

Implication of facies relationships of Upper Carboniferous/Lower Permian sediments in the Southern Alps (Carnic Alps/Karavanke Mts.) for Late Paleozoic paleogeography and Neogene tectonics

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Paleogeographic relationships of the Late Paleozoic successions in the Carnic Alps and Karavanke Mts. were for a long time obscured by the problems of lithologic characterization, terminology, and correlation of the sedimentary successions along the Austrian/Italian/Slovenian border triangle.

Joint investigation in both areas led to a revised lithostratigraphic scheme, a sedimentologic reinterpretation, and an updated biostratigraphic correlation of Upper Carboniferous/Lower Permian deposits. Facies relationships reveal an inner shelf, lagoonal setting (Goggau Limestone) during the late Early Permian present in the area from Tarvis, Kranjska Gora to Jesenice. Contemporaneous outershelf/slope environments (Trogkofel Limestone) are found about 30-40 km towards WNW in the Trogkofel/Gartnerkofel area of the Carnic Alps. We believe that this present-day geographical setting of the two facies belts is a result of dextral slip along the Schwarzwipfel/Fella/Sava Fault (SFSF). The 30-40 km of separation matches with independent slip estimates obtained from the displacement of exhumed Oligocene tectono-metamorphic units (Eder Unit, Mauthner Klamm Unit) and several other lithostratigraphic units (Devonian Feldkogel Limestone, Oligocene vulcanogenic complexes in Slovenia). The dextral displacement along the SFSF is regarded to be related to continuous northward drift and counterclockwise rotation of the Adriatic microplate in late Neogene (VRABEC & FODOR 2006).

According to other authors, Upper Carboniferous/Lower Permian sediments in the Southern Alps have been deposited in several narrow, fault-bounded pull-apart basins. Our considerations, however, suggest a wider, contiguous, (south?)eastward opening basin, which has been strongly modified and disintegrated during the Alpine orogeny.

VRABEC, M. & FODOR, L. (2006): Late Cenozoic tectonics of Slovenia: structural styles at the Northeastern corner of the Adriatic microplate. - In: PINTER, N., GRENERCZY, G., WEBER, J., STEIN, S., MEDAK, D. (eds.): The Adria microplate: GPS geodesy, tectonics and hazards, NATO Science Series, IV, Earth and Environmental Sciences, vol. 61, 151-168, Dordrecht, Springer.

Die post-variszische Schichtfolge der Karnischen Alpen (Österreich/Italien) Teil 1: Oberkarbon-Unterperm

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In den kürzlich erschienenen Erläuterungen zur Geologischen Karte des Jungpaläozoikums der Karnischen Alpen (SCHÖNLAUB & FORKE 2007) wurde aufgrund neuer Daten zur Biostratigraphie, Sedimentologie und Strukturgeologie eine überarbeitete lithostratigraphische Gliederung der post-variszischen Sedimentabfolge (Oberkarbon bis Trias) vorgestellt. Im Rahmen dieser Arbeit

wurden einige Schichtglieder revidiert, bzw. neu definiert. Die post-variszische Sedimentation beginnt im Oberkarbon (unteres Kasimovium) lokal mit verschiedenen Lydit- und Kalkbrekzien-/konglomeraten (Collendiaul Formation, Malinier Formation, Auernigalm-Kalkbrekzie), die mit deutlicher Winkel diskordanz auf variszisch verfaultem Altpaläozoikum liegen. Darüber folgt die vorwiegend marine ausgebildete Auernig Formation mit einer Wechselfolge aus Quarzkonglomeraten, Sandsteinen, Schiefern, und Kalken, die in verschiedene Subformationen unterteilt wurde (Meledis, Pizzul, Trögl, Watschig, Krone, Gugga, Garnitzen). Aufgrund der Unsicherheiten bei der lithologischen Korrelation einzelner Profilabschnitte in den oberkarbonen Ablagerungen und fehlender Definitionen der Typusprofile, erscheint uns, im Gegensatz zu früheren Autoren, die hierarchische Zuordnung der Auernig Formation in den Rang einer „Gruppe“ derzeit als nicht gerechtfertigt.

Die Auernig Formation wird von der überwiegend kalkig ausgebildeten Schulterkofel Formation (ehem. „Unterer Pseudoschwarzerinenkalk“) konkordant überlagert. Lithologisch kennzeichnend sind die eingeschalteten, massigen Riffkörper (Algenmounds), die häufig von dunklen, chert-reichen Horizonten überlagert werden sowie dem Auftreten von groß-fusiformen Foraminiferen in den gebankten Kalken. An der Wende zum Unterperm findet ein erneuter Wechsel zu siliziklastisch dominierten Sedimenten (Grenzland Formation) statt. Die dazwischen eingeschalteten Kalkbänke sind durch Onkoide und groß-kugelige Foraminiferen charakterisiert. Im oberen Teil der Grenzland Formation geht der Anteil an Siliziklastika stark zurück und es bilden sich mächtige, gebankte Kalke der Zweikofel Formation (ehem. „Oberer Pseudoschwarzerinenkalk“). Sie finden ihren Abschluß in den massigen Riffkalken des Trogkofels. Nach längerer Sedimentationsunterbrechung im oberen Unter/Mittelperm setzt ein neuer Sedimentationszyklus ein, beginnend mit lokalen, basalen Brekzienhorizonten (Tarfiser Brekzie/Trogkofel Konglomerat), gefolgt von der Grödner Formation und Bellerophon Formation.

SCHÖNLAUB, H.P. & FORKE, H.C. (2007): Die post-variszische Schichtfolge der Karnischen Alpen - Erläuterungen zur Geologischen Karte des Jungpaläozoikums der Karnischen Alpen 1:12.500. - Abh. Geol. B.-A., 61: 3-157, Wien.

Early Cretaceous and Jurassic events in the Western Carpathians

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In the W-Carpathians high pressure lithologies of Jurassic age from the Meliatic domain overlie the Palaeozoic Gemic units (counterpart of Grauwackenzone, Eastern Alps). Therefore it was often postulated that the Gemic was involved in a Jurassic closure history and early nappe stacking process. A similar evolution history was proposed to explain the Jurassic tectonics within the Northern Calcareous Alps S of Salzburg.

New geochronological results from the Western Carpathians yield arguments that Early Cretaceous crustal shortening, decollement and early nappe stacking process in the Gemic realm is independent from the Jurassic accretion in the Meliatic realm, both driven by rather different palaeogeographic – tectonic settings. In the Gemic a large number of Ar/Ar data from fine fractions can be interpreted as formation ages bracketing the growth history of white micas associated with the dominating penetrative crenulation cleavage between 110-125 Ma. This coincides perfectly with the end of sedimentation prograding northwards in the Mesozoic cover sediments. The Cretaceous cleavage fan

dominates the Gemic units, especially the central part. Variscan ages of a first slaty cleavage are only partially preserved at the northern border zone and to a lesser extent in the South. Summarizing, a Jurassic metamorphic event is not justified by the available data.

To understand the evolution in the Meliatic, new information from localized amphibolite grade rocks which have been overprinted under HP-LT conditions (glaucophane bearing) are of importance. These slivers of various micaschist were earlier interpreted as Variscan relics, because the muscovite yielded very uniform Ar/Ar ages at 370 Ma. However, it turned out that these muscovites are not relics but porphyroblasts, grown in a matrix of basic tuffitic material and showing Jurassic Rb/Sr ages.

As a consequence the Meliatic Jurassic subduction started with (localized) prograde amphibolite-facies conditions and later reached HP-LT conditions. Furthermore, these peculiar rocks with identical excess Ar ages are now exposed in form of small slivers along a distance of 60 km along strike direction. Most probably they have been dismembered during a strike slip event after their HP-metamorphism and prior to their emplacement in the present tectonic architecture.

A key for understanding may be the Late Jurassic palaeogeographic situation. Almost at the same time when the Meliatic HP event was active, the crustal block of Tisia - formerly located roughly at the area occupied nowadays by the W-Carpathians - was dismembered from the European continent. The consequences of these movements will be discussed.

Weiters wurden representativ Sedimentproben in Harz gehärtet, Dünnschliffe hergestellt und mittels Durchlichtmikroskopie mineralogisch untersucht.

WINKELBAUER, T. (1992): Von Hüttenmeistern und Glasmachern, Aschenbrennern und Flussiedern. - Das Waldviertel, Jahrgang **41** (52.), Heft 3, 225-252, Horn.

MATURA, A. (2006): Böhmisches Masse. - In: WESSELY, G. (Ed.): Geologie der österreichischen Bundesländer - Niederösterreich., Geologische Bundesanstalt, 25-39, Wien.

Rates and mechanisms of Miocene to Pleistocene exhumation in the Central Himalayas

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The interplay between erosionaly controlled and tectonically driven exhumation is a key factor that determines the morphology of mountain ranges and there is much debate how these processes interact and how exhumation velocities vary with time. We use information from five isotopic systems (⁴⁰Ar/³⁹Ar white mica, zircon fission track, (U-Th)/He zircon, apatite fission track, (U-Th)/He apatite) to constrain mode and velocity variation of central Himalayan exhumation between Miocene to Pleistocene times. Cooling ages are transferred to exhumation velocities by use of a new technique that relates the spatial distribution of geochronological data with isotherm geometries, displacement paths and topography. Data elaborated across Tethys Himalayas (north), Higher Himalayan Metamorphics, Lesser Himalayan Metamorphics and Lesser Himalayan Sediments (south) display non-steady state exhumation. Tectonically driven Miocene extrusion of the Himalayan metamorphic sequences at velocities of ~ 1-3 mm/yr and a climax around 13 Ma was followed by a period of tectonic calmness. Around 2.5 Ma ago, the high mountain relief was established by accelerated heterogeneous exhumation. Coevally, sediment delivery rates to the southern foreland increased. Within the Higher Himalayan metamorphic wedge exhumation velocities decreased southwards from >5 mm/yr, immediately south to the reactivated South Tibetan Detachment Zone to ~ 3 mm/yr at the Main Central Thrust Zone. At ~ 0.5 Ma tectonic exhumation ceased and erosion rates (~ 2.5-4 mm/yr) exceeded tectonic exhumation.

Grain boundary systematic by use of fractal geometry

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The shape of mineral grains and grain boundaries during deformation is a result of different operating deformation mechanisms which in turn reflect physical and chemical conditions during