



Asian dust storm observed at a rural mountain site in southern China: chemical evolution and heterogeneous photochemistry

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Abstract. Heterogeneous processes on dust particles are important for understanding the chemistry and radiative balance of the atmosphere. This paper investigates an intense Asian dust storm episode observed at Mount Heng (1269 m a.s.l.) in southern China on 24–26 April 2009. A set of aerosol and trace gas data collected during the study was analyzed to investigate their chemical evolution and heterogeneous photochemistry as the dust traveled to southern China. Results show that the mineral dust arriving at Mt. Heng experienced significant modifications during transport, with large enrichments in secondary species (sulfate, nitrate, and ammonium) compared with the dust composition collected at an upwind mountain top site (Mount Hua). A photochemical age “clock” ($-\text{Log}_{10}(\text{NO}_x/\text{NO}_y)$) was employed to quantify the atmospheric processing time. The result indicates an obvious increase in the abundance of secondary water-soluble ions in dust particles with the air mass atmospheric processing time. Based on the observations, a 4-stage evolution process is proposed for carbonate-containing Asian dust, starting from fresh dust to particles coated with hydrophilic and acidic materials. Daytime-enhanced nitrite formation on the dust particles was also observed, which indicates the recent laboratory result of the TiO_2 photocatalysis of NO_2 as a potential source of nitrite and nitrous acid.

1 Introduction

Mineral dust is injected into the atmosphere under specific meteorological conditions in deserts or semiarid areas, generating dust storms. In East Asia, the Taklimakan Desert in western China and the Gobi Desert in Mongolia and northern China are the two major source regions of dust, with annual emissions between 100 Mt yr^{-1} and 460 Mt yr^{-1} (Laurent et al., 2006). Upon entering the atmosphere, especially when transported over heavily polluted regions, the chemical and surface nature of mineral dust particles undergo significant changes, in turn influencing a number of atmospheric processes (Zhuang et al., 1992; Usher et al., 2003; Formenti et al., 2011, and the references therein).

In the past few decades, several large field campaigns (e.g., TRACP, ACE-Asia and INTEX-B) have studied the atmospheric processes of Asian dust (Jacob et al., 2003; Arimoto et al., 2006; Singh et al., 2009). The evolutionary processes of dust particles during their eastward transport to the western Pacific have been already investigated in north-east Asia. These studies have significantly improved understanding of the heterogeneous reactions between Asian pollution and mineral dust. Sullivan et al. (2007a) found that nitrate tended to accumulate on calcium-rich dust, while sulfate tended to accumulate on aluminosilicate-rich dust. Li and Shao (2009) demonstrated nitrate being coated onto dust particles. The direct uptake of chlorine on dust particles by

heterogeneous reaction with HCl was clearly verified by Sullivan et al. (2007b). Sullivan and Prather (2007) observed internally mixed organic dicarboxylic acid (DCA) with dust particles and proposed a mechanism of heterogeneous oxidation of DCA on the surface of Asian dust. To our knowledge, there have been very few studies of the chemical evolution of Asian dust during its transport to southern China, where the dust is subject to higher temperature, higher humidity, and possibly different emissions compared with eastward transport.

Heterogeneous reactions on dust particles have been extensively investigated, and most previous studies have focused on nighttime chemistry (Usher et al., 2003). Photo-induced/enhanced reactions of some reactive gases on dust particles have recently been demonstrated by laboratory studies (Cwiertny et al., 2008 and the references therein). Nicolas et al. (2009) observed photo-enhanced O_3 decomposition on mimic mineral dust particles of TiO_2/SiO_2 mixtures; Ndour et al. (2008) detected the photo-enhanced uptake of NO_2 on both TiO_2/SiO_2 mixed particles and real Saharan and Arizona dust particles. The latter reaction received great attention because it may be a missing source of daytime nitrous acid (HONO). However, the extent of that reaction has not been observed in the real atmosphere.

As part of China's National Basic Research Project on acid rain, a comprehensive field campaign was conducted from March to May 2009 at Mt. Heng in Hunan province, southern China (Fig. 1), during which an intense dust storm was observed on 24–26 April. The aim of the specific work is to investigate the impact of this Asian dust storm on the atmospheric chemistry at rural and remote mountain regions in southern China. Special attention is given to the chemical evolution of the dust by comparison with the aerosol composition concurrently observed at another mountain site near the dust source region (Wang et al., 2011), as well as the role of heterogeneous photochemistry in the production of nitrite. The general features and transport pathway of this dust storm are described first. The chemical evolution of the dust particles during transport from Mt. Hua to Mt. Heng is then investigated. A photochemical age “clock” ($-\log_{10}(NO_x/NO_y)$) is employed to examine the changes in the major secondary water-soluble ions with the atmospheric processing time. The evidence of the daytime production of nitrite on dust particles is also presented, and its possible mechanism and potential contribution to the gas-phase HONO is discussed.

2 Experimental methodologies

2.1 Site description

Mt. Heng is situated in Hunan Province, southern China, and is approximately 500 km and 900 km from the South China Sea and the East China Sea, respectively (Figs. 1a and 2). The field campaign was conducted at a meteorological station at

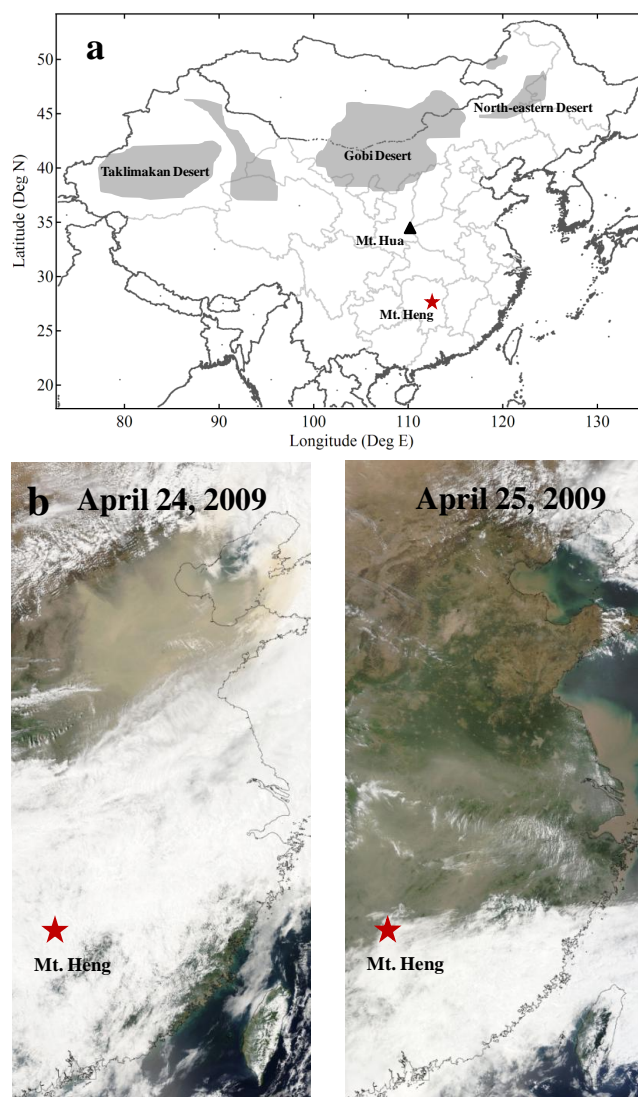


Fig. 1. (a) Map showing Mt. Heng and Mt. Hua. The major deserts in East Asia are marked by the shadows in the map. (b) MODIS true-color imagery of eastern China on 24 and 25 April 2009.

the summit of Mt. Heng ($27^{\circ}18' N$, $112^{\circ}42' E$, 1269 m a.s.l.). There are a few temples and a road around the mountain top of Mt. Heng. Emissions from incense burning and traffic may occasionally affect the trace gas and aerosol measurements at the site. There are few other local emission sources around the measurement site. The closest cities are Hengyang and Xiangtan, which are approximately 50 km to the south and 70 km to the north, respectively.

2.2 Field campaign, sample collection and measurement techniques

The field campaign covered a 3-month period from March to May 2009. A comprehensive suite of trace gases, aerosols, and rain and cloud water samples was measured/collected

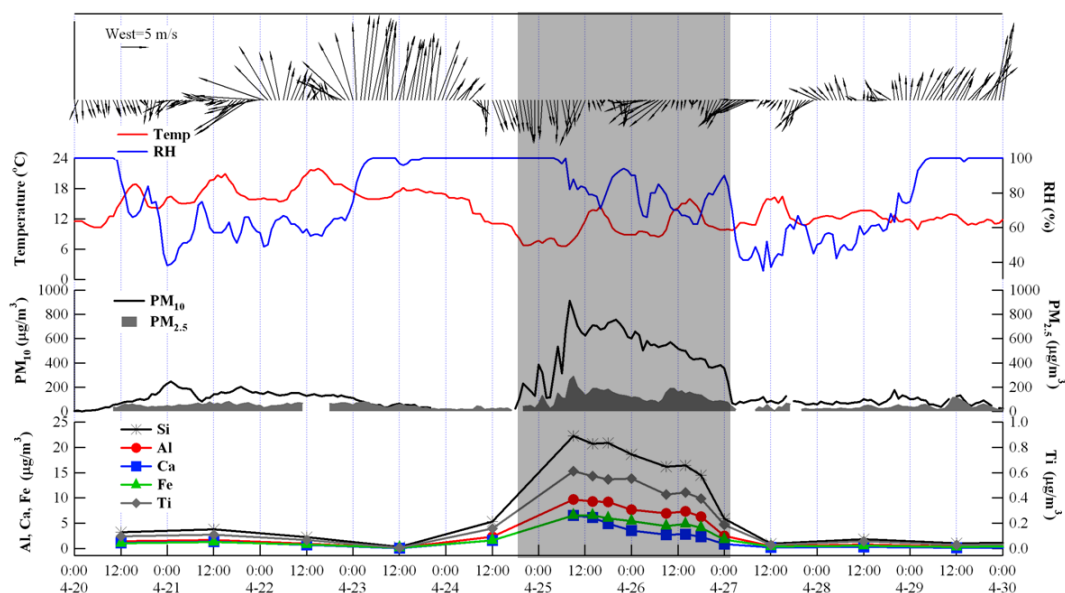


Fig. 2. Temporal variations of PM_{10} , $\text{PM}_{2.5}$, mineral elements (Si, Al, Ca, Fe, Ti) in $\text{PM}_{2.5}$ and meteorological parameters at Mt. Heng during 20 to 29 April 2009.

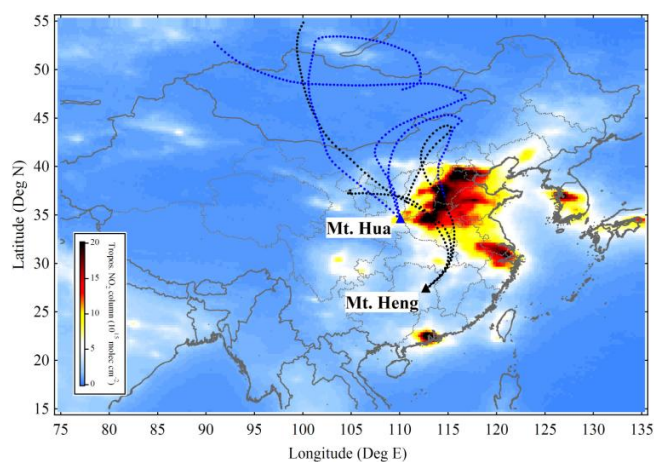


Fig. 3. The 72-h backward trajectories for the dust event at Mt. Hua and Mt. Heng. The NO_2 column data from GOME and SCIAMACHY (averaged over 1996–2007, Wang et al., 2009) are shown in the shaded areas.

(Sun et al., 2010; Gao et al., 2012; Zhou et al., 2012). The present study is focused on the dust storms observed during 20–29 April 2009 (Fig. 3). The dataset employed in this paper includes continuous data of SO_2 , NO_x , NO_y , $\text{PM}_{2.5}$, PM_{10} and black carbon (BC); filter-based element data of Si, Al, Fe, Ca and Ti in $\text{PM}_{2.5}$ and Al in TSP; and size-resolved water-soluble ions (SO_4^{2-} , NO_3^- , NO_2^- , NH_4^+ , Ca^{2+}).

Measurements of trace gases were done as follows: SO_2 was measured with a pulsed UV fluorescence analyzer (TEI model 43C); NO_y (defined as NO_x and its oxidation products and intermediates) was converted to NO with heated

molybdenum oxide (MoO) and measured with a chemiluminescence analyzer (TEI model 42CY); NO_2 was converted to NO with a blue light converter (Meteorologieconsult GmbH) and measured with another NO analyzer (TEI model 42I). The NO_2 conversion efficiencies, determined by the method of gas-phase titration, were approximately 40 % on average during the whole sampling period. The calibration methods followed those reported in Wang et al. (2001). In this study, hourly-averaged data were employed for all the gaseous species and the uncertainties were estimated to be within $\pm 10\%$ (Wang et al., 2001, 2003).

Sampling and analysis of aerosols were done as follows: $\text{PM}_{2.5}$ and PM_{10} mass concentrations were continuously measured with a tapered element oscillating microbalance (TEOM, Thermo Electron Corporation, East Greenbush, NY, USA). BC was measured using an Aethalometer (Magee Scientific, Berkeley, California, USA, Model AE-21) at a wavelength of 880 nm. $\text{PM}_{2.5}$ filter samples were collected using a four-channel particle sampler (Thermo Andersen Chemical Speciation Monitor, RAAS2.5-400, Thermo Electron Corporation) with Teflon filters (TeflonTM, 2 μm pore size and 47 mm diameter, Pall Inc.) at a flow rate of 16.7 liters per minute (LPM) (Wu and Wang, 2007). Total Suspended Particulates (TSP) samples were collected using medium-volume air samplers (TH-105C, Wuhan Tianhong Intelligent Instrument Corp. Co., Ltd., China) at flow rate of 100 LPM on glass fiber filters (GFFs, 90 mm in diameter). The size-resolved particles were collected with a Micro-Orifice Uniform Deposit Impactor (MOUDI, MSP Company) with aluminum substrates (MSP Company) at a flow rate of 30 LPM (Nie et al., 2010). The impactor collected

Table 1. Comparison of chemical species in non-dust and dust storm periods during the Mt. Heng study.

Species	Dust		Non-dust		PM size	Sampling instruments
	Concentrations	Sample numbers	Concentrations	Sample numbers		
PM ₁₀ (μg m ⁻³)	502 ± 196	55	45.6 ± 32.4	1195	PM ₁₀	TEOM
PM _{2.5} (μg m ⁻³)	125 ± 55	54	32.3 ± 22.6	1273	PM _{2.5}	TEOM
PM _{2.5} /PM ₁₀	0.25 ± 0.08	54	0.63 ± 0.23	621		TEOM
Si (μg m ⁻³)	18.5 ± 2.9	7	1.1 ± 0.6	7	PM _{2.5}	RAAS
Al (μg m ⁻³)	8.1 ± 1.3	7	0.47 ± 0.25	7	PM _{2.5}	RAAS
Fe (μg m ⁻³)	5.4 ± 1.0	7	0.39 ± 0.17	7	PM _{2.5}	RAAS
Ca (μg m ⁻³)	4.2 ± 1.7	7	0.20 ± 0.10	7	PM _{2.5}	RAAS
Ti (μg m ⁻³)	0.5 ± 0.1	7	0.03 ± 0.02	7	PM _{2.5}	RAAS
Sulfate (μg m ⁻³)	20.0 ± 3.9	4	13.8 ± 6.1	21	TSP	MOUDI
Nitrate (μg m ⁻³)	16.9 ± 3.4	4	8.2 ± 5.4	21	TSP	MOUDI
NO _y (ppbv)	6.5 ± 2.1	55	3.7 ± 5.4	1505		42CY
SO ₂ (ppbv)	4.2 ± 2.2	55	2.1 ± 2.4	1396		43C
NO _x (ppbv)	2.1 ± 0.7	55	1.0 ± 0.8	671		42I

particles in nine size ranges: > 18 μm (inlet), 10–18 μm, 5.6–10 μm, 3.2–5.6 μm, 1.8–3.2 μm, 1–1.8 μm, 0.56–1 μm, 0.32–0.56 μm, and 0.18–0.32 μm. Submicron particle (PM_{<1}) concentrations are calculated by the sum of last three stages. Supermicron particle (PM_{>1}) concentrations are calculated by the sum of first five stages. The sampled filters were stored at –5 °C in order to minimize artifacts. In this paper, the discussions on the chemical evolution of mineral dust particles focus on the supermicron dust particles (PM_{>1}).

The elements in PM_{2.5} and TSP were measured using X-ray fluorescence (XRF) and inductively coupled plasma mass spectrometry (ICP-MS), respectively. The water-soluble ions were analyzed using ion chromatography (IC) (Dionex 90) (Wu and Wang, 2007). The measurement uncertainties of water-soluble ions and elements were estimated to be ± 10 % (Vecchi et al., 2008; Zhou et al., 2009). The evaporation loss of ammonium nitrate can be ignored due to the low ambient temperature (< 20 °C, Fig. 2) (Nie et al., 2010). During the dust storm of 24–26 April, we collected four sets of size-resolved aerosol samples using MOUDI: sample-I, sample-II, sample-III and sample-IV, with sample-I and sample-III collected during daytime (08:00–20:00 LT) and sample-II and sample-IV at night (20:00–08:00 of next day, LT) (Fig. 5).

3 Results and discussion

3.1 General description of the dust episodes

Figure 2 shows the temporal variations of meteorological parameters, particle mass, and chemical species observed at Mt. Heng during 20–29 April 2009, including wind speed, wind direction, temperature, RH, PM₁₀, PM_{2.5}, and major mineral elements (Si, Al, Ca, Fe, Ti) in PM_{2.5}. During this period, a moderate dust storm was observed on 21–22

April, with hourly averaged PM₁₀ and PM_{2.5} concentrations of 153 μg m⁻³ and 57 μg m⁻³, respectively. A more intense dust storm attacked the sampling site on 24–26 April, as highlighted by the shadow in Fig. 2, with the hourly averaged PM₁₀ and PM_{2.5} concentrations reaching 502 μg m⁻³ and 125 μg m⁻³, respectively. Prevailing northerly winds occurred during the dust storm periods (Fig. 2). The second dust storm on 24–26 April was also observed at three other sites at Mt. Hua in the Shaanxi province of central China (Fig. 1a), Mt. Tai in the Shandong province of central-eastern China (Wang et al., 2011) and the island of Taiwan. This dust event was also clearly shown in the MODIS true-color image for 24 and 25 April (Fig. 1b). These observations indicate a large geographical region of southern and eastern China being influenced by this dust storm. Due to the significant particle loading and more extensive transport, the specific analysis is focused on the second dust storm event (24 to 26 April).

Table 1 compares a set of species/parameters measured in non-dust periods and dust storm periods (24–26 April), including PM₁₀, PM_{2.5}, PM_{2.5}/PM₁₀ ratio, mineral elements (Si, Al, Ca, Fe and Ti) in PM_{2.5}, sulfate and nitrate in TSP, NO_y, NO_x and SO₂. Compared to non-dust periods, all the species in the dust storm event showed higher concentrations: PM₁₀ and mineral elements in PM_{2.5} increased by 10–20 fold; PM_{2.5}/PM₁₀ ratio decreased from 0.63 to 0.25; anthropogenic pollutants (sulfate, nitrate, NO_y, SO₂ and NO_x) increased only moderately, by 0.4–1.2 fold. These results reveal that the dust storm of 24–26 April brought to Mt. Heng high loading mineral dusts that had been mixed with moderate anthropogenic pollutants. Gao et al. (2012) have examined the size distributions of major water-soluble ions in non-dust periods, which showed the presence of sulfate and ammonium as a single peak in the fine particle mode and nitrate as dual peaks in the fine and coarse mode. As to be shown

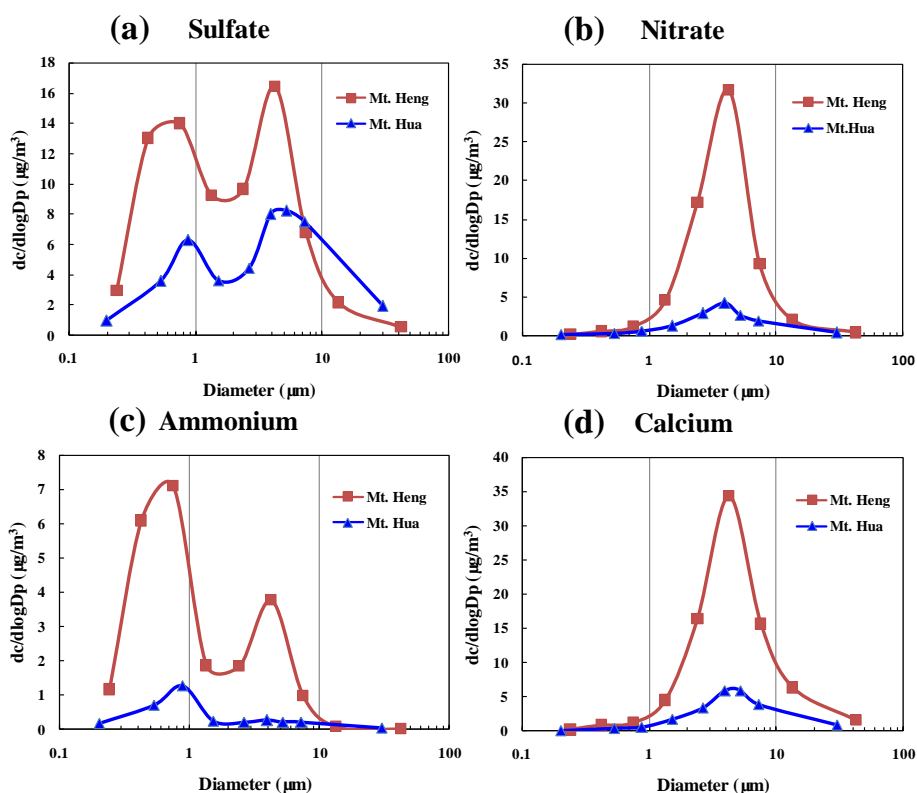


Fig. 4. Size distributions of (a) sulfate, (b) nitrate, (c) ammonium and (d) calcium at Mt. Hua and Mt. Heng during the dust event on 24–26 April.

in Sect. 3.2.1, during the dust storm of 24–26 April, the size distributions of these ions were significantly modified.

In order to characterize the origin and transport pathway of the air masses, a NOAA HYSPLIT4 model (Hybrid Single-Particle Lagrangian Integrated Trajectory) was employed to calculate 72-h backward trajectories arriving at Mt. Hua and Mt. Heng during the dust event of 24–26 April. As illustrated in Fig. 3a, the dust originating in the Gobi Desert moved southward and passed over central China (Shaanxi province), the North China Plain (Hebei province, Shanxi province and Henan province) and southern China (Hubei province and Hunan province) and is expected to travel farther to the South China Sea.

3.2 Chemical evolution of mineral dust particles during transport

3.2.1 Comparison with samples collected near the dust source region

The concurrent campaigns during the dust storm of 24–26 April at Mt. Hua (Wang et al., 2010) and Mt. Heng (this study) provide an opportunity to directly investigate the evolution of mineral dust particles from a near source region in the north to a southern mountainous area. Figure 4 illustrates the size distributions of four major water-soluble

ions, including sulfate, nitrate, ammonium and calcium, at Mt. Hua (sample of 24 April) and Mt. Heng (sample of 25 April). The size distribution patterns of sulfate, nitrate and calcium at both mountain sites were similar: sulfate had a bimodal size distribution, with one peak in the submicron range and another larger peak in the supermicron range; nitrate and calcium showed single peaks in the coarse mode. For ammonium, the size distribution exhibited different patterns at the two sites, being unimodal at Mt. Hua, but bimodal at Mt. Heng. Compared with Mt. Hua, all four ions in the supermicron particles at Mt. Heng showed significant enhancement (Fig. 4): sulfate increased by approximately 56 %, but the other three ions were enhanced up to 5–7 times. The enrichments of water-soluble ions on dust particles at Mt. Heng can be generally attributed to the injection of anthropogenic emissions as the dust plumes passed over the polluted North China Plains (see Fig. 3).

The different degrees of enrichment for the different ions at Mt. Heng may be influenced by the relative abundance of SO_2 , NO_x and NH_3 in the atmosphere and the factors affecting their uptake/conversion on dust particles. Due to a higher hygroscopicity of $\text{Ca}(\text{NO}_3)_2$ than CaSO_4 , the formation of $\text{Ca}(\text{NO}_3)_2$ would significantly enhance the uptake of water (and other soluble gases), resulting in a positive feedback until all the calcium transformed to calcium nitrate (Formenti

et al., 2011). This process may offer an explanation for the much larger increase in nitrate than that of sulfate on the mineral particles during the transport from Mt. Hua to Mt. Heng (Fig. 4). The enrichment of ammonium will be discussed in Sect. 3.2.3.

3.2.2 Changes of mineral dust characteristics with atmospheric processing time

In order to further investigate the evolution of the mineral dust, the atmospheric processing time of the air masses was calculated by the photochemical age ratio $-\text{Log}_{10}(\text{NO}_x/\text{NO}_y)$ (Kleinman et al., 2008), and then was related to the changes in water-soluble ions on the mineral dust particles. As illustrated in Fig. 5a, the photochemical age generally showed an upward trend during the dust event (from 19:00, 24 April to 08:00, 27 April, LT). The sudden decrease in the values of $-\text{Log}_{10}(\text{NO}_x/\text{NO}_y)$ was related to the changes of northerly wind to easterly wind (Fig. 2), which brought freshly emitted NO_x from the incense burning in the nearby temples. The photochemical ages of the first three samples increased gradually, while the values of the sample-III and sample-IV were similar (approximately 0.7).

The ratios of sulfate, nitrate and ammonium to the PM mass in the supermicron range are shown in Figs. 5b. These results indicate overall increases in the relative abundances of all three ions with the atmospheric processing time of the air parcels. During the time scale of the four samples, sulfate increased from 1.7 to 2.4 %, nitrate from 2.5 to 4.3 %, and ammonium from 0.2 to 1.2 %. The increased water-soluble ions could significantly enhance the solubility of mineral particles, thus affecting their lifetimes and cloud condensation nuclei (CCN) or ice nuclei (IC) activity. It is worth noting that the large increases in the abundance of the secondary ions in sample III and sample IV with similar photochemical age indicate the contribution of additional factor(s) to the formation of these ions.

Acidity is another important factor that affects the chemical and physical properties of aerosols (Jang et al., 2002). For Asian dust particles, calcium is the major cation. Therefore, the ratios of the sum of major anions to calcium in equivalence in the supermicron range particles were calculated to investigate the changes in acidity of the mineral particles with the atmospheric processing time of the air mass. As illustrated in Fig. 5c, evident acidification occurred with the aging of the air mass. The ratio increased from 0.48 in the first sample to 0.96 in the third sample and to 1.20 in the fourth sample.

3.2.3 Ammonium enrichment and a conceptual model for dust evolution

In our study, large enrichments in ammonium on supermicron range particles were observed at Mt. Heng. The results show an absence of coarse-mode ammonium at Mt. Hua

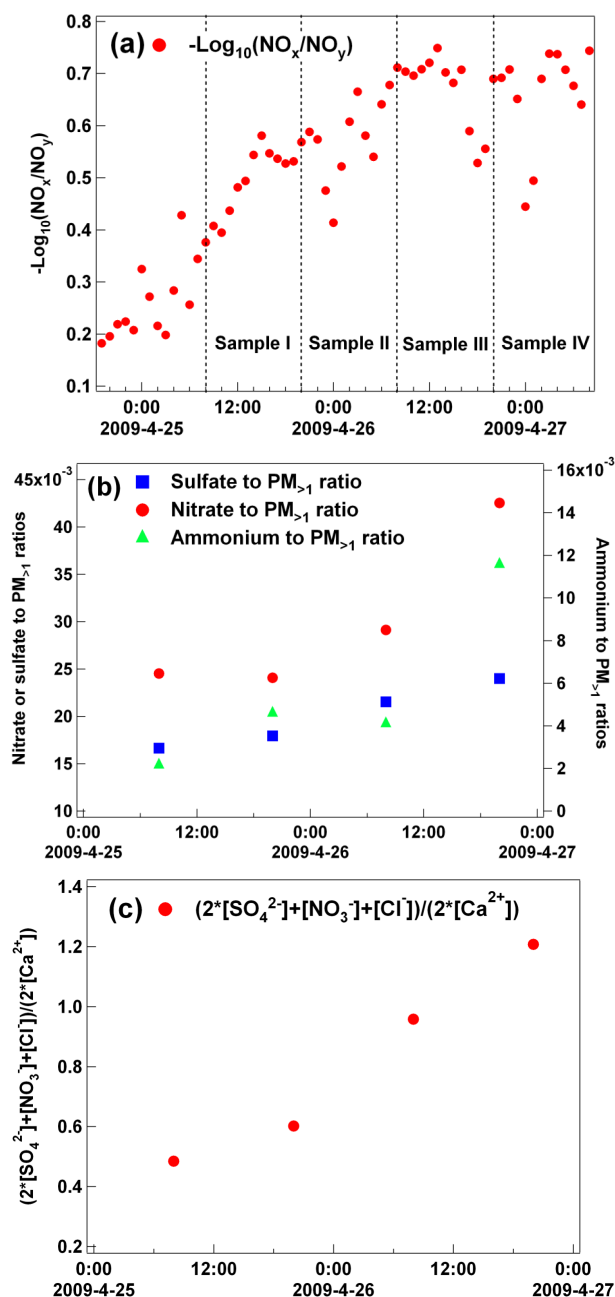


Fig. 5. (a) Photochemical age (defined as $-\text{Log}_{10}(\text{NO}_x/\text{NO}_y)$), (b) concentration ratios of sulfate, nitrate and ammonium to PM in the supermicron particles, (c) sum of major anions to calcium in equivalence in the supermicron particles, during the dust event on 24–26 April at Mt. Heng.

(Fig. 4c), an increase to a content of 2–5 % in sample I to sample III, and a further increase to ~ 12 % in sample IV at Mt. Heng (Fig. 5c). These observations are in contrast to previous that which suggest suppression of the formation of ammonium aerosol by the invasion of a dust storm due to dilution of the precursor ammonia and the uptake of acidic gases on dust (Song and Carmichael, 1999; Huang et al., 2010).

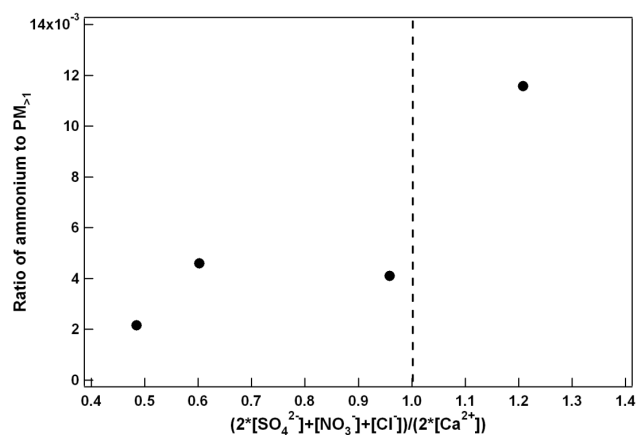


Fig. 6. Scatter plots of ratios of NH_4^+/PM and $(2 \times [\text{SO}_4^{2-}] + [\text{NO}_3^-] + [\text{Cl}^-]) / (2 \times [\text{Ca}^{2+}])$ in $\text{PM}_{>1}$ during the dust event on 24–26 April at Mt. Heng.

Previous field observations have indicated that, for ammonium to form, mineral dust should first take up some acidic species that ammonia can partition to (Sullivan et al., 2007a). Carbonate-containing dust particles have a high alkalinity initially, and the uptake of secondary acids would first neutralize the alkaline carbonate and uptake of ammonia would follow. Thus some relationship between the amount of accumulated ammonium and the neutralized carbonate would be expected. In Asian dust, carbonates are mainly combined with calcium in the form of calcite (Matsuki et al., 2005). Figure 6 shows the relationship of ammonium abundances and the ratio of the sum of anions to calcium, which can be used to indicate the degree of neutralization of carbonate in the dust particles of Mt. Heng. It shows that the alkaline carbonates were not fully neutralized in the first three samples ($(2 \times [\text{SO}_4^{2-}] + [\text{NO}_3^-] + [\text{Cl}^-]) / (2 \times [\text{Ca}^{2+}]) < 1$), but some ammonium was still observed in the supermicron mode. This observation can be explained as follows: after the depletion of surface carbonate, the absorbed acidic species would uptake the gas-phase ammonia in parallel with neutralizing the carbonate in the inner core. Then, after the carbonate was fully consumed, ammonium was enhanced significantly in the fourth sample (Fig. 6).

Based on the above discussions, a conceptual model is proposed to describe dust evolution into a four-stage process (Fig. 7): freshly emitted dust (stage 1), the uptake of secondary acidic species with little ammonium (stage 2, as represented by the sample of Mt. Hua; see Fig. 4c), dust with ammonium accumulated on but without the depletion of carbonate (stage 3, represented by sample-I to III of Mt. Heng; see Figs. 5c and 6), and dust with all carbonate neutralized and significantly enhanced ammonium (stage 4, represented by sample-IV of Mt. Heng). This 4-stage evolution has important implications for atmospheric chemistry and the impact of dust on climate. When the Asian dust particles have

reached the third stage (i.e., with accumulated ammonium), the surface properties of the dust have changed from hydrophobic and alkaline to hydrophilic and acidic, thus significantly enhancing their CCN/IN capacity. At the fourth stage, the dust particles would be no different from the secondarily formed particles and could serve as favorable media to promote further reactions. It should be noted that the conceptual model can only apply to the carbonate-containing dust particles and that ambient ammonia concentrations will influence the development of the four stages.

3.3 Photo-enhanced nitrite formation

High nitrite concentrations on dust particles were observed during the dust event of 24–26 April. Figure 8 shows the nitrite concentrations in TSP in the four samples following the temporal order during the dust episode of 24–26 April. The average nitrite concentration during this dust event was $2.5 \mu\text{g m}^{-3}$, much higher than that during the non-dust days (the mean value = $0.3 \mu\text{g m}^{-3}$), indicating an enhancement of nitrite formation on the dust particles. Furthermore, the enhanced nitrite formation during the dust event occurred only in the daytime samples (with a concentration up to $6 \mu\text{g m}^{-3}$ in sample-III; see Fig. 8), suggesting a photochemical pathway yielding the observed nitrite.

Daytime HONO (as well as nitrite) can be produced from gas phase sources and photo-related heterogeneous sources (Kleffmann, 2007). The reaction of NO and OH radicals is the dominant gas phase process yielding HONO. However, due to the fast photolysis of HONO, it has been demonstrated that this reaction is not a net source and cannot explain the frequently observed high daytime HONO concentrations (e.g., Kleffmann, 2007; Sörgel et al., 2011). The insignificance of this gas-phase source in our study is supported by the low NO concentrations (< 0.5 ppbv) observed during the dust event. Therefore, there must be other photo-related heterogeneous processes contributing to the observed daytime nitrite.

Heterogeneous photochemical reactions producing HONO are generally thought to occur via four different processes: the surface photolysis of nitric acid (HNO_3) (Zhou et al., 2003; Ramazan et al., 2004); photo-induced NO_2 conversion on soot surfaces (Aubin and Abbatt, 2007); the heterogeneous photochemistry of NO_2 on some organic surfaces, such as aromatic compounds and humic acids (George et al., 2005; Stemmler et al., 2006); and the surface TiO_2 photocatalysis of NO_2 (Gustafsson et al., 2006; Ndour et al., 2008). In order to determine the key process contributing to the observed nitrite formation, the concentrations of nitrite, nitrate and BC (representing soot) during the dust storm of 24–26 April were compared with those during a non-dust pollution episode (denoted as PE) occurring on 8–9 May (Table 2). Compared to those during PE, the nitrate and BC concentrations during the dust storm were lower, but the nitrite concentrations were much higher.

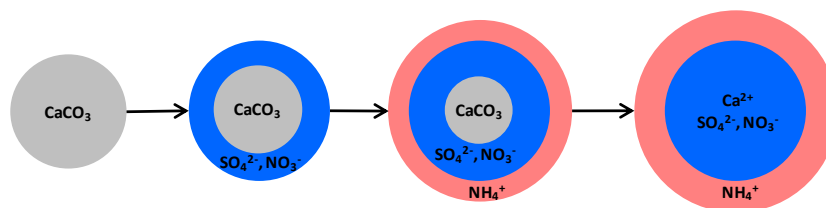


Fig. 7. Conceptual model showing the four aging stages for the carbonate-containing dust particles.

Table 2. Concentrations of nitrite in TSP, nitrate in TSP and BC in $PM_{2.5}$ measured during the dust storm of 24–26 April and a non-dust pollution episode that occurred on 8–9 May at Mt. Heng.

	DS-II on 25–26 April	Pollution episode on 8–9 May
Nitrite in TSP ($\mu\text{g m}^{-3}$)	2.5	0.6
Nitrate in TSP ($\mu\text{g m}^{-3}$)	15.6	20.1
BC in $PM_{2.5}$ ($\mu\text{g m}^{-3}$)	3.2	3.5

This result suggests that the first two reaction pathways were not the major contributors to the observed nitrite formation. Despite the fact that aromatic compounds and humic acids were not measured, their concentrations were not presumed to increase much during the dust storm because mineral dust particles are not their sources. Considering that the daytime enhancement of nitrite production was only observed in the dust storm event, the third process should also be excluded as the key contributor.

The fourth process of TiO_2 photocatalysis of NO_2 yielding HONO was recently demonstrated in the laboratory (Gustafsson et al., 2006; Ndour et al., 2008). The reaction mechanism can be simply described as nitrogen dioxides being reduced to nitrite ions by photo-produced electrons (the detailed mechanism was proposed in Ndour's study, 2008). In the present study, the titanium concentrations (in the form of TiO_2 , Chen et al., 2012) indeed increased significantly during the dust storm of 24–26 April (see Fig. 2), implying that the surface TiO_2 photocatalysis may be responsible for the daytime nitrite formation. To further support this hypothesis, the correlations of daytime nitrite with titanium and with nitrate were also examined. Noting that the titanium concentration in the coarse mode particles ($PM_{>2.5}$) was unavailable, their values were calculated via the formula $\text{Ti}_{\text{coarse}} = \text{Ti}_{\text{fine}} \times (\text{Al}_{\text{coarse}}/\text{Al}_{\text{fine}})$ under the assumption that aluminum and titanium have the same size distributions (Zhang et al., 2003). As tabulated in Table 3, for sample-I, the ratio of $\text{NO}_2^- (\text{coarse})/\text{NO}_2^- (\text{fine})$ was approximately 3.0, which was similar to that of titanium (approximately 2.8), but largely differed from that of nitrate (approximately 9.1). The result for sample-III was similar. These results further indicate the role of TiO_2 photocatalysis on the observed daytime nitrite formation.

Table 3. Concentrations of titanium and nitrite in both fine particle mode ($PM_{2.5}$) and coarse particle mode ($PM_{>2.5}$) in the two daytime samples during the dust storm of 24–26 April at Mt. Heng.

Sample	Fine particles ($\mu\text{g m}^{-3}$)			Coarse particles ($\mu\text{g m}^{-3}$)		
	NO_2^-	Ti	NO_3^-	NO_2^-	Ti*	NO_3^-
04:25 Daytime	0.68	0.58	2.0	2.02	1.61	18.2
04:26 Daytime	1.5	0.42	1.0	4.49	1.12	10.7

* Ti in coarse particles was calculated by $\text{Ti}_{\text{coarse}} = \text{Ti}_{\text{fine}} \times (\text{Al}_{\text{coarse}}/\text{Al}_{\text{fine}})$.

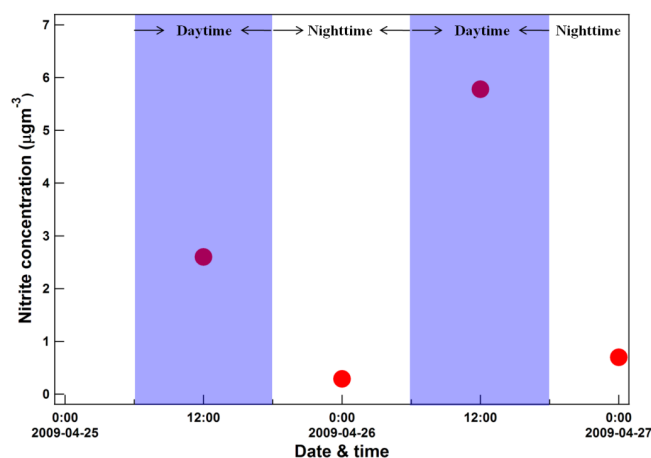


Fig. 8. Nitrite concentrations in TSP collected during the dust event on 24–26 April at Mt. Heng.

The partition between gas-phase HONO and aerosol nitrite is affected by acidity of dust particles. In the two daytime samples (sample-I and sample-III), the supermicron particles showed strong alkalinity (Fig. 5d). Under such a condition, the nitrous acid is normally thought to favor partitioning to the aerosol phase. However, Ndour et al. (2008) observed 2.3 ppbv of gas-phase HONO produced from 3 ppbv of NO_2 on irradiated real Sahara dust particles, suggesting the potential source of the gas-phase HONO from the heterogeneous photochemical reactions of dust particles even under the condition of strong alkalinity. In addition, the presence of coarse mode ammonium (samples I to III, the third stage, Fig. 7) may suggest the surface acidification of the dust particles, which would promote the partitioning of nitrous acid to the gas phase in samples I–III. For sample IV, strong acidity was

indicated (the fourth-stage, Fig. 7); the newly formed daytime nitrous acid would mainly partition to the gas phase as HONO and therefore would significantly contribute to the radical pool during daytime.

In summary, photo-enhanced nitrite formation on dust particles was observed, which indicated lab-demonstrated TiO₂ photocatalysis of NO₂ as an important source of daytime HONO under dust storm conditions. Given that there is increasing evidence for anthropogenic sources of photocatalyst-containing (including TiO₂) nanodusts emitted from industrial nanomaterials (Vicki, 2009; Chen et al., 2012), this process may play a potentially important role in atmospheric chemistry on a global scale. We therefore recommend more comprehensive studies to better evaluate the contribution of photoactive particles to gas-phase HONO and, in turn, OH radicals.

4 Summary and conclusions

An intense Asian dust storm that originated from the Gobi Desert was captured at a mountain site (Mount Heng) in southern China. The mineral dust particles experienced significant modifications before arriving at Mt. Heng, with large enrichments in secondary species, such as sulfate, nitrate and ammonium. The dust particles became more acidic and hydrophilic with the atmospheric processing time of air parcels. The accumulation of ammonium on dust particles was verified, which is related to the neutralization of carbonate. Based on the observations, a four-stage conceptual model is proposed to describe the aging process of the carbonate-rich Asian dust.

Photo-enhanced nitrite formation on the dust particles was also observed, indicating that the TiO₂ photocatalysis of NO₂ is a potentially important source of HONO (and nitrite) under dust storm conditions. In view of the potential for nanodust to become a global source of photocatalysts (e.g., TiO₂), this process may be an important daytime HONO source. More studies are needed in order to examine further the extent of photo-enhanced nitrite formation in the atmosphere and its implications for atmospheric chemistry.

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