosol models are important tools for studying climatic feedbacks and for estimating their impact on the climate system. They also provide boundary conditions for higher resolution regional models. In both cases, a reasonable representation of the concentration and optical properties of aerosols over the Arctic is required. However, the verification of these models has been performed mainly on the global scale, focussing at low and mid latitudes, and their skill in the Arctic is still largely unexplored. Aerosol concentration measurements in the Arctic are sparse, and the main source of data for model verification is provided by measurements of atmospheric optical properties collected by a network of a small number of measuring stations and by satellite observations.

In this work we compare the aerosol optical properties as modeled by two state-of-the-art aerosol models with ground-based station measurements in the Arctic. We consider the atmospheric aerosol transport and chemistry model TM5 (Krol et al., 2005) and the global climate model ECHAM5 (Roeckner et al., 2003), which has been shown to perform well in Arctic areas (Walsh et al., 2008), coupled with the aerosol transport and dynamics module HAM (Stier et al., 2005). A comprehensive comparison in terms of aerosol burdens and optical properties on a global scale with other models has been performed in the framework of the aerosol model inter-comparison initiative AeroCom (Textor et al., 2006; Kinne et al., 2006). A comparison focused on the Arctic has been reported in Shindell et al. (2008). Here we compare modeled aerosol optical depths and Ångström coefficients with ground-based measurements (Tomasi et al., 2007), analyzing the role of different aerosol components. In order to compare spatial distributions of optical depth, we also consider remote sensing data from the MODIS satellites and reconstructed fields from the recent reanalysis project MACC. We explore the sensitivity of modeled Arctic aerosol optical depths to different emission datasets and model resolutions, considering also a simple change in the parametrization of wet scavenging for HAM which has been recently suggested (Bourgeois and Bey, 2011).

In the following, Sect. 2 provides information on the available station, satellite and reanalysis data and describes the ECHAM-HAM and TM5 models. In Sect. 3 we report and discuss the comparison between observed and modeled aerosol optical properties. Concluding remarks are provided in Sect. 4.

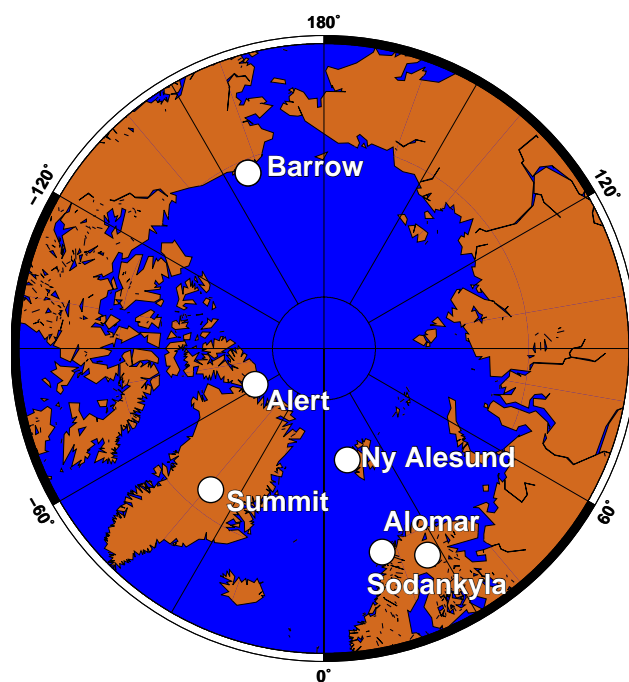


Fig. 1. Location of the six measurements stations considered in this work.

2 Models and data

2.1 Ground-based measurements

We focus on long timeseries of sun photometer measurements of the daily mean values of AOD (500 nm) collected by groups participating in the POLAR-AOD programme (Aerosols Optical Depth in Polar regions; see Tomasi et al. (2007) and <http://polaroad.isti.cnr.it:8080/Polar/index.jsp>) at six Arctic sites, in the period 2001–2006 (Fig. 1). The measurement sites and periods which are available are: Barrow (Alaska, 71°19' N, 156°36' W) from 2002/2 to 2005/10, Alert (Canada, 82°28' N, 62°30' W) from 2004/8 to 2006/5, Summit (Greenland, 72°20' N, 38°45' W) from 2001/3 to 2005/7, Ny Ålesund (Norway, 78°58' N, 11°54' E) from 1994/3 to 2006/4 (used from 2001 to 2006 in this work), ALOMAR (Norway, 69°17' N, 16°00' E) from 2002/5 to 2006/4 and Sodankylä (Finland, 67°22' N, 26°38' E) from 2004/5 to 2006/4. The average errors on daily AOD measurements at these sites due to instrumental characteristics, calibration errors, and atmospheric corrections for Rayleigh scattering and gaseous absorption, range from 0.018 (Summit) to 0.035 (Barrow). The Ångström turbidity parameter α was derived by fitting a power law to spectra of AOD measured at different wavelengths. A complete description and analysis of these data is provided by Tomasi et al. (2007). In order to allow for a comparison with modeled, reanalysis and satellite AODs (which are all available at 550 nm), we interpolate measured AODs to 550 nm using

the available Ångström coefficients (the resulting values are lower by about 10 % on average).

2.2 Satellite measurements and reanalysis

Satellite observations of AOD and Ångström parameter are provided by the Moderate Resolution Imaging Spectroradiometer (MODIS) aboard the Terra and Aqua satellites. Specifically the Aerosol Cloud Water Vapor Ozone Daily L3 Global 1Deg CMG collection products were used. In this work we use data starting in 2001 (Terra) and in July 2002 (Aqua). Ångström parameters are based on AODs at 470 nm and 660 nm over the land and 550 nm and 865 nm over the ocean.

We also consider an aerosol reanalysis product provided by the MACC (Monitoring Atmospheric Composition and Climate) project (Benedetti et al., 2009), which uses the ECMWF IFS cycle 36R1 model with a prognostic aerosol scheme at resolution T255L60, assimilating MODIS AOD observations. Aerosol reanalysis data are available starting from 1 January 2003. To our knowledge a complete validation of the MACC reanalysis product in the Arctic is still missing. In this work we treat the MACC reanalysis as an additional source of spatially extended AOD observations and we use it as an ‘interpolator’ to provide reference data also in months when satellite observations at high latitudes are scarce.

2.3 ECHAM5-HAM and emissions

The ECHAM5-HAM model couples the global climate model ECHAM5 (Roeckner et al., 2003) with the HAM module (Hamburg Aerosol Module; Stier et al. (2005); Lohmann and Hoose (2009)), which models the dynamics, the microphysics and the transport of the main atmospheric aerosols and their radiative feedbacks. In particular HAM contains the microphysical core M7 (Vignati et al., 2004), based on the representation of particle distributions as the superposition of log-normal modes peaked at different particle size classes, and reproduces the main aerosol emission, sedimentation and wet and dry scavenging processes. The aerosol compounds included are sulfates, black carbon, organic matter, sea salt and mineral dust. The emissions of dust, sea salt and oceanic dimethyl sulfide (DMS) are computed on-line, while other natural and anthropogenic emissions are prescribed. Secondary Organic Aerosols (SOA) emissions are prescribed monthly and added to Particulate Organic Matter emissions, as described in Stier et al. (2005) (19 Tg C yr^{-1} globally). The main optical and microphysical properties of the aerosols, such as optical depth, and their number and mass concentrations are simulated by the model. In this study we set the spatial resolution to T42 in spectral space (corresponding to a resolution of about $2.8^\circ \times 2.8^\circ$ on a Gaussian grid). The vertical resolution is set to 19 vertical levels from the surface up to 10 hPa.

We integrate the ECHAM-HAM model using a “nudging” technique to force the model to stay close to the dynamical wind fields provided by the ECMWF ERA-Interim database (Dee et al., 2011) in the period 2000–2006. The nudging fields were prepared using the IN-TERA package (Ingo Kirchner, <http://wekuw.met.fu-berlin.de/~IngoKirchner/nudging/nudging/>). The initial year is used for spinup, and we consider model results in the period 2001–2006. Since transport processes play an important role in determining the concentration of Arctic aerosols (Shindell et al., 2008), nudging was used in order to allow the model runs to reproduce as close as possible the main tropospheric winds in the period of interest and to allow a direct comparison with observed data over a limited range of years.

In this work we explore the use of different aerosol emission databases as boundary conditions for ECHAM-HAM, which are summarized in Table 1 together with the labels used to identify them. Table 2 details the total anthropogenic and wildfire emissions for these datasets, both globally and for the northern hemisphere excluding the equatorial area (north of 10° N). In particular we use the AeroCom-I (Den-tenner et al., 2006) and the ACCMIP (Atmospheric Chemistry and Climate Model Intercomparison Project, Lamarque et al., 2010) inventories, both for the year 2000. Anthropogenic fossil-fuel and bio-fuel emissions for sulfur, black and organic matter are annual data. Wildfire burning emissions are represented as monthly climatologies. Since fires in Siberia, Canada and Alaska have been found to provide a significant contribution to Arctic pollution (Law and Stohl, 2007; Stohl, 2006; Stocks et al., 1998; Koch and Hansen, 2005), with important interannual variability, the choice of biomass burning emissions datasets may play an important role. Variability in these emissions, interacting with interannual changes in circulation, may lead to significant changes in Arctic aerosol concentrations. While ACCMIP includes data from the Global Fire Emission Database 2 (GFED2), we also explore ACCMIP emissions using the more recent GFED3 monthly biomass burning emission database (van der Werf et al., 2010) and we allow these emissions (sulfur, black carbon and organic matter) to vary annually instead of using a climatological mean.

Koch and Hansen (2005) showed that, while the largest contribution in terms of sulfate aerosols to Arctic AOD comes from Russia, South-East Asia contributes with a significant fraction. Since there is a significant difference in terms of sulfate emissions between the REAS and ACCMIP emission datasets, we explore the possible impact of a change in anthropogenic SO_2 emissions in the South-East Asia region. To this end, we consider the latter emissions (ACCMIP+GFED3), rescaling ACCMIP 2000 anthropogenic sulfate emissions in the region $95\text{--}135^\circ \text{ E}$ – $10\text{--}50^\circ \text{ N}$ to the REAS database (Ohara et al., 2007) averaged over the years 2001–2006 (i.e. we multiply the ACCMIP anthropogenic SO_2 emissions in this region by a factor 1.7). In all configurations, injection heights of emissions are prescribed

Table 1. Emission datasets used for ECHAM5-HAM runs.

Label	Emission inventories	Description	Reference
EAERO	AeroCom I	Anthropogenic: annual climatology (year 2000). Wildfire emissions have a monthly climatology.	Dentener et al. (2006) Stier et al. (2005)
EIPCC	ACCMIP	Anthropogenic: annual climatology; Wildfires: monthly clim.	Lamarque et al. (2010)
EGFED	ACCMIP+GFED3	ACCMIP + wildfires (monthly) with interannual variations.	van der Werf et al. (2010)
EREAS	ACCMIP+GFED3+REAS	Sulfur emissions in South-East Asia scaled to REAS average.	Ohara et al. (2007)

Table 2. Anthropogenic emissions for each emission dataset (in [Tg yr⁻¹]).

		Global emissions				Emissions north of 10° N			
		EAERO	EIPCC	EGFED	EREAS	EAERO	EIPCC	EGFED	EREAS
BC	Fossil fuel	3.03	3.06	3.06	3.06	2.58	2.53	2.53	2.53
	Biofuel ^a	1.63	2.09	2.09	2.09	1.14	1.57	1.57	1.57
	Wildfire	3.04	2.61	2.07	2.07	0.54	0.61	0.45	0.45
OC ^d	Fossil fuel	2.44	4.25	4.25	4.25	1.96	3.17	3.17	3.17
	Biofuel ^a	6.47	8.45	8.45	8.45	4.53	6.28	6.28	6.28
	Wildfire	24.7	23.2	18.1	18.1	5.38	6.99	5.32	5.32
SU	High ^b	95.2	91.0	91.0	104.4	85.5	77.6	77.6	91.1
	Low ^c	13.0	12.7	12.8	15.7	11.5	10.9	10.9	13.8
	Wildfire	4.09	3.83	2.29	2.29	0.74	1.07	0.62	0.62

^a Biofuel = biomass burning + agricultural waste + fuelwood; ^b SU high = sulfur from industry + powerplants + shipping; ^c SU low = sulfur from domestic + roads + off road; ^d Particulate organic matter = 1.4 × Organic Carbon.

in the model as described in Stier et al. (2005) and Dentener et al. (2006). We use the Balkanski et al. (2004) scheme for on-line dust emissions and Schulz et al. (2004) for on-line sea-salt emissions. Other emissions and parametrizations are as described in Stier et al. (2005) and in Lohmann and Hoose (2009).

In the following we also explore a simple change in the model wet scavenging parametrization introduced recently by Bourgeois and Bey (2011) to better reproduce the observed optical properties and concentrations of aerosols in the Arctic region. HAM parametrizes wet scavenging using the precipitation formation rate of the ECHAM5 cloud scheme and computing the fraction of tracer that is embedded in cloud water using simple size-dependent and cloud-type dependent scavenging coefficients, based on measurements from Henning et al. (2004). Bourgeois and Bey (2011) explored reducing these coefficients, based on a re-evaluation of the results in Henning et al. (2004) and following papers, leading to an increase in BC and sulfate lifetimes and to increased burdens in the Arctic, while the global and annual scavenged masses remained similar. We apply this modification using the same parameters as described in their paper.

2.4 TM5

TM5 is a global three-dimensional atmospheric chemistry and transport model (Krol et al., 2005). Aerosol microphysical processes are modeled using the aerosol dynamics module M7 (Vignati et al., 2004). It represents sulphate, black carbon, organic carbon, sea salt and mineral dust in seven internally mixed soluble or insoluble log-normal size modes. Gas-particle partitioning of ammonium nitrate is calculated using the EQSAM thermodynamic equilibrium model, as described in Aan den Brugh et al. (2011). The gas-phase chemistry scheme is based on the Carbon Bond Mechanism 4 (CMB4) and is given in Huijnen et al. (2010). In our setting, the horizontal resolution is 3 × 2 degrees and the vertical grid comprises 34 hybrid σ -pressure levels. Atmospheric dynamical fields are provided by ECMWF ERA-Interim reanalysis data (Dee et al., 2011).

Anthropogenic and biomass burning emissions are taken from the CMIP5/ACCMIP datasets. The year-2000 values from the historical dataset described in Lamarque et al. (2010) are combined with scenario estimates for the year 2005 and 2010 from the representative concentration pathway RCP4.5 (van Vuuren et al., 2011). Linear interpolation is applied for the intermediate years.

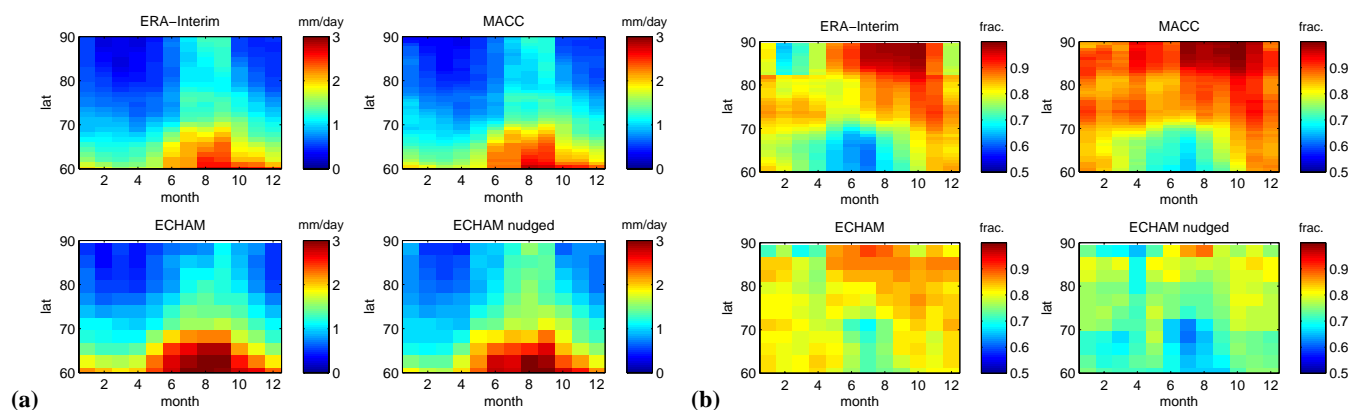


Fig. 2. Zonal averages of precipitation and total cloud cover as a function of month and latitude for models and reanalyses used in this study. The total precipitation fields are monthly averaged in the period 2003–2006 (common to all datasets) for ERA-Interim, MACC, ECHAM (the EAERO run) and for ECHAM nudged to ERA-Interim (EAERO). (a) Total precipitation. (b) Total cloud cover fraction.

Emissions of sea salt, oceanic DMS and nitrogen oxides (NO_x) from lightning are calculated online, while other natural emissions are prescribed. Sea-salt emissions are parameterized as described in Vignati et al. (2010). The schemes applied for DMS and lightning NO_x are the same as in Huijnen et al. (2010). The emissions of mineral dust are prescribed using the AeroCom-I dataset for the year 2000 (Dentener et al., 2006). Biogenic emissions of isoprene and other natural emissions are prescribed based on yearly and monthly datasets compiled by the MACC project (total SOA emission fields correspond to 19 Tg yr⁻¹). The implementation of the emission heights for the different sources and anthropogenic sectors has been revised compared to the description given in Huijnen et al. (2010).

3 Results

3.1 Model precipitation

In the Arctic wet removal processes are the main mechanisms for aerosol removal (gravitational sedimentation plays a minor role because the aerosol is dominated by small particles). For these reasons the ability of a global climate model in reproducing reasonably the atmospheric branch of the water cycle is crucial, in particular cloud and precipitation fields at high latitudes and their seasonality. We compare Arctic precipitation climatologies in ERA-Interim, MACC and in the ECHAM model, reporting in Fig. 2a zonally averaged total precipitation fields, as a function of the month, averaged over the period 2003–2006, common to all datasets. Both a nudged and free simulation of ECHAM are shown. While average precipitation is comparable in winter months, ECHAM shows slightly higher precipitation in summer, compared both to MACC and ERA-Interim, mainly south of 70° N, particularly for the non-nudged simulation. Investigation of spatial maps (not shown) reveals that this excess summer pre-

cipitation occurs mainly over the areas from eastern Siberia to Alaska. This difference may be particularly relevant for removal of OC, BC and sulfur aerosols emitted by summer fires in these areas. Figure 2b also reports a comparison of total cloud covers (as a fraction) in the different models. While MACC shows the highest cloud fraction in all months at high latitudes, and while cloud cover in ECHAM, in particular in the nudged run, is lower compared to ERA-Interim and MACC, the differences remain of the order of 0.1.

3.2 Timeseries and monthly aerosol climatologies

Figure 3a and b compare daily values of AOD observed at the measurement sites with those modeled at the nearest gridpoint by ECHAM-HAM (using the EIPCC emissions) and TM5. As the figures clearly show, AOD measurements are available only for limited periods and with several gaps, different for each site. For the Ny Ålesund, Sodankyla and ALOMAR sites, both models are capable of reproducing extreme values of AOD comparable to observations. At Barrow, Alert and Summit a significant underestimation of observed values is apparent for both models. In evaluating these plots it is important to recall the extreme intermittency observed in Arctic haze, fire smoke and Asian dust transport episodes, which often take place over few days. While the model simulations are based on realistic wind fields (with TM5 using ERA-Interim fields and ECHAM nudged to the same), and may be able to reproduce correctly transport episodes from low latitudes, aerosol emissions used here are based on monthly climatologies, so that the daily variability of local sources and the interaction between individual fire episodes and specific transport patterns may not be reproduced.

In Fig. 4a we compare monthly means of observed AOD with climatologies from the ECHAM-HAM and TM5 models in the period 2001–2006. To indicate the data availability and interannual variability we report the individual monthly means of the station data for every year, instead of their

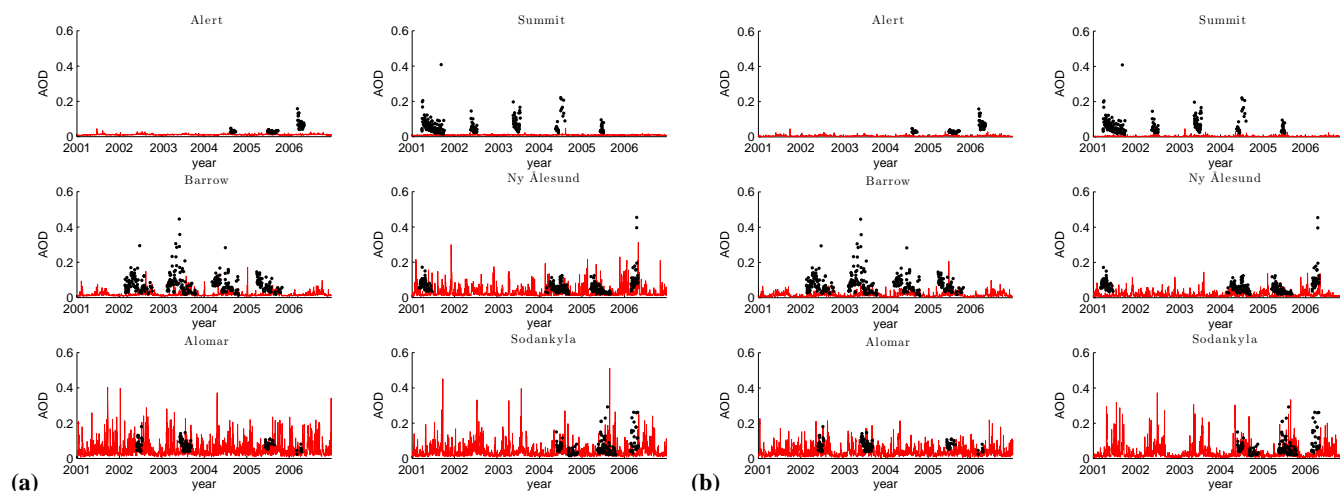


Fig. 3. Time series of daily mean of AOD (550 nm) at the indicated Arctic sites. The black symbols are the observations and the continuous lines the model results: panel (a) ECHAM5-HAM (EIPCC emissions) results; panel (b) results obtained using TM5. The station data for AOD were measured at 500 nm and interpolated to 550 nm using the observed Ångström parameters.

average. The figure also reports the results from the MACC reanalysis and the MODIS Terra and Aqua climatologies. At some sites (Ny Ålesund, Sodankyla and Barrow) ground-based observations, MODIS and MACC data display similar values. For these sites we observe that the MACC dataset agrees well with the observations, also in terms of seasonality. A significant late spring–early summer maximum of AOD is clearly visible, with maximum average AOD values exceeding 0.1. At these sites, both ECHAM-HAM and TM5 produce values of AOD which are significantly lower than the observations, particularly during the spring maximum. Only at the rather meridional station of Sodankyla the modeled AOD reaches 0.06 for both TM5 and the ECHAM-HAM models, but with a peak in July instead of March–April as for the observations and the MACC reanalysis. The other Scandinavian station, ALOMAR, shows larger differences between station data and satellite data, particularly in spring. In this particular case though, the agreement between station data and modeled AOD is good, particularly for ECHAM-HAM. At Alert the modeled AOD is significantly smaller than station data, MODIS and MACC observations (there are some significant differences also between these datasets, presumably for the lack of reliable satellite data in this area). At Summit, MACC estimates are very different from observed data during the whole year, presumably because there are no MODIS satellite data available for this location. Modeled AODs are significantly lower than both ground-based measurements and the MACC reanalysis. To summarize, the models underestimate AOD, except at Sodankyla and ALOMAR, and they peak in summer, while the observations, generally, show a maximum in spring (as documented also in Shaw (1995)).

It has been suggested in Bourgeois and Bey (2011) that the amount of aerosols transported to the Arctic is sensitive

to the magnitude of wet removal, suggesting a revision of the wet scavenging parameters used in the ECHAM-HAM model. We tested this approach, using the parameter values suggested in that paper, and the resulting AOD climatology is shown as a blue line in Fig. 4. We see a significant increase of AOD at all latitudes, leading to a reasonable agreement of summer AOD values with observations at all sites. Like the observations, the AOD modeled with the Bourgeois and Bey (2011) parametrization displays a peak in spring, but instead of presenting a minimum in winter as suggested by MACC and by the available measurements, it is minimum in summer.

The monthly climatology of the Ångström parameter is reported in Fig. 4b. While the exact values of the Ångström parameter shown here can be compared only with difficulty, as they were all obtained with slightly different methods (a caveat discussed in Tomasi et al. (2007)), comparing their seasonal variations is certainly of interest. While the observed AODs have been found to peak around April–May, the Ångström parameter is found to peak in June–July in the measurements, suggesting that there is a summer shift towards finer particles. The TM5 and ECHAM-HAM results are close to the observations and the reanalysis, particularly in the summer months, at all sites, except for Summit. Both models show a seasonal variation with a single main maximum which is reached around June–July for Barrow, Ny Ålesund, ALOMAR and Sodankyla. TM5 displays a stronger variability than ECHAM-HAM, showing lower values from October to April which are in better agreement with the MACC reanalysis and MODIS data. At Alert and Summit, ECHAM-HAM does not show a summer maximum, which is present for TM5. Introduction of the Bourgeois and Bey (2011) parametrization in ECHAM-HAM leads to the appearance of a summer maximum also at Alert and Summit,

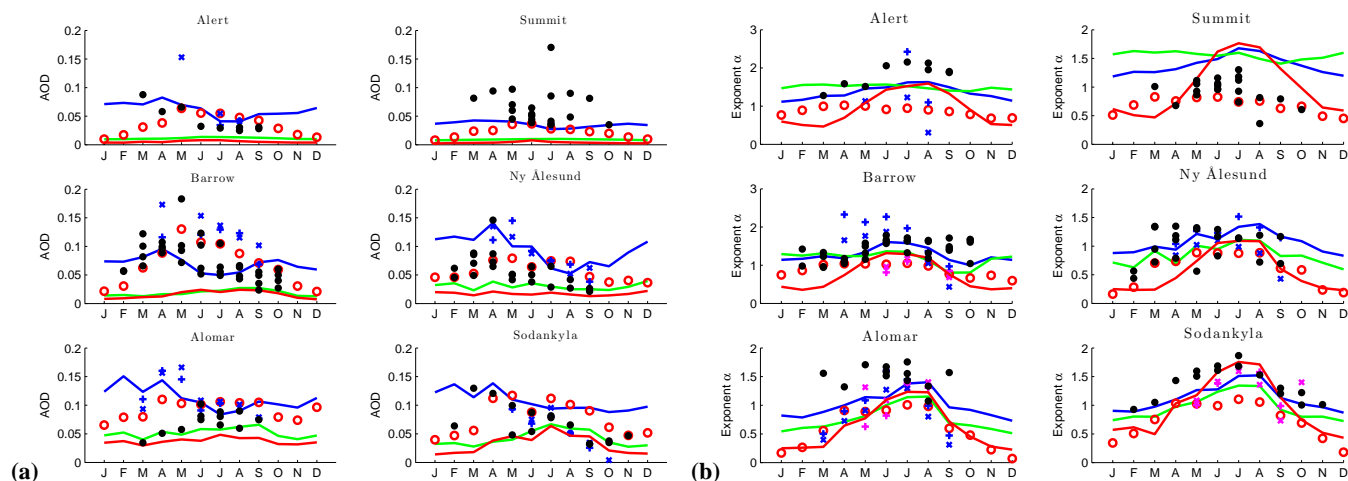


Fig. 4. Observed and modeled seasonal climatologies of AOD at 550 nm (panel (a)) and of the Ångström parameter (panel (b)) at the measurement sites. Observations are shown as black dots (to illustrate interannual variability, we report the individual values for each year). Model results are represented by the continuous lines: ECHAM5-HAM (EIPCC emissions) (green), ECHAM5-HAM with the correction suggested by Bourgeois and Bey (2011) (blue), TM5 (red). MACC data are shown for comparison (red circles) together with MODIS-Aqua (+) and MODIS-Terra (x) observations. For the Ångström parameter both the land (magenta) and the ocean (blue) MODIS products are reported. All model data and MODIS-Terra observations (where available) are averaged over the years 2001–2006. MODIS-Aqua observations are averaged over the years 2002–2007 and MACC data are averaged over the years 2003–2006. The station data for AOD were measured at 500 nm and interpolated to 550 nm using the observed Ångström parameters.

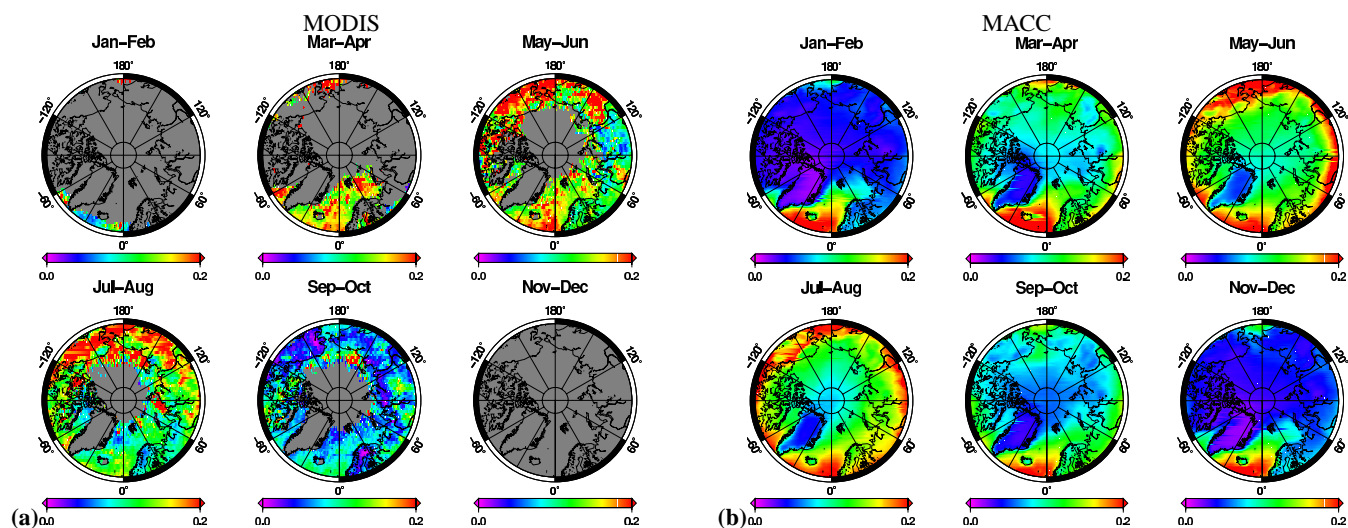


Fig. 5. Maps of average AOD in different periods of the year. Monthly means are calculated in the period 2003 to 2006, common to all datasets. Panels: (a) MODIS Terra, (b) MACC. Grey areas indicate missing data.

even if the modeled values are still far from the observations for all months. At other sites the modification leads only to a marginal increase in the value of the Ångström parameter. To summarize, the Ångström parameter is reasonably reproduced at most stations by both models, indicating that the distribution of particle sizes is captured correctly, together with its seasonality, characterized by a peak in summer. The main exception is the Summit station, located at very high altitude on the Greenland ice sheet, where both TM5 and ECHAM-HAM display an excess of fine particles.

3.3 Spatial distributions

The spatial distributions of the AOD climatologies are reported in Figs. 5, 6 and 7, all averaged over the years 2003–2006, which are common to all datasets. Panels 5a and b report the values of the AOD from MODIS data (Terra) and MACC reanalysis. MODIS data are missing from November to February. In the other months, the MACC reanalysis and MODIS data display a similar spatial structure. Some differences are evident, such as the MACC AODs being lower

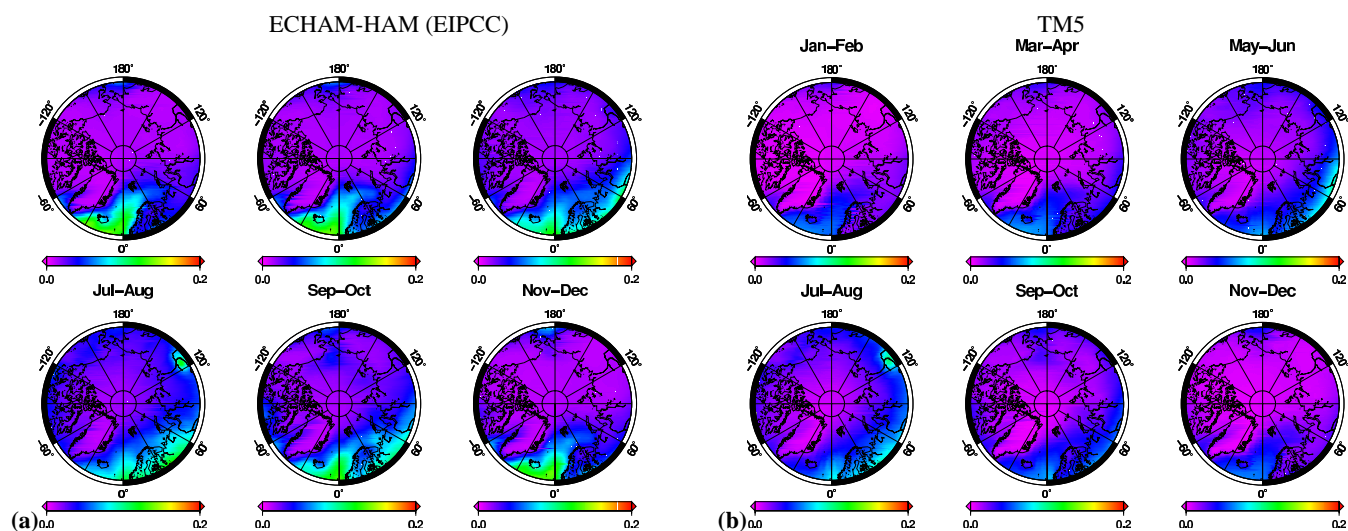


Fig. 6. Maps of average AOD in different periods of the year. Monthly means are calculated in the period 2003 to 2006, common to all datasets. Panels: (a) ECHAM-HAM (EIPCC emissions), (b) TM5.

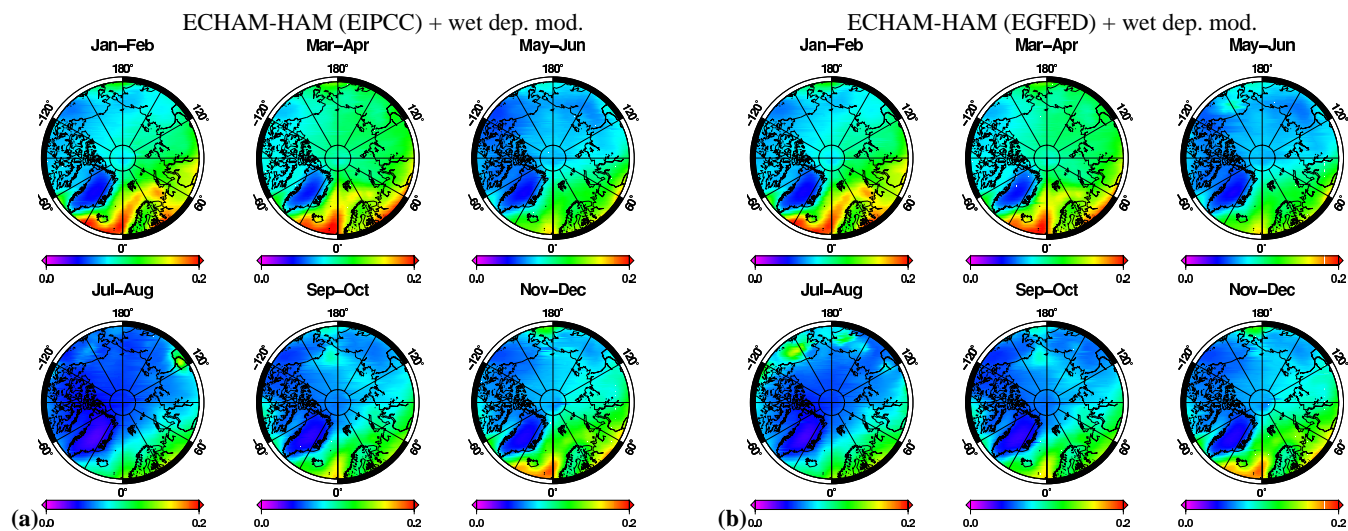


Fig. 7. Maps of average AOD in different periods of the year. Monthly means are calculated in the period 2003 to 2006, common to all datasets. Panels: (a) ECHAM-HAM (EIPCC) with the Bourgeois and Bey (2011) modification. (b) ECHAM-HAM (EG FED) with the Bourgeois and Bey (2011) modification.

than reported by MODIS in Scandinavia in July–August and higher over the Atlantic in July–October; nonetheless there is a general agreement between the two datasets, both in spatial distributions and amplitudes, suggesting that the MACC reanalysis provides a reasonable interpolation of available observations in these areas. Both the satellite and the reanalysis datasets are characterized by a large spatial variability in AOD, with significant differences between the Atlantic and the Pacific sector and an average decrease of AOD with latitude. MACC predicts low AOD values from November to February, which, while they cannot be verified with MODIS observations, find some confirmation from station data (see Fig. 4). For these reasons, in the following we are encour-

aged to use the MACC reanalysis as a reference to compare with modeled AOD distributions.

Figure 6 reports the AOD climatologies observed for ECHAM (EIPCC emissions) and TM5. A severe underestimation of AOD over the entire Arctic area and in all months is apparent for both models. A high AOD tongue over the Atlantic in all months is reproduced, particularly in ECHAM-HAM, but its amplitude is too weak when compared with MACC and with MODIS from March to June. This suggests that sea-salt aerosol emissions are possibly well reproduced but that the AOD associated with other components is underestimated. Spatial maps of the AOD due to only sea-salt (not shown) show an intense contribution of the Atlantic which

peaks in December–January and confirm this conclusion. The contribution of the Atlantic in winter is much weaker in TM5 than in ECHAM-HAM possibly due to differences in the emission parametrization. Increased values of AOD over northern America and Russia in May–August, possibly associated mainly with fires, are present as spatial features in both model runs, but underestimated in amplitude. Spatial maps of the contribution of organic carbon (not shown) confirm this view showing contributions over Siberia and Northern Canada in these months. Overall, as already seen for the monthly climatologies at the station locations, the modeled data do not show the observed seasonality of Arctic AOD, characterized by a pronounced peak in late spring and early summer. Average amplitudes over large areas are underestimated by almost one order of magnitude in some months.

Introduction of the Bourgeois and Bey (2011) modification, in Fig. 7a, leads in general to larger amplitudes which are, in all seasons, closer to the observed (MODIS) and reanalysis (MACC) data. However, the spatial distribution and seasonality still display significant points of disagreement with the observations. Very high Atlantic AOD from January to March compares well with the MACC reanalysis (but does not find strong support in MODIS observations), and AOD over Scandinavia and western Russia appears to be better reproduced. The high AOD observed over northern America and east Russia in May to August is underestimated by more than a factor of two. The spatial structure observed in the data, with high AOD concentrated particularly over the continents, is totally absent in the model results. The spatial maps show clearly how the modeled AOD reaches a minimum in summer, while the extreme values observed in late winter/spring are still underestimated. Overall the seasonal variability seems lower than observed, compare for example the change between July–August and September–November with the MODIS observations.

Previous works (Law and Stohl, 2007; Stohl, 2006; Stocks et al., 1998) suggest that fires in spring and summer in these regions can be the dominant source of sulfur and black and organic carbon. In Fig. 7b we report also results from the ECHAM-HAM model using the EGFED emissions, which include a better representation of fires in these regions and interannual variability. The overall picture does not change much, even if some patches of higher AOD can now be found on the continental masses of North America and Russia in summer.

3.4 Sensitivity to emissions, resolution and nudging

We further explore the ECHAM-HAM model sensitivity to changes in the emission databases in Fig. 8, which reports, for the different emissions, the monthly AOD climatologies at the measurements sites for which ground-based data are available. We find clearly that, while these emission databases differ significantly in terms of annual average emissions (see the budgets reported in Table 1), their

impact on Arctic AOD is scarcely significant in the model simulations, irrespective of whether or not the Bourgeois and Bey (2011) modification is used. Neither a better representation of Arctic fire emissions (EGFED), including interannual variability, nor the increased emissions in Southeast Asia, do impact the modeled Arctic AOD. In principle it is possible that interannual variations in atmospheric circulation may interact with the timing of major fires in the Arctic, leading to differences in summer aerosol concentrations. Our results indicate that this is not the case. A simulation (not shown) using a monthly climatology based on the EGFED emissions leads to results which are almost indistinguishable from the EGFED simulation with interannual variations.

For comparison we also report the results obtained (for the same years) from a climatological, non-nudged ECHAM-HAM simulation using the EAERO emissions. The results display some small differences with respect to the nudged runs, particularly at the European stations of Ny Ålesund, ALOMAR and Sodankyla, in keeping with the view that the details of the atmospheric circulation and of the associated transport processes influence the aerosol distribution. However, the differences between the nudged and non-nudged runs are rather small, and in all cases the simulated AOD remains smaller than the observations at all sites.

Two recent publications (Kaiser et al., 2012; Huijnen et al., 2012) recommend to enhance particulate emissions from wildfires with a global corrective factor of 3.4 in order to compensate for an observed discrepancy between bottom-up and top-down aerosol emission estimates. We tested this suggestion performing an experiment with the EGFED emissions, in which OC and BC emissions are multiplied by this factor, also reported in Fig. 8. We find a significant increase from June to August for Barrow, where modeled AOD becomes comparable with observations. Indeed in this area summer fires in Siberia, Alaska and Canada can have a significant impact. The differences compared to the other simulations remain extremely small at all other stations.

Finally, the role of model resolution is also explored, considering a nudged ECHAM-HAM simulation (with EAERO emissions) at a higher horizontal and vertical resolution (T63 with 31 vertical levels), also reported in Fig. 8, averaged for the years 2001 to 2003. We find that, while there are some differences in the average AOD observed in individual months (possibly due to the shorter period over which the climatology has been computed), these do not cure in any way the underestimation of the AOD which we already discussed. This result is in agreement with Shindell et al. (2008), where it has been shown that, in an intermodel comparison, horizontal model resolution does not show any clear effect on aerosol transport towards the Arctic.

3.5 Contribution of aerosol compounds

To get an insight into the contribution of individual aerosol compounds, Fig. 9a–d report the average over the entire

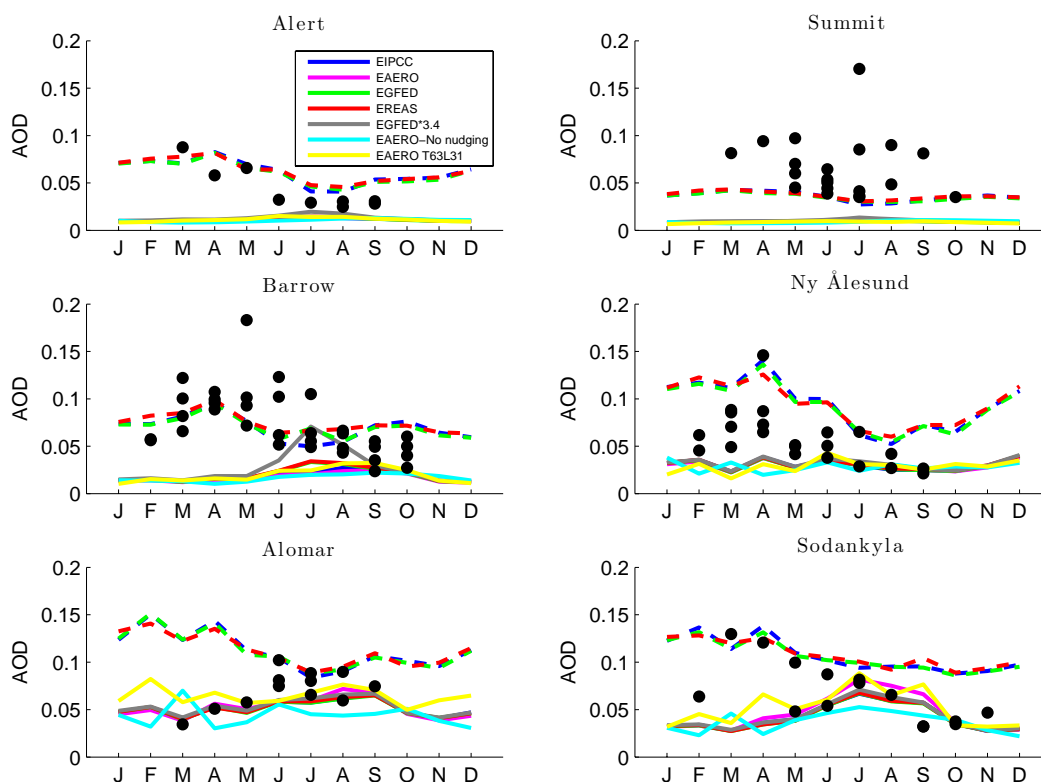


Fig. 8. Comparison of AOD modeled with ECHAM-HAM for different emission databases, averaged over the years 2001–2006. Emissions used are EAERO (continuous magenta line), EIPCC (blue), EGFED (green), EREAS (red). The same emissions, using also the Bourgeois and Bey (2011) parametrization are dashed. A simulation with the EGFED dataset, in which OC and BC emissions from wildfires have been multiplied by a factor 3.4, is reported in gray. A simulation with the EAERO emissions and without nudging is reported in cyan. We also report one simulation (ECHAM-HAM, EAERO) at a higher resolution of T63L31 (in yellow- averaged over the years 2001–2003). Ground-based station observations are reported with black dots as in Fig. 4.

Arctic (defined as the area north of 60° N) of the total AOD and of the fraction associated with each aerosol compound (organic carbon, black carbon, sea salt, sulfur and dust). For MACC and TM5, ambient aerosol contributions to total AOD are reported (i.e. the contribution of their water content is included) while for ECHAM-HAM we report dry compound contributions and diagnosed aerosol water as a separate component. As the MACC results (Fig. 9a) clearly show, there is a predominant role of sulfates in determining the total AOD, accounting for about half of it. An “Arctic-haze” spring maximum in April–May is clearly visible, with a progressive decay during summer, reaching a minimum in December. Sea salt has an opposite seasonality, reaching a minimum in summer and a maximum in winter, compatible with a contribution of Atlantic sea salt emissions in winter which we have described above, possibly due to stronger mesoscale perturbations. Organic carbon peaks and contributes significantly to total AOD from May to August. As already mentioned, spatial maps of its distribution (not shown) confirm an association with wildfires in Siberia and northern Canada in that period. Also black carbon peaks in summer, but shows

a secondary peak in the Arctic haze months of April and May. There is also a significant contribution from dust in the MACC reanalysis. We can compare these results with the AODs attributed to individual compounds in the TM5 simulation in Fig. 9b (note the different vertical scale). For reference, in this figure we also report AOD due to water in aerosols, even if the shown AOD fractions are for ambient aerosols. We see that, as already discussed, the total average AOD is underestimated by a factor of more than four and there are significant differences in seasonality and variability. Sulfates still play the dominant role, accounting for almost half of the observed total AOD in summer, but instead of reaching a maximum in April–May, they peak in summer, from June to August. Like the other compounds their AOD is significantly smaller than reported by MACC. Organic and black carbon peak in July and August, showing a seasonality similar to MACC, which can be understood in terms of seasonality of wildfires (as confirmed by spatial maps of its distribution, not shown). The contribution of dust in TM5 is significantly lower than in MACC. Sea salt shows the same seasonality as in MACC but also these values are lower by

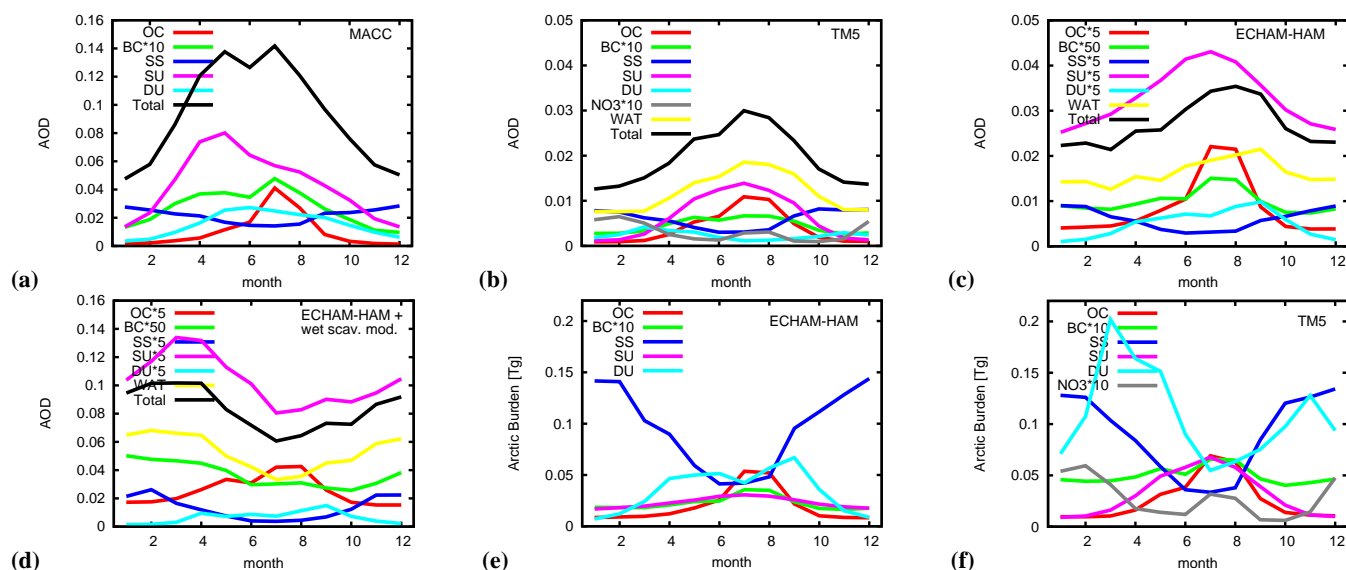


Fig. 9. (a) Averages over the Arctic (north of 60° N) of the contribution to total AOD of different compounds for the MACC reanalysis (these are ambient aerosol contributions, i.e. including water); (b) average AOD of different compounds in the TM5 model (ambient aerosols). (c) the same for the ECHAM-HAM (EIPCC) simulation (a direct comparison is not possible since diagnosed aerosol water is treated as a separate compound in HAM diagnostics and the dry aerosol contributions are shown); (d) the same as (c) for the ECHAM-HAM (EIPCC) simulation with the (Bourgeois and Bey, 2011) wet scavenging modification. (e) average total burdens (in Tg) over the Arctic for the ECHAM-HAM (EIPCC) simulation. (f) average total burdens (in Tg) over the Arctic for the TM5 simulation. A factor 1.4 has been used to convert particulate organic matter to organic carbon for TM5. MACC data are averaged over the years 2003–2006. TM5 and ECHAM-HAM data are averaged over the years 2001–2006. AOD and burdens of black carbon have been multiplied by 10 for clarity in all panels. An additional factor of 5 has been applied to show the seasonalities of dry AOD contributions of individual compounds for ECHAM-HAM in panels (c) and (d).

about a factor of three. Nitrates provide a very small contribution to total AOD, comparable to that of black carbon in winter. The aerosol water contribution to total AOD, shown in the figure, appears to be dominated mainly by the seasonality of sulfates, with a contribution by sea-salt in winter. These effects in terms of AOD are associated with corresponding seasonalities in terms of total Arctic loads (burdens) which we report in Fig. 9f. The seasonal cycles of sea-salt, organic carbon, nitrate and sulfate burdens correspond to similar cycles in terms of AOD. While a significant dust load is present, we have seen above that it does not contribute much to the total AOD. Overall the partition in terms of the contribution of individual aerosol compounds to total AOD is similar between TM5 and MACC, leading to the conclusion that (provided the MACC reanalysis is reliable), the underestimation of total AOD may be due to mechanisms which affect all compounds, such as errors in long range transport or underestimated aerosol residence times for this model.

The AOD contributions of individual compounds for the ECHAM-HAM simulations (using EIPCC emissions) are reported in Fig. 9c and the corresponding Arctic burdens are reported in Fig. 9e. In this case, since aerosol water content is treated as an additional optically active compound in HAM and its contribution is saved separately in the diagnostics, we report the dry contributions to the total AOD of each compound and, separately, the contribution of diagnosed aerosol

water. The total AOD, while on average comparable to that reported by TM5 and much smaller than the MACC average, presents a smaller seasonal variability compared to MACC. Also in this case AOD peaks in August, while high values in April are absent. Sulfates follow the same pattern, with a weak summer maximum and no sign of a spring peak. We find again that organic and black carbon are concentrated in summer months. The total burdens, reported in Fig. 9e are, compared to TM5, lower in terms of sulfates, black carbon and dust and similar for organic carbon and sea salt.

If we introduce the Bourgeois and Bey (2011) wet scavenging modification, total AODs change significantly, as reported in Fig. 9d. The average value is now comparable to what reported by MACC, but with a very different seasonality. The peak in AOD has now shifted to February–March, while July corresponds to a minimum. Aerosol water and sulfates follow the same pattern. Organic carbon still peaks in summer, while black carbon AOD and concentrations (not shown) reach now a peak in winter and a minimum in summer. Overall the wet scavenging modification, while leading to more realistic values of Arctic AOD as a time average, leads to a different seasonality in which a summer minimum in AOD appears, instead of a minimum in winter months as indicated by the MACC reanalysis. This leads to significantly higher AOD estimates in winter and very low values in summer. The seasonal variability in total AOD produced

by ECHAM-HAM with the wet scavenging modification remains smaller than indicated by MACC.

4 Discussion and conclusions

An underestimation of modeled concentrations of sulfates and black carbon in the Arctic has already been evidenced in Shindell et al. (2008) for several current models, including TM5 and ECHAM-HAMMOZ (ECHAM-HAM + gas phase chemistry module MOZART; Pozzoli et al. (2008)). Comparison with a large number of other models in that study has allowed to determine a great diversity in model results, attributed mainly to differences in aerosol physical and chemical processing mechanisms, while emissions and transport processes play a minor role. Particularly the bad reproduction of black carbon surface concentrations at Alert and Barrow shown in that study, suggested that there may exist also issues in the interpretation and comparison of measurements and modeled concentrations. Low BC concentrations in the Arctic and a wide model diversity have also been reported in the AeroCom model comparison in Koch et al. (2009). In Textor et al. (2006) ECHAM-HAM and TM5 show significantly lower total polar burdens of aerosols, for all compounds, compared with other AeroCom models. Bourgeois and Bey (2011) verified that ECHAM-HAMMOZ underestimates the export of aerosols from the planetary boundary layer to the free troposphere and consequently underestimates long-range transport from polluted areas in the mid-latitudes to the Arctic. In comparison the model is successful in reproducing long-range transport of CO, suggesting that dry transport processes are reasonably well modeled. They suggested a reduction of the wet scavenging parameters in the model, finding significant improvement.

In this work, we explored the skill of TM5 and ECHAM-HAM in these areas in deeper detail, focusing on the optical parameters AOD and Ångström, which are directly observed by ground-based and satellite measurements, comparing both their spatial structure and the amplitudes measured at the location of six Arctic stations. The Ångström parameter is reasonably reproduced at most stations by both models, indicating that the distribution of particle sizes is captured correctly, together with its seasonality, characterized by a peak in summer. The main exception is the Summit station where both TM5 and ECHAM-HAM overestimate the parameter; this station is located at high altitude on the Greenland ice sheet, so that orographic effects and model resolution may play a role. The AOD results from in-situ stations show a severe underestimation in amplitude of observed values and an absence of the strong seasonality found in the observations, with a peak in summer rather than in late spring. This underestimation is also confirmed by spatial maps of AOD compared with currently available satellite and reanalysis products. The latter are highly uncertain in polar areas, but together with the station measurements they confirm an under-

estimation of AODs and burdens by the models over most of the Arctic. Spatial maps show also that some observed features, such as the appearance of areas with very high AOD over north America and northern Russia from May to August, are not reproduced in the model simulations. While specific properties of individual measurement stations such as their position and elevation may of course play a role, the satellite and reanalysis comparisons suggest that the problem is mainly over the entire region and that other mechanisms, such as emissions, transport and scavenging processes, may be at the origin of this mismatch. We verified, for ECHAM-HAM, that changes in emission databases or in model resolution do not have a significant impact on the Arctic distributions of modeled AOD. The only exception is an experiment following Kaiser et al. (2012), in which particulate emissions from wildfires were enhanced with a corrective factor, which provides AOD values comparable to station observations, but only for Barrows in summer and with little effect for other stations. There are also no significant differences if the model is free to run in a climatological mode or forced by nudging to follow observed winds. When we test the modification of wet scavenging suggested by Bourgeois and Bey (2011) we find much higher AODs, comparable in a yearly and spatially averaged sense with observations, but still with very different seasonality and spatial structure. In particular, values in autumn and winter appear too high, and a summer minimum appears, while areas with high AOD over northern Russia and America in summer are not reproduced. These differences in seasonality may be linked with the higher summer precipitation in ECHAM compared to MACC and ERA-Interim, reported in Sect. 3.1. While the wet-scavenging parameters are reduced in the modification, summer precipitation at high latitudes can still contribute to significant wet scavenging, leading to low summer aerosol concentrations.

The results reported above indicate that changes in wet scavenging, while they do not produce a correct seasonality and spatial aerosol distribution, allow to reach yearly averaged values of AOD in the Arctic which are more realistic. Efforts towards a better representation of aerosols in the Arctic, for ECHAM-HAM and possibly TM5, should aim at improving the representation of features and processes involving aerosol removal and transformation. In particular, two aspects should be considered: On one hand, the atmospheric models have to be capable, at high latitudes, to represent correctly physical features related to wet removal, such as cloud coverage, structure and precipitation. On the other hand current parametrizations for wet removal have to be further developed and tuned, taking into account some of the specificities of the Arctic environment. Several improvements that may be beneficial are being developed and need to be investigated in detail. A recent study (Browse et al., 2012) associates the seasonal cycle of Arctic aerosols with scavenging processes, in particular with the passage from inefficient scavenging of soluble aerosols from ice clouds in winter to more efficient scavenging from low, warm liquid clouds in summer.

Croft et al. (2010) compared different in-cloud aerosol scavenging parametrization schemes for ECHAM-HAM, finding large variabilities in aerosol mass and number burdens, and improved agreement with recent diagnostic and prognostic scavenging schemes, particularly at high latitudes. Improved schemes for below-cloud scavenging by snow in ECHAM-HAM are discussed in Croft et al. (2009).

Another source of discrepancy can be related to differences in transport mechanisms between the models and the real atmosphere. While large-scale winds are similar, mesoscale dynamics and transport barriers are poorly represented by the global models considered here. Future developments should thus also consider aerosol dynamics in high resolution regional and global models capable of describing mesoscale processes.

It is important to notice that a comparison between aerosol measurements and model results in the Arctic area is particularly difficult. Indeed, aerosol measurements are particularly challenging and scarce in polar areas. As shown in this work, ground-based, satellite and reanalysis datasets may display significant differences which can be caused by insufficient statistics, uncertainty associated with the measurements, issues in their interpretation and conversion, and differences in spatial and temporal resolution. We conclude noting that the spatial maps of observed AOD, from satellites and reanalysis, show a large spatial variability of this parameter over the Arctic, suggesting that measurements at single stations can be hardly representative of larger areas. On the other hand, ground-based measurements are essential to calibrate and validate satellite measurements and reanalysis datasets. Due to the impact of changes in the Arctic for the global climate, the extensions of the currently available measurement network, the collection and availability of long-term measurements and the exploration of methods to bound data reliability are important future goals.

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