AN UPPER TRIASSIC LIMESTONE PEBBLE WITH "SPIRIGERA" DESLONGCHAMPSI SUESS FROM THE SOUTHERN PART OF THE WESTERN CARPATHIANS, SOUTHERN SLOVAKIA

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Abstract: The brachiopod "Spirigera" deslongchampsi Suess has so far been known only in the Northern Calcareous Alps. It is the first find of this fossil in the Western Carpathians. The brachiopod belongs to the family Spirigerellidae Grunt 1965 (sensu Dagys 1974) and it is most probably a new genus. It is a very rarely occurring fossil (only a very few speciments were found altogether) so its generic identification for the time being remains doubtful. According to macro and mainly micro fossils the pebble material is of the Early Rhaetian age. We assume that it derives from a transitional type between Hallstatt limestones and Zlambach Beds facies.

Key words: Western Carpathians, Late Triassic, microfacies, microfossils, "Alpine" brachiopods.

Introduction

A limestone pebble with "Spirigera" deslongchampsi has been found in conglomerates of Egerian age on the northern margin of the Rimavská kotlina Depression. It is the first find of this species on the Western Carpathians territory and the first occurrence outside the Northern Calcareous Alps. The locality is situated 2 km SW from Chvalová village in the Banská Valley, see Fig. 1. The first petrographical analysis of these conglomerates was carried out by Marková 1959. The conglomerates are formed by pebbles of Mesozoic rocks (Triassic and Jurassic) of the Silica Nappe and partly of rocks of the Meliata Formation as reported by Mišík & Sýkora 1981. They identified Lower Triassic limestones with Meandrospira iulia Premoli-Silva, limestones of Wetterstein facies, "Tisovec" Limestone with Clypeina besici Radoicic, Hallstatt Limestone and limestone pebble with the brachiopods. Results of the brachiopod study are given in the present contribution. All these rocks mentioned belong to Silica Nappe succession. Metamorphic limestone and radiolarites of the Meliata Formation were rarely found.

Description of investigated pebble

The pebble studied is well rounded with (a) axis of 15 cm. The limestone is fine-grained and of rusty-brown to yellowish colour and nondescriptly spotted. The limestone is biomicrite — wackestone to packstone (Pl. II: Fig. 9). The biodetritus is mainly formed by calcified silicisponge spicules. Ossicles of echinoderms — crinoids, echinoid spines, ossicles of ophiuria, planctonic crinoids of the genus *Osteocrinus*, holoturian sclerites — *Theelia* sp. are present. Fragments of punctate brachiopods, foraminifers and ostracods occur relatively often. Sporadically zoospores — *Globochaete alpina* Lombard,

G. tatrica Radwanski and G. gregaria Schäfer & Senowbari-Daryan are found. Fragments of dasycladacean thalli (Pl. II: Fig. 3), Halicoryne (Pl. II: Fig. 4), Thaumatoporella parvovesiculifera (Raineri) as well as fragments of juvenile ammonite shells, minutegastropods and lamellibranch fragments are rare. Fragments of bryozoan zooecia of the order Cyclostomata are scattered, mainly belonging to a form of genus Stomatopora. Aeolisaccus tintinniformis Mišík, A. cf. amplimuralis Pantic, A. cf. inconstans Radoicic, Didemnoides moreti (Durand Delga) (Pl. II: Fig. 5) are rare. The organic detritus is prevailingly fragmentary and mainly of the size of fine-grained sand. Besides organic remains, a terrigenous admixture (clay and clastic quartz of silt size) is present. It contains several shells of "Spirigera" only but one specimen is relatively well preserved. We emphasize that this facies has not been found in the outcrops of the Silica Nappe yet.

Description of the brachiopod specimen

Order: **Athyridida** Boucot, Johnson & Staton 1964 Superfamily: **Athyridacea** Davidson 1881 Family: **Spirigerellidae** Grunt 1965

"Spirigera" d'Orbigny 1847

"Spirigera" deslongchampsi Suess 1855 (Pl. I: Figs. 1-2, Text. Fig. 2)

1855 Spirigera Deslongchampsi Suess — Suess, p. 26, Pl.1, Fig. 3. 1890 Spirigera Deslongchampsi Suess — Bittner, p. 243, Pl. 15, Figs. 3–5. 1988 "Spirigera" deslongchampsi Suess — Siblík, p. 83, Pl. 6, Fig. 5 (with note on using the name "Spirigera").

Holotype: It is deposited in the collection of the Geologische Bundesanstalt in Vienna under No. 1855/5/2 (Pl. I:

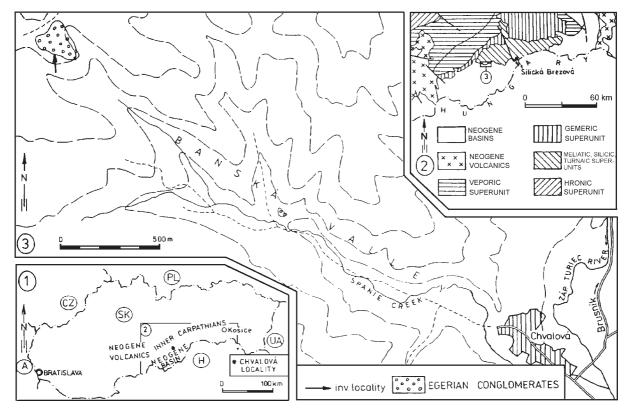


Fig. 1. Location of finding place of the examined pebble.

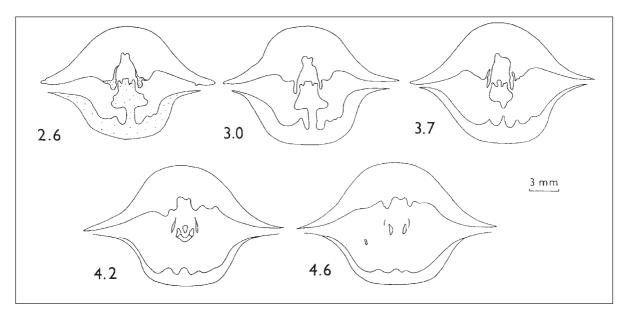


Fig. 2. "Spirigera" deslongchampsi Suess. Five transverse sections through posterior part of shell. Chvalová. Magnified.

Fig. 2). It measures ca. $23.0 \times 31.0 \times 14.8$ mm and derives from the Norian "Hallstätterkalk" of the Steinbergkogel near Hallstatt — the Hallstatt Zone, Juvavic Superunit, Upper Austria.

Material: One damaged specimen without posterior part of pedicle valve, measuring ca. $24.0 \times$ ca. 34.0×16.2 mm (Pl. I: Fig. 1) and one incomplete pedicle valve, deposited in the collection of Slovak National Museum under No. SNM Z 21978.

Diagnosis and remarks: Equibiconvex smooth shells of subpentagonal outline. Shallow sulcations developed both in

pedicle valve and on corresponding fold of brachial valve, perceptible as a flattening near umbones. Low uniplication relatively broad, rising steeply from commissural plane, and straightly limited on dorsal side. Due to the bad preservation of the only available specimen, no more than several sections through the interior structure could be gained. They showed quite different characters of the cardinalia (Fig. 2) in comparison with those known in other large, smooth athyridid genera as *Oxycolpella* Dagys, *Majkopella* Moisseiev or *Ochotathy*-

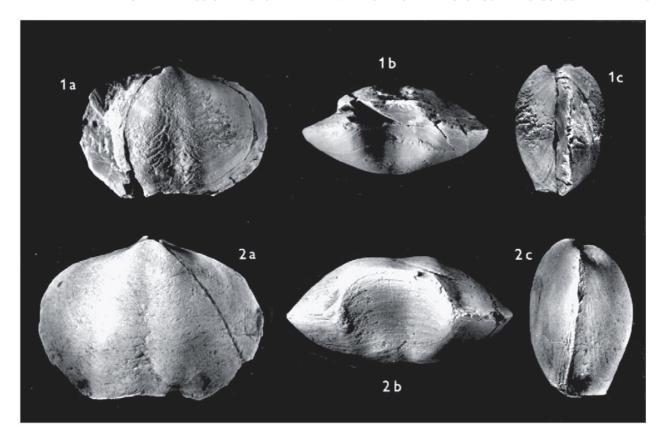


Plate I: Fig. 1 — "Spirigera" deslongchampsi Suess, Chvalová, ×1.5. Slovak National Museum No. SNM Z 21978. Fig. 2 — "Spirigera" deslongchampsi Suess. Steinbergkogel near Hallstatt. Holotype. Geologische Bundesanstalt No. 1855/5/2, ×2.

ris Dagys. In this respect, former Dagys' presumption of the possible appurtenance of "deslongchampsi" to Oxycolpella (Dagys 1965, p. 132) seems incorrect. Our material is characterized by absence of the pedicle collar, by a tripartite, very strong cardinal process and by poorly developed dental lamellae. In these characters it differs substantially from Oxycolpella. However the appurtenance to Spirigerellidae Grunt 1965 (sensu Dagys 1974) seems legitimate. The generic identification of "Spirigera" (= "Athyris") deslongchampsi regrettably remains doubtful (most probably new genus) until further, more favourably preserved finds. For the time being, the original generic name for "deslongchampsi" is used even if invalid.

The species has been only very rarely reported from the Northern Calcareous Alps. Our specimens agree well to the original material that we had at our disposal for comparison. Apart from the holotype (the only specimen known to Suess 1855) the collection of the Geologische Bundesanstalt in Vienna exhibits three specimens figured by Bittner (1890) on Pl. 15 (Fig. 3 with the dimensions $22.2 \times 30.5 \times 14.2$ mm; and young specimens on Fig. 4: $10.1 \times 9.9 \times 4.8$ mm and on Fig. 5: $10.3 \times 12.2 \approx 5.2$ mm). They all derive from Steinbergkogel near Hallstatt. New collecting done at this locality by L. Krystyn yielded a well-preserved pedicle valve (measuring 20.5×29.0 mm) that we had at our disposal, too.

Occurrence: Chvalová, Southern Slovakia. According to Bittner (1890) the most specimens came from Steinbergkogel near Hallstatt (marly Hallstatt Limestone of Late Norian age, and light-coloured crinoid limestone of Vandaites stu-

erzenbaumi Subzone = Early Rhaetian sensu Wiedmann & Krystyn in Wiedmann et al. 1979, p. 145). Siriuskogel near Ischl; other occurrences mentioned by Bittner (1890) are uncertain (Nassköhr near Neuberg, Teltschen near Bad Aussee and Barmsteine near Hallein).

Microfossils and their stratigraphic evaluation

The stratigraphic position of the analysed limestone pebble was established by determination of macro and micro fauna. "Spirigera" deslongchampsi has rarely been reported from Hallstatt Limestone of Late Norian age and from light coloured crinoid limestones of Early Rhaetian age, see part Description of brachiopod specimen. A part of pebble (about 25 dkg) was dissolved in 10% acetic acid. In insoluble residue, the following statigraphically important microfossils were found: conodonts, sclerites of holothurians, spines of ophiuroids, pedicellarian valves of echinoids and siliceous spicules of sponges.

Determinated forms:

a/ Conodonts — Misikella hernsteini (Mostler), (Pl. IV: Fig. 1); M. posthernsteini Kozur & Mock (Pl. IV: Fig. 2, 3); Oncodella paucidentata (Mostler), (Pl. IV: Fig. 4). M. hernsteini is predominant. The first determination of the conodonts in our sample was already made by Mock (1980, p. 133), and it was confirmed by our present study.

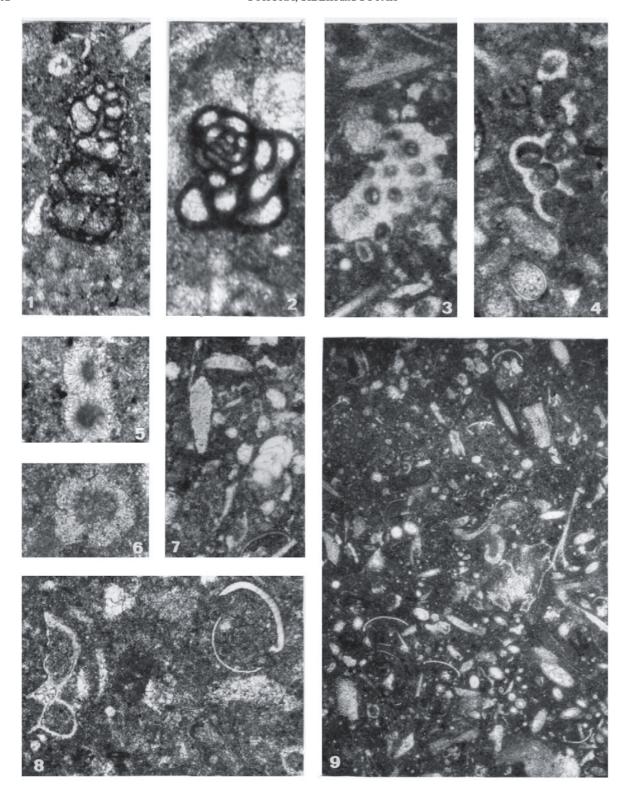


Plate II: Fig. 1 — Tetrataxis nanus Kristan-Tollmann, thin sec. 11211, ×86. Fig. 2 — Meandrospiranella? aff. planispira Oravecz-Scheffer, thin sec. 20834, ×133. Fig. 3 — Fragment of dasycladacean alga, thin sec. 20834, ×56. Fig. 4 — Fragment of Halicorynean thalli, thin sec. 20833, ×56. Fig. 5 — Muranella parvissima (Dragastan), thin sec. 11211, ×190. Fig. 6 — Didemnoides moreti (Durand Delga), thin sec. 11211, ×190. Fig. 7 — Osteocrinus sp., thin sec.11211, ×50. Fig. 8 — Bryozoan fragment and valves of ostracods, thin sec. 11211, ×100. Fig. 9 — Fossils assemblage in the analysed limestone, thin. sec. 11211, ×27.

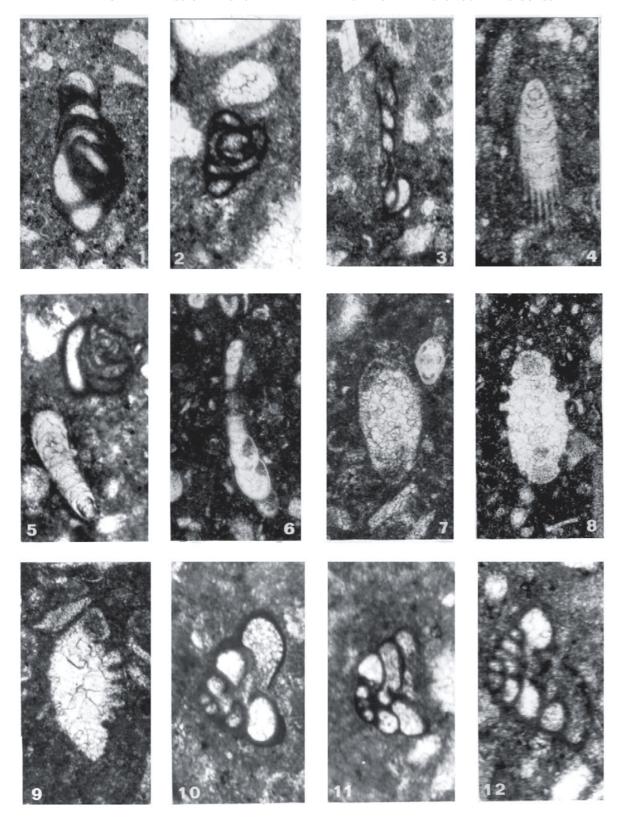


Plate III: Fig. 1 — Agathamminoides spiroloculiformis (Oravecz-Scheffer), thin sec. 11211, ×86. Fig. 2 — Agathammina incostans (Michalík, Jendrejaková & Borza), thin sec. 11292, ×129. Fig. 3 — Planiinvoluta carinata Leischner, thin sec. 20833, ×86. Fig. 4 — Austrocolomia canaliculata Oberhauser, thin sec. 20834, ×127. Fig. 5 — Frondicularia woodwardi Howchin and Quinqueloculina sp., thin sec. 20834, ×86. Fig. 6 — Rectoglandulina aff. polyarthra Kristan-Tollmann, thin sec. 20833, ×45. Fig. 7 — Aulotortus communis (Kristan), thin sec. 20834, ×86. Fig. 8 — Involutina cf. turgida Kristan, thin sec. 11211, ×45. Fig. 9 — Trocholina granosa (Frenzen), thin sec. 11211, ×110. Figs. 10, 11 — Tetrataxis inflata Kristan, thin sec. 20833 ×117, 20834, ×86. Fig. 12 — Tetrataxis nanus Kristan-Tollmann, thin sec. 11296, ×119.

b/ Sclerites of holothurians — *Theelia stellifera* Zankl (Pl. VI: Fig. 1); *Theelia simoni* Kozur & Mock (Pl. VI: Fig. 2); *Theelia corbula* Zankl (Pl. VI: Fig. 3); *Theelia variabilis* Zankl (Pl. VI: Fig. 4); *Theelia sp.* (Pl. VI: Fig. 5); *Punctatites sp.* (Pl. VI: Fig. 6).

c/ Ophiuroid spines — "Loch" types sensu Mostler (1972, Pl. 3) — (Pl. V: Figs. 7, 8). These forms were reported by Mostler l.c. from Late Norian and Rhaetian age. The other types (Pl. V: Figs. 6, 9, 10) are similar to those found and figured by Mostler (1971) from the Upper Triassic limestones.

d/ Pedicellarian valves — ☐ tridentate (Pl. IV: Figs. 5, 6); ophicephalous (Pl. IV: Figs. 7, 8), similar types of valves were described by Mostler (1972, Pl. 1, 2) from limestones of Norian age.

e/ Sponge spicules — tetraxons, orthodichotriaene forms (Pl. IV: Fig. 10; Pl. V: Fig. 1); tetraxons ?triaene form with reduction of rhabds (Pl. V: Fig. 3); desmas, rhabdoclone types (Pl. IV: Fig. 11; Pl. V: Fig. 2); diactinal monoaxon, amphitilote type (Pl. V: Fig. 5) and criccostyl (Pl. V: Fig. 4). Forms of the sponge spicules on Pl. IV: Figs. 10, 11; Pl. V: Figs. 1, 2, 3, 4 are similar to those figured by Mostler (1976) from limestones of Sevatian age.

In the thin sections we identified stratigraphically important foraminifers. Their association is characterized by prevailing ophthalmidia and sessile forms. These groups of forams are represented by the following species: *Ophthalmidium lucidum* (Trifonova), *Oph. triadicum* (Kristan), *Oph. fusiformis* Trifonova, *Paleonubecularia? floriformis* Ciarapica & Zaninetti, *Agathammina incostans* (Michalík, Jendrejáková & Borza) (Pl. III: Fig. 2), *A. austroalpina* Kristan-Tollmann & Tollmann, *Agathamminoides spiroloculiformis* (Oravecz-Scheffer) (Pl. III: Fig. 1), *Planiinvoluta multitabulata* (Kristan-Tollmann), *P. carinata* Leischner (Pl. III: Fig. 3), and *P. regularis* Salaj, Borza & Samuel.

In the associations also Milioporidae represented by the species *Galeanella? tollmanni* (Kristan) are found sporadically. Nodosarian foraminifers like *Austrocolomia canaliculata* Oberhauser (Pl. III: Fig. 4), *Dentalina hoi* Trifonova, *Pseudonodosaria* sp., *Frondicularia xiphoidea* Kristan-Tollmann, *F. woodwardi* Howchin (Pl. III: Fig. 5) and *Rectoglandulina* aff. *polyarthra* Kristan-Tollmann (Pl. III: Fig. 6) are among the common species.

Involutinid and glomospiroid foraminifers occur less frequently and have subtle growth. They are represented by Aulotortus communis (Kristan) (Pl. III: Fig. 7) A. tumidus (Kristan-Tollmann), Gandinella falsofriedli (Salaj, Borza & Samuel), Coronipora austriaca (Kristan), Involutina cf. turgida Kristan (Pl. III: Fig. 8) and Trocholina granosa (Frentzen) (Pl. III: Fig. 9). From further groups the species Trochammina januensis Brönnimann & Page, Tetrataxis inflata Kristan (Pl. III: Figs. 10, 11), T. nanus Kristan-Tollmann (Pl. III: Fig. 12), Meandrospiranella? aff. planispira Oravecz-Scheffer (Pl. II: Fig. 2), Turrispirillina minima Pantic, Oberhauserella ovata Fuchs and Earlandia amplimuralis (Pantic) were indentified. The described assemblage of foraminifers with numerous forms of ophthalmidia, nodosaria (mainly Austrocolomia canaliculata), points to the Norian age s.l. (i.e. including the Early Rhaetian). From the point of view of the indistinct occurrence of involutid forams which does not correspond to their explosive extension in the Norian s.s., the age of the limestone studied would rather correspod to the Early Rhaetian. This range is also supported by the sporadic presence of the Early Rhaetian index species *Gandinella falsofriedli*, Early Rhaetian Angulodiscus friedli Zone according to Salaj et al. (1983) and the appearance of Early Rhaetian *Milioporidae* (species *Galeanella? tollmanni*). In the associations of foraminifers the first involutinas (*Involutina* cf. *turgida*) already appear. They are mentioned by Salaj et al. l.c. from the Rhaetian Triassina hantkeni Zone. According to all the mentioned microfossils the pebble is of Early Rhaetian age. According to conodonts it belongs to Misikella hersteini-Misikella posthernsteini Subzone sensu Kozur & Mock (1991).

Comparison of the analysed limestone with Upper Triassic sediments of the Silica Nappe near Silická Brezová village

The microfacies of the limestone with "Spirigera" was compared with the microfacies of Upper Triassic limestones and marls near Silická Brezová village (situated approximately 26 km E of Chvalová village, see Fig. l) at two localities:

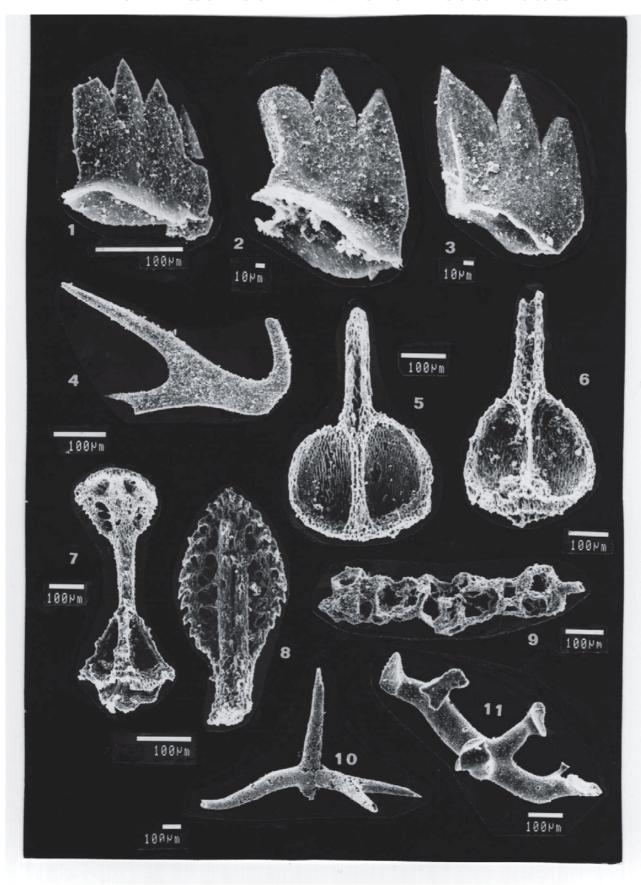
A. New section SW of old quarries — red and pink Hall-statt Limestone in its uppermost part probably corresponds according to unpublished stratigraphic data to the Early Rhaetian

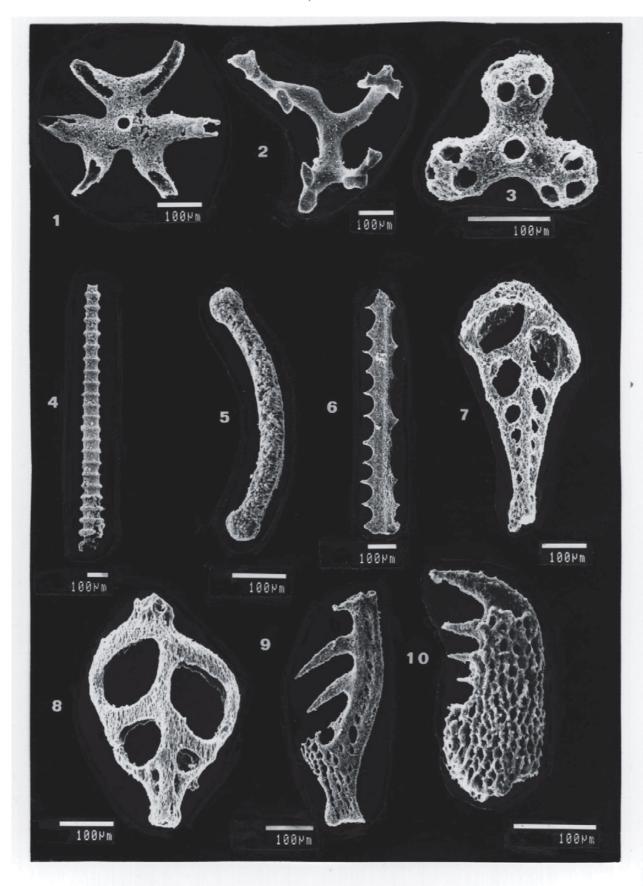
B. Malý Mlynský vrch Hill — locality situated l km east of Silická Brezová village where the Zlambach Beds of Late Norian to Rhaetian age occur (their age according to Mock 1973, 1980 and Kozur & Mock 1974a,b).

In another place — Bohúňovo village — the sedimentation of Zlambach Beds facies continued into the Early Jurassic (Mello 1997), as in the Northern Calcareous Alps (see e.g. Zankl 1971).

The red Hallstatt Limestone (loc. A) is represented by wackestones to packstones. They contain a relatively abundant association of fossil remains: mainly shells of bivalves, juvenile ammonites, ostracods, sponge spicules, ossicles of echinoderms, foraminifers in various proportions. Allochems are not sorted. Terrigenous admixture is very low, about 2 grains of quartz per thin section (compare also Mišík & Borza 1976).

Plate IV: Fig. 1 — Misikella hernsteini (Mostler), neg. 1454, × 370. Fig. 2 — Misikella posthernsteini Kozur & Mock, neg.1421, × 450. Fig. 3 — Misikella posthernsteini Kozur & Mock, neg.1419, ×450. Fig. 4 — Oncodella paucidentata (Mostler), neg, 1456, × 220. Fig. 5 — Tridentate pedicellarian valve of a sea urchin without a tip part, neg. 1448, ×150. Fig. 6 — Tridentate pedicellarian valve of a sea urchin without a tip part, neg. 1428, ×190. Fig. 7 — Ophiocephalous pedicellarian valve of a sea urchin, neg. 1443, × 150. Fig. 8 — Dentated tip of the ophiocephalous pedicellarian valve, neg. 1444, ×230. Fig. 9 — Agglutinate foraminifera, neg. × 150. Fig. 10 — Sponge spicule, tetraxon - orthodichotriaene, neg. 1457, ×80. Fig. 11 — Sponge spicule, desmarhabdoclone, neg. 1463, ×150.





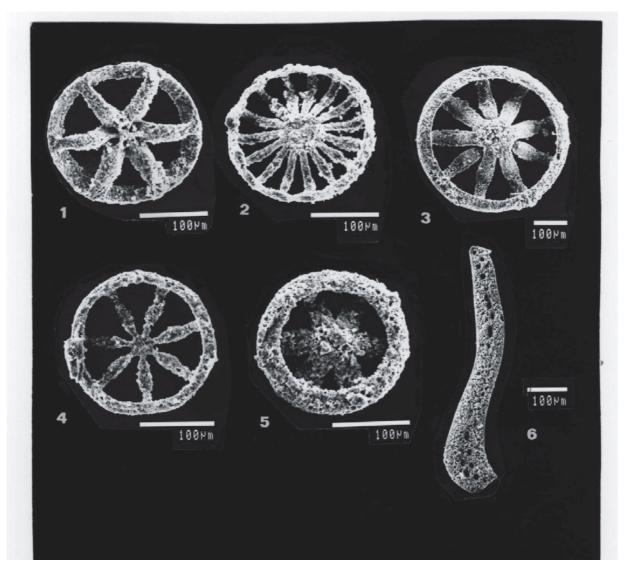


Plate VI: Fig. 1 — Theelia stellifera Zankl, neg. 1436, ×300. Fig. 2 — Theelia simoni Kozur & Mock, neg. 1426 ×300. Fig. 3 — Theelia corbula Zankl, neg. 1449, ×150. Fig. 4 — Theelia variabilis Zankl, neg. 1455, ×330. Fig. 5 — Theelia sp., neg. 1438, ×350. Fig. 6 — Punctatites sp., neg. 1434, ×180.

Grey to grey-brownish rocks of the Zlambach Beds (loc. B) are wackestones, they contain small fragmentary skeletal remains (silt to very fine sand size mainly). Thus the identification of allochems is often difficult. Fragments of bivalve shells, echinoderm ossicles, foraminifers and ostracods were distinguished. Peloids, small intraclasts and epigenetic pyrite are present, too.

Plate V: Fig. 1 — Sponge spicule, tetraxon - orthodichotriaene with reduction of rhabds, neg. 1425, ×190. Fig. 2 — Sponge spicule, desma - rhabdoclone, neg. 1418, ×150. Fig. 3 — Sponge spicule, tetraxon - ?triaene with reduction of rhabds, neg. 1452, ×350. Fig. 4 — Sponge spicule, criccostyl neg. 1431, ×90. Fig. 5 — Sponge spicule, diactinal monoaxon - aphitylote, neg. 1433, ×230. Fig. 6 — Thorny spine of ophiuroid, neg. 1427, ×120. Fig. 7 —"Loch" spine of ophiuroid, neg. 1422, ×190. Fig. 8 — "Loch" spine of ophiuroid, neg. 1447, ×200. Fig. 9 — Hooked spine of ophiuroid, neg. 1440, ×350. Fig. 10 — Hooked spine of ophiuroid, neg. 1462, ×350.

The rock also contains terrigenous admixture — clay and silt mainly (quartz: silt to 0.3 mm), it is often bioturbated and spotted. In the upper parts of the formation, beds of hybrid calcareous sandstones with a quartz content of 40-50 % sporadically appear. They differ from the Zlambach Beds of the Northern Calcareous Alps mainly in the smaller content and reduced diversity of their fossils, and in the absence of corals (compare Zankl 1971; Pistotnik 1972; Matzner 1986).

According to the association of fossil remains, the studied limestone with "Spirigera" is similar to Hallstatt Limestone near Silická Brezová.

It differs from it in the content of terrigenous component (admixture), colour and indistinct spottedness. It differs from rocks of the Zlambach Beds near Silická Brezová in the assemblage, preservation and size of organic remains, also in colour and the smaller content of terrigenous component (compare description of the pebble studied). We suppose that the analysed limestone pebble

represents a facies type, transitional between the facies of Hallstatt Limestone and Zlambach Beds.

Although papers, describing the lithology and microfacies of the Hallstatt Limestone (containing "Spirigeras") in detail are not familiar for us (localities Steinbergkogel near Hallstatt, Siriuskogel near Ischl, Nassköhr near Neuberg, see Suess 1855; Bittner 1890), we suppose that the Upper Triassic limestones in some localities of the Juvavic Superunit (Northern Cacareous Alps) and Silicic Superunit (Western Carpathians) sedimented under similar biofacial conditions.

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