



Late Ordovician Conodonts from the Olistostromal Wisemans Arm Formation (New England Region, Australia)

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2 Text-Figures, 2 Tables and 4 Plates



Australia
Ordovician
Conodonts
Olistoliths
Biochronology

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Oberordovizische Conodonten aus der olistostromen Wisemans Arm Formation (New England, Australien)

Zusammenfassung

Allochthone Kalke von sieben Lokalitäten der olistostromen Wisemans Arm Formation der New-England-Faltenkette haben eine Fauna von einigen 700 Conodonten erbracht, die 24 Arten zugeteilt werden. Die Vergesellschaftung gleicht anderen, die kürzlich vom Lachlan Fold Belt im zentralen New South Wales dokumentiert wurden, insofern als sie eine Mischung von Arten enthält, die für mehrere spätordovizische Faunenprovinzen charakteristisch ist. Die Fauna wird dominiert durch *Panderodus gracilis* und *Belodina confluens* und enthält *Taoqupognathus tumidus*, *Yaoxianognathus ani* und *Panderodus nodus*, die ein spätes Eastonium-Alter (Ea3) angeben. Die Wisemans-Arm-Fauna ist vergleichbar mit derjenigen, die von den Trelawney-Schichten des Tamworth Belt bekannt sind, indem sie eine Verbindung zwischen den zwei wesentlichen tektonischen Einheiten des New-England-Faltenzuges vermitteln. Die Conodontenfaunen und Kalke weisen auf die Präsenz – während des späten Ordoviziums – von ozeanischen Vulkaninseln mit assoziierten Kalkriffen in diesem Gebiet hin.

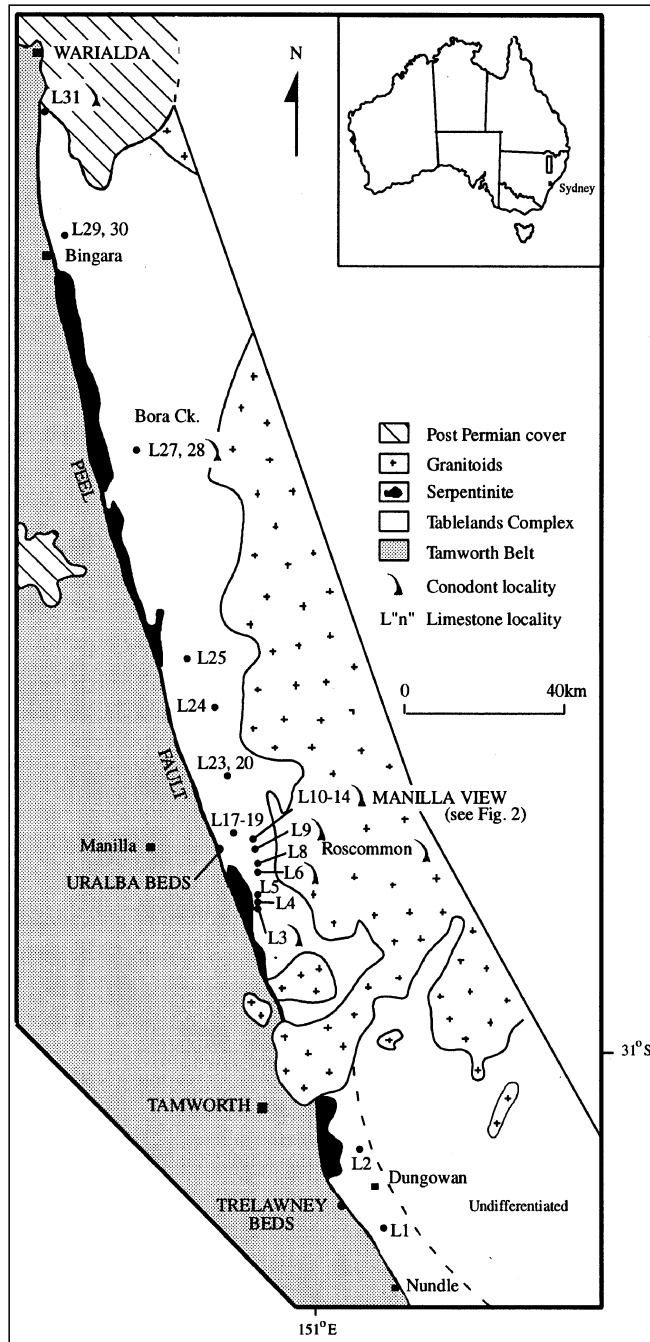
Abstract

Allochthonous limestone from seven localities in the olistostromal Wisemans Arm Formation of the New England Fold Belt has yielded a fauna of some 700 conodonts, assigned to 24 species. The assemblage is similar to others recently documented from the Lachlan Fold Belt in central New South Wales in comprising a mixture of species characteristic of several Late Ordovician faunal provinces. The fauna is dominated by *Panderodus gracilis* and *Belodina confluens* and includes *Taoqupognathus tumidus*, *Yaoxianognathus ani* and *Panderodus nodus*, indicating a Late Eastonian (Ea3) age. The Wisemans Arm fauna is comparable with that known from the Trelawney Beds of the Tamworth Belt, providing a link between the two major tectonic elements of the New England Fold Belt. The conodont fauna and limestone indicates the presence, during the Late Ordovician, of oceanic volcanic islands with associated limestone reefs in this region.

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1. Introduction

Late Ordovician rocks crop out intermittently between Nundle and Warialda, east of the Peel Fault in the New England Fold Belt, northeastern New South Wales (Text-Fig. 1). Although a number of palaeontological reports concerning rocks in this area have documented ages ranging from Ordovician through Early Carboniferous (eg. CHAPPELL, 1961; PICKETT, 1973, 1982; HALL, 1975, 1978; LEITCH & CAWOOD, 1980; ISHIGA, 1988; ISHIGA et al., 1988; AITCHISON, J.C., 1993), none was based on investigation of more than isolated or a few localities. The need for a better data base prompted investigation of limestone outcrops along 250 km of strike with the principal aims of improving precision and understanding of the ages of the carbonates and evaluating their tectonic implications for the New England Fold Belt.



Text-Fig. 1. Location of area investigated herein, major structural elements and representative limestone/conodont localities.

The presence of limestone outcrops in the tract just east of the Peel Fault has been reported sporadically for over a century (eg., DAVID, 1891; BENSON, 1913; CARNE & JONES, 1919; CHAPPELL, 1961; LISHMUND et al., 1986; LEITCH & CAWOOD, 1980). These, and a number of unpublished reports have led to the identification of more than 50 localities. Present work has involved the sampling of these (some, poorly specified localities, could not be located) and a number of additional outcrops have been found. Text-Fig. 2 indicates numerous carbonate occurrences east of Manilla, reflecting a number of detailed mapping efforts in that area (eg. CHAPPELL, 1961 and LEITCH & CAWOOD, 1980), and attests to the likelihood that small limestone bodies in the area of concern are relatively common.

Table 1 summarises locality details and outcrop data. The limestones examined herein are all considered to be allochthonous, often exposed as debris flows (Text-Fig. 2), and have a clast size ranging from sub-cm in conglomerates to olistoliths of several hundred meters. Limestones producing conodonts are characteristically recrystallised fine sands probably deposited in shallow settings, and contain rare small, rounded volcanic grains. The limestone blocks are hosted by a matrix of volcanoclastic sediments comprising a spectrum from conglomeratic to fine silt and mudstones. Results of investigation into sediment lithological and depositional characteristics will be published elsewhere.

2. Geological Setting

The southern part of the New England Fold Belt is divided by the Peel Fault System into a western magmatic arc and forearc basin succession (Tamworth Belt), and an eastern subduction-accretion complex (Tablelands Complex) (LEITCH, 1974), that is the focus of this investigation.

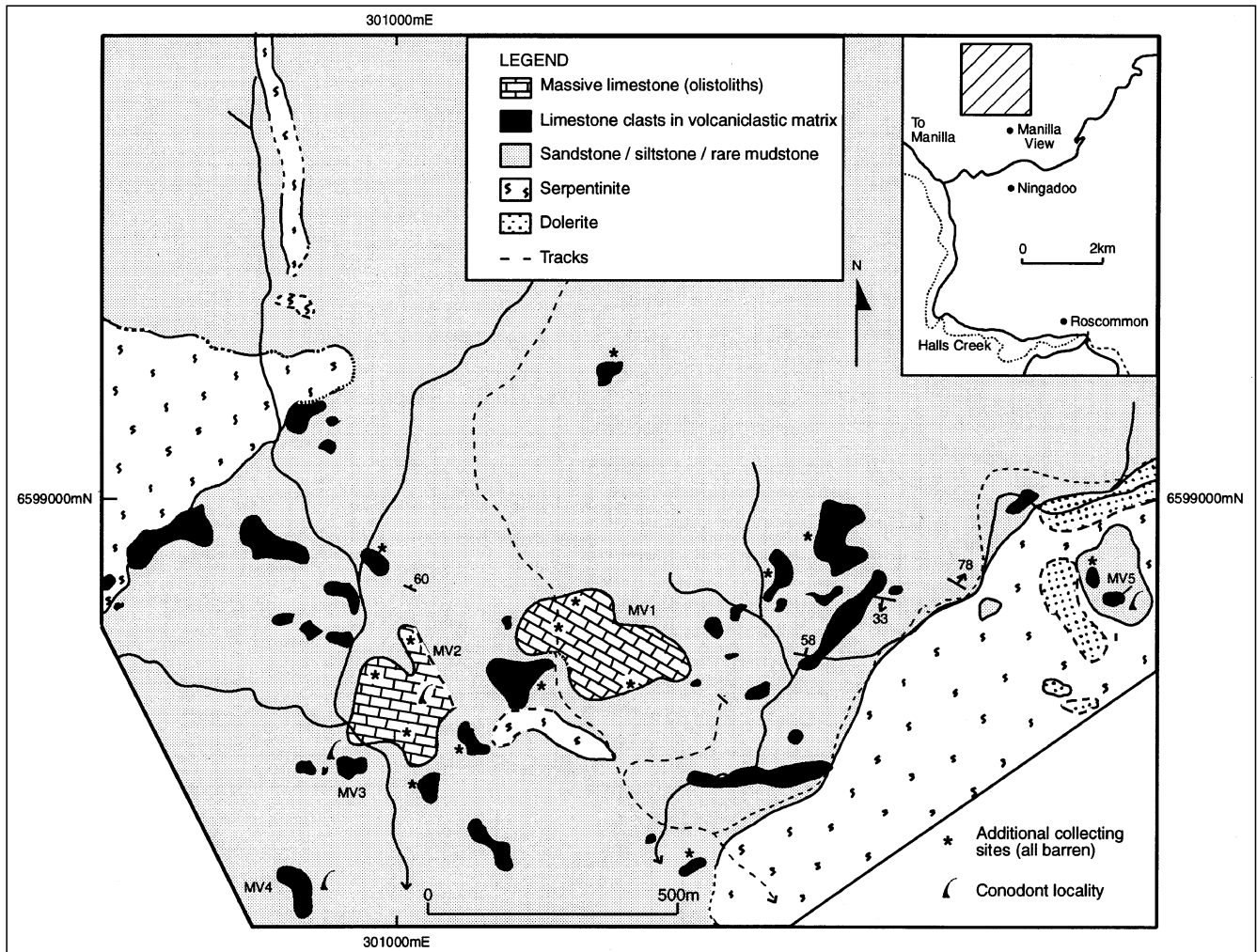
Stratigraphic nomenclature used by various authors in the intensely deformed Tablelands Complex immediately east of the Peel Fault (Text-Fig. 1) was summarised by CROOK (1961), he applied the name Woolomin Beds to the area. This has since been revised. CAWOOD (1980) and LEITCH & CAWOOD (1980) recognised two main units in the Attunga-Halls Creek area, the Woolomin and Wisemans Arm formations. The structurally lowest Woolomin Formation and its lateral correlatives extend for approximately 250 km from Warialda in the north to the Manning River headwaters, bounded by the Wisemans Arm Formation (WAF) and lateral correlatives to its west and the New England granitoids to the east, and is characterised by repeated slices of metabasalt, chert and siltstone with rare outcrops of sandstone and pebbly siltstones (CAWOOD, 1980). The WAF was defined by LEITCH & CAWOOD (1980) as a fault bounded unit cropping out between the Peel Fault in the west and the Mount Abundance Fault in the east. Lithological components include siltstone, tuff, sandstone and olistostromal conglomerates that contain olistoliths of chert and limestone in a volcanoclastic matrix of andesitic and basaltic provenance (CAWOOD, 1983). Present field work extends the WAF tract north to the Bungendore Spur area (FUREY-GREIG, in prep). CAWOOD (1980) mapped correlative rocks to the south between Nundle and Piallamore; AITCHISON et al. (1988) correlated the WAF with the Nangarah Formation in the Woodsreef-Bingara area to the north. Lithological similarities were considered by LEITCH & CAWOOD (1980) as evidence for provenance links between chert olistoliths in the WAF and

Table 1.

Summary of locality, outcrop and microfossil details.

cr = crinoid fragments; cnd = conodonts; crl = corals; oc = ostracoda; sc = scolecodonts; b = brachiopoda.

LOCALITY	GRID & 1:25,000 Sheet	FORMATION	DESCRIPTION	OUTCROP AREA	FOSSILS	AGE	COMMENTS
1 Wallaby Mtn.	2050.3500 Woolomin	Wisemans Arm	deformed clasts	c. 50x50m	barren	no data	scattered clasts on W side of ' Mtn.
2 Piallamore	1580.5115 Dungowan	Wisemans Arm	altered clasts	20x5m	barren	no data	loose boulders W of fence
3 Springwater	0110.8560 Klori	Wisemans Arm	olistoliths & clasts	35x10m	cnd.	indet.	c. 350m upslope from creek
4 Serpentine Hill	0065.8660 Klori	Wisemans Arm	sheared clasts	40x20m	barren	no data	c. 300 upslope from bend in road
5 Parish Boundary	0085.8735 Klori	Wisemans Arm	olistolith & clasts.	20x20m	cr	no data	adj. to boundary fence
6 Lone Boulder	0190.9140 Klori	Wisemans Arm	single megaclast	1x1m	cnd.	Eastonian	probably = Cawood'80&Pickett'73
7 Wisemans Arm Rd.	0180.9160 Klori	Wisemans Arm	clasts in congl.	5x5m	barren	no data	next to fence
8 Roscommon 1	0230.9280 Klori	Wisemans Arm	clasts, 1cm-2m	500x100m	cnd,sc,crl	Eastonian	extends N from new road cutting
9 Ningadoo South	0250.9585 Klori	Wisemans Arm	coarse sand & clasts	200x75m	cnd	Eastonian	on SW side of creek
10 Manilla View 1	1340.9880 Mundowey	Wisemans Arm	olistolith & clasts	300x100m	barren	no data	straddles intersection of 4wd tracks
11 Manilla View 2	1330.9870 Mundowey	Wisemans Arm	olistolith & clasts	250x100m	cd,cr,oc	indet	extends S from new dam
12 Manilla View 3	1200.8960 Mundowey	Wisemans Arm	clasts+ ?olistolith	25x10m	cd	Eastonian	? SW extension of MV2
13 Manilla View 4	0120.9860 Mundowey	Wisemans Arm	scattered clasts	75x30m	cnd	Eastonian	S & adjacent to fence
14 Manilla View 5	0150.9870 Mundowey	Wisemans Arm	clasts in skarn	20x20m	cnd	Eastonian?	blocks adjacent to dolerite
15 WG 116	1500.9900 Mundowey	Wisemans Arm	clasts in congl.	200x50m	barren	no data	c. 200m upslope from creek
16 Glen Bell 1	0230.9930 Mundowey	Wisemans Arm	clasts in congl.	5x5m	barren	barren	next to track, 25m E of gate
17 Ukolan 1	9830.9940 Mundowey	Wisemans Arm	megaclasts & clasts	10x25m	barren	no data	on lower part of slope, N of creek
18 Ukolan 2	9790.9950 Mundowey	Wisemans Arm	megaclasts & clasts	50x80m	barren	no data	adjacent & on E side of gully
19 Killarney	9710.9990 Mundowey	Wisemans Arm	megaclasts & clasts	60x10m	barren	no data	E of windmill;
20 Windfall Valley	9540.1230 Welsh	Wisemans Arm	olistolith & clasts	c.200x75m	barren	no data	W of watertank
21 Dunmore S-1	9430.0450 Mundowey	Wisemans Arm	clasts	10x5m	barren	no data	on W side of Glenbara Road
22 Dunmore S-2	9380.0650 Mundowey	Wisemans Arm	deformed clasts	15x50m	barren	no data	up steep slope, SW side of track
23 Dunmore N-1	9585.1195 Welsh	Wisemans Arm	megaclasts (c. 2m)	20x2m	barren	no data	2 small O/C's, adj. to fence on hill
24 Pera South	9325.2285 Welsh	?Wisemans Arm	olistolith & clasts	500x80m	cnd.	indet.	extends from S boundary of "PERA"
25 Rocky Creek	8565.3720 Linton	?Wisemans Arm	clasts in red matrix	10x3m	barren	no data	in W bank of Rocky Ck.
26 Gulf Cr. smelter	7680.5730Cobbadah	?Woolomin	stockpile, local l/s	N/A	barren	no data	quarry inaccessible
27 Bora 1	7070.8360 Dinoga	Nangarah	thermally altered	2x5m	barren	no data	S of & next to fence, 500m E of creek
28 Bora 2	7070.8350 Dinoga	Nangarah	thermally altered	2x5m	cnd	indet	c. 250m SSW from Bora 1
29 Lochiel 2	6760.9600 Bingara	Nangarah	blasted blocks	5x5m	barren	unknown	small quarry close to fence
30 Lochiel 1	6750.9650 Bingara	Nangarah	blasted blocks	5x5	barren	unknown	small quarry on 4wd track
31 Balfour	6370.1680 Mt. Rodd	?Nangarah	megaclasts & clasts	250x75m	cnd	Late Ord.	c. 1000m N from road
32 Kellys Gully	6365.2455 Warialda	Undifferentiated	altered clasts	300x50m	cnd	indet.	extends S from railway



Text-Fig. 2. Geological map of the Manilla View area (Localities 10–14), showing debris flow of limestone olistoliths and conglomerate in the Wisemans Arm Formation.

the Woolomin Formation, and between intermediate volcanic detritus and the Tamworth Belt arc volcanics.

Age determinations of limestones from the WAF vary from Ordovician to Early Devonian for limestones (CHAPPELL, 1961; HALL, 1975, 1978; PICKETT, 1982; LEITCH & CAWOOD, 1980); an Early Carboniferous date was advanced for allochthonous chert, based on radiolaria (ISHIGA et al., 1988). At least one of the two reported Early Devonian limestones (LEITCH & CAWOOD, 1980), Locality #6 herein, is in fact Eastonian. Samples from another locality proved barren, and the third was not located. The Early Devonian age reported should be regarded as unreliable as the diagnostic fossils quoted (never illustrated and now lost) have superficially identical forms in the late Ordovician (see Discussion). As recent research in the Tamworth Belt on Early Devonian conodonts (FUREY-CRAIG, 1995; MAWSON et al., 1995) has demonstrated the unreliability of biostratigraphic correlation using radiolaria to date New England sediments, Early Carboniferous dates in the WAF need to be substantiated by conodont data. Siliceous matrix sediments hosting the olistoliths are currently being processed for conodonts using the method described by ORCHARD (1987) and FOWLER & IWATA (1995).

Mapping by LEITCH & CAWOOD (1980) and present field work indicate that sub-units within the WAF that include either chert or limestone olistoliths appear to be mutually exclusive. Silurian age assignments based on conodonts,

from Woolomin Formation chert (ISHIGA et al., 1988) and from one limestone questionably within the WAF (PICKETT, 1982), with Silurian (Late Llandoveryan) corals reported from the same locality (HALL, 1978) [Locality 4 herein], might restrict the age of the WAF to no older than Early Silurian, but structural complexity in the area investigated raises the possibility that thin slices of early Palaeozoic strata are preserved within the WAF.

The current examination of siliceous rocks and associated detailed mapping should assist in further resolving the age of the WAF. Correlation between the WAF conodont fauna and conodonts from the Uralba Beds east of Manilla (FUREY-GREIG, in prep), and a fauna reported from the Trelawney Beds (PHILIP, 1966), south-east of Tamworth (see Discussion) has elucidated the palaeogeography of New England during the Late Ordovician.

3. Biostratigraphy

It is evident from data presented in this report that the Wisemans Arm conodont fauna is strikingly similar to material recently documented from Lachlan Fold Belt localities, namely the Malongulli Formation (TROTTER & WEBBY, 1994) and the Dowerry Member of the Balingoole Limestone, Bowan Park Group (ZHEN et al., 1998).

TROTTER & WEBBY (1994) used graphical methods to correlate the Malongulli fauna with the *Oulodus velicuspis* zone

of the North America Mid-continent system, with that determination underpinned by co-occurring graptolites of Late Eastonian (Ea3) age. The WAF fauna includes several species present in the Malongulli Formation, eg. *Belodina confluens* SWEET (12 % of the WAF material) and *Periodon grandis* (ETHINGTON) (2 %). More importantly, *Taoqupognathus tumidus* TROTTER & WEBBY has only previously been described from the abovementioned Lachlan Fold Belt occurrences and from the Carriers Well Formation (SIMPSON, 1997) and Fork Lagoons Beds (PALMIERI, 1978) in Queensland. *Panderodus nodus* ZHEN et al. and *Yaoxianognathus ani* ZHEN et al. are reported from the same units, as well as from Late Ordovician Chinese faunas (ZHEN et al., 1998). All are here described from the WAF fauna, indicating a Late Eastonian (Ea3) age for these limestones.

The WAF fauna, including species previously considered as representative of several "Faunal Provinces" as recognised by BERGSTRÖM (1990), adds to a growing list of similar eastern Australian faunas. NOWLAN et al. (1997) assessed such faunas and established a new "Australasian Province", within the Midcontinent Faunal Region (their Fig. 2).

3.1. Correlation with Other Eastern Australian Occurrences

PHILIP (1966) documented conodonts from coralline limestone of the "Trelawney Beds" southeast of Tamworth, including *B. confluens*, *Phragmodus undatus* and *Panderodus gracilis*, a fauna similar to that from the WAF. The Trelawney Beds, represented by loose blocks scattered in a field were interpreted by PHILIP as derived from a fault bounded remnant of a Late Ordovician unit adjacent to the Peel Fault System. CAWOOD (1980) remapped the area and considered that the limestones were Ordovician olistoliths within the Drik Drik Formation of the Tamworth Belt, known to be of Early Devonian (Emsian) age (FUREY-GREIG, 1994, 1995). The presence of such clasts elsewhere in the Drik Drik Formation has been demonstrated by CAWOOD (1980).

HALL (1975) reported a Late Ordovician coral fauna from the Uralba Beds, east of Manilla, an apparently fault bounded unit within the Peel Fault System. Recent collecting from the same locality has produced a conodont fauna that is correlatable with the WAF material (FUREY-GREIG, in prep.).

PALMIERI (1978) described a large fauna from the Fork Lagoons Beds in east-central Queensland that included *B. confluens*, *P. undatus* and *T. ?tumidus*. The stratigraphic and structural context of that unit is poorly understood, but the limestone may be correlated with the WAF material.

SIMPSON (1997) reported a fauna including *T. tumidus* and *B. confluens* from the Carriers Well Formation of north Queensland.

3.2. Comments on Supposed Silurian and Devonian Reports from the Wisemans Arm Formation

3.2.1. Silurian

A reported Silurian limestone (Loc. 4, Serpentinite Hill herein) proved barren in this investigation. BRADLEY (in PICKETT, 1982) reported *Kockelella variabilis* WALLISER from the locality and HALL (1978) documented a coral fauna also from this outcrop including *Acanthohalysites pycnoblatooides* (ETHERIDGE) and halysitids. This outcrop lies within a large

mass of serpentinite mapped within the WAF (LEITCH & CAWOOD, 1980) and considered by HALL (1978) to lie on a strand of the Peel Fault System, and thus cannot be demonstrated to be a component of the WAF.

3.2.2. Devonian

An Early Devonian age was advanced for some limestones in the WAF (CAWOOD, 1980; LEITCH & CAWOOD, 1980).

The material on which the age determinations were based has since been lost. "*Belodella resima*" was reported from an outcrop equivalent to Loc. 6 (Lone Boulder) herein, a single olistolith that yielded Eastonian conodonts. It seems likely that the element was misidentified, probably being a denticulate element of *Ansella*. Similarly, "*Icriodus* sp." and "*B. triangularis*" reported from CAWOOD's (1980) #UP 79138, Loc. 1, (Wallaby Mtn.) herein, may represent elements of *Icriodella* and *Ansella* respectively (this outcrop proved barren herein). Similarly, "*B. triangularis*" from #UP 79129, not located in the present investigation, cannot be verified.

3.2.3. Palaeoenvironment and Tectonic Setting

Conodont producing clasts from the WAF are generally recrystallised, but preserve evidence of relatively shallow conditions. They are fine grained detrital limestones with allochems including ostracods, echinoid fragments, brachiopod and trilobite remains and approximately 80 % fine, recrystallised grains with micritic envelopes, probably deposited at no greater than shelf depths. Clasts of well washed, favositid coralline limestone (barren when processed for conodonts) also occur at some localities (eg. Roscommon). The two lithologies represent a depth range that accords with presence in the fauna of species indicative of shallow environments, eg. *Belodina confluens* SWEET and others characteristic of deeper water, eg. *Periodon grandis* (ETHINGTON) and *Phragmodus undatus* BRANSON & MEHL (TROTTER & WEBBY, 1994).

Some limestone clasts from Loc. 8 (Roscommon) contain rounded felsic volcanic grains, evidence for input of volcanic detritus at time of limestone deposition. This, and the considerable lateral extent of Late Ordovician limestone outcrops (Fig. 1) [Late Ordovician conodonts were reported from the vicinity of Loc. 23, herein barren, (PICKETT, 1973, 1982) extending the extent of Late Ordovician outcrops east of Manilla], provide evidence for a chain of volcanic edifices with flanking reefal limestone platforms and associated carbonate fans, possibly located west of the area studied.

SCHEIBNER (1973) suggested that Ordovician limestone occurrences east of the Peel Fault System were off-scraped guyot cappings. Preserved guyot elements are characterised by associations of metabasalt, limestone and spicular & radiolarian cherts (eg. SANO & KANMERA, 1980). Such associations have not, thus far, been observed in any WAF limestone outcrops examined.

Correlation between the WAF fauna and that from the Trelawney Beds (PHILIP, 1966), and the close affinity of both with the Lachlan Fold Belt occurrences discussed earlier provides evidence that the two major tectonic elements of the New England Fold Belt developed in relative juxtaposition. This supports arguments that Early Palaeozoic elements in the New England Fold Belt are unlikely to have an exotic origin (LEITCH & CAWOOD, unpub. manuscript).

Table 2.
Distribution of conodont elements.

LOCALITY SPECIES	SPRING WATER	LONE BOULDER	ROSCOMMON					NINGADOO SOUTH	MANILLA VIEW				PERA SOUTH	BORA 2	BALFOUR	KELLY'S GULLY
			Clast No.						2	3	4	5				
			1	6	7	11	21									
<i>Ansella sp.</i>				7		2	4			1					1	
<i>Belodina confluens</i>		14	1	35	7	22	10	2		4	9	2				
<i>Belodina sp. A</i>															1	
<i>Belodina sp. B</i>		1														
<i>Belodina sp. C</i>		4														
<i>Belodina sp. D</i>		1				1					1					
<i>Belodina? sp. E</i>					1											
<i>Drepanoistodus suberectus</i>				3		6	7									
<i>Panderodus gracilis</i>	16	3		36	2	47	55	9	7	5	6			9	3	
<i>Panderodus panderi</i>				10		7	7				1					
<i>Panderodus nodus</i>				9		10	11			2						
<i>Panderodus sp. 1</i>				2		2	5									
<i>Panderodus sp. 2</i>				1		4	4									
<i>Peridon grandis</i>				5		3	8									
<i>Phragmodus undatus</i>			1	6	16	1	7									
<i>Pseudobelodina dispansa</i>			1	5		6	7			2						
<i>Pseudooneotodus beckmanni</i>								1								
<i>Pseudooneotodus mitratus</i>				1	1	1		1								
<i>Strachanognathus parvus</i>				1												
<i>Taoqupognathus tumidus</i>			2	10		12	14			1		2				
<i>Yaoxianognathus ani</i>				5		3	6									
<i>Yaoxianognathus? tunguskaensis</i>				7	1	6	7							?1		
Gen. et sp. indet. A								1								
Gen. et sp. indet. B				1	2											
Unassigned elements/fragments	8	7	5	41	6	32	30	33	9	5	7	5	1	25	6	4

4. Methods

From 33 localities collected, 128 samples of 8–10 kg were processed. Eleven produced conodonts with 7 yielding age diagnostic faunas numbering some 700 elements. Several samples were collected from each olistolith, a number of individual clasts from conglomeratic localities. Limestones were digested in acetic acid (10 %) at the Macquarie University Center for Ecostratigraphy and Palaeobiology acid leaching facility. Separation was by centrifuging with sodium polytungstate at 2.75 SG using the method described by ANDERSON et al. (1995).

Conodont illustration using a JEOL 35CF SEM at the University of Technology Sydney Microstructural Analysis Unit, was entirely digital. Images were collected from SEM via Moran Scientific Instruments Software and saved to disk, each image requiring 0.3 MB. Processing for clarity and background was done using Paint Shop Pro Ver. 3, plates were prepared in Corel Draw Ver. 6.0 and printed by a Fuji digital Pictography system.

5. Systematic Palaeontology

Phylum: Conodonta PANDER, 1856

Class: Conodonta PANDER, 1856

Genus: *Ansella* FÄHRAEUS & HUNTER, 1985

Type species: *Belodella jemtlandica* LÖFGREN, 1978.

Ansella sp.

(Pl. 1, Figs. 1–3)

Material: Thirteen elements from Loc. 8, one from Loc. 31, one from Loc. 12.

Remarks: Elements assigned to *Ansella* sp. herein include a denticulate and a single denticulate form. Adenticulate elements are similar to those illustrated by TROTTER & WEBBY (1994, Pl. 2, Figs. 5, 6) and to non-striate elements assigned to *Serraculodus?* sp. from the Cliefden Caves Limestone (ZHEN & WEBBY, 1995; Pl. 5, Figs. 1, 2). The adenticulate elements have deep basal cavities and faint surface striations on the upper half of their postero-lateral faces and are very similar to elements of *Ansella jemtlandica* (LÖFGREN) (FÄHRAEUS & HUNTER, 1985, Pl. 1, Figs. 3a, 3b). The denticulate element (Pl. 1, Fig. 1), despite poor preservation, resembles the asymmetrical element of *Ansella nevadensis* (ETHINGTON & SCHUMACHER) (FÄHRAEUS & HUNTER, 1985, Fig. 2, c).

Genus: *Belodina* ETHINGTON, 1959

Type species: *Belodus compressus* BRANSON & MEHL, 1933.

Belodina confluens SWEET, 1979

(Pl. 1, Figs. 4–9)

- 1966 *Belodina grandis* (STAUFFER); PHILIP, Fig. 2.
1966 *Belodina wykoffensis* (STAUFFER); PHILIP, Fig. 3.
1966 *Eobelodina fornicata* (STAUFFER); PHILIP, Figs. 4, 8.
1978 *Belodina compressa* (BRANSON & MEHL); PALMIERI, Pl. 3, Figs. 12–15, 20, 21, 23–25; Pl. 4, Figs. 1–17.
1979 *Belodina confluens* SWEET, Fig. 5, 10, 17; Fig. 6, 9.
1981 *Belodina confluens* SWEET, in ZIEGLER, Pl. 2, Figs. 8–14.
1988 *Belodina confluens* SWEET; NOWLAN & MCCrackEN, in NOWLAN et al., Pl. 1, Figs. 16–21.
1990 *Belodina confluens* SWEET; BERGSTRÖM, Pl. 3, Figs. 8–12.
1990 *Belodina confluens* SWEET; PICKETT & INGPEN, cover illustration, K.
1990 *Belodina confluens* SWEET; UYENYO, Pl. 1, Fig. 9, ?Fig. 8.
1990 *Belodina confluens* SWEET; SAVAGE, Fig. 9, 1–6.

1994 *Belodina confluens* SWEET; TROTTER & WEBBY, Pl. 2, Figs. 7, 18–20, 24, 25, 27–30.

1995 *Belodina confluens* SWEET; ZHEN & WEBBY, Pl. 1, Figs. 16–18.

1998 *Belodina confluens* SWEET; ZHEN et al., Pl. 1, Figs. 7–9.

Material: 118 elements; 95 from Loc. 8, 11 from Loc. 6, 2 from Loc. 12, 8 from Loc. 13 and 2 from Loc. 14.

Remarks: Compressiform elements display the diagnostic smoothly arcuate anterior margin described by SWEET (1979). These specimens show some variation in degree of compression and number of posterior denticles. A small number of specimens (Pl. 1, Figs. 10–11) have a lobe-like cusp (in lateral view), resembling *B. confluens?* illustrated by ZHEN & WEBBY (1995, Pl. 1, Fig. 20) and are tentatively assigned to this species. *B. confluens* constitutes 12 % of the WAF fauna, is common in central New South Wales faunas (ZHEN & WEBBY, 1995; PERCIVAL, 1997) and appears to be relatively abundant in New England occurrences, having also been reported from the Trelawney Beds (PHILIP, 1966), and recovered in material from the Uralba Beds (FUREY-GREIG, in prep.).

Belodina sp. A

(Pl. 1, Fig. 12)

Material: One compressiform element from Locality 31, Balfour.

Remarks: A robust compressiform element distinguished by a short, rapidly tapering heel and prolate adjacent denticles.

Belodina sp. B

(Pl. 1, Fig. 13)

Material: One grandiform element from Locality 6, Lone Boulder.

Remarks: The specimen is broken and displays an unusual, blunt heel with numerous, alate denticles that commence on its upper surface and extend along the cusp's posterior margin.

Belodina sp. C

(Pl. 1, Figs. 14–15)

Material: Three compressiform and one eobelodiniform element from Locality 6, Lone Boulder.

Remarks: The elements display a distinct, wide, short heel, while compressiform elements have blade like, horizontal denticles. They are similar to *Belodina* spp. E & F from the Malongulli Formation (TROTTER & WEBBY, 1994).

Belodina sp. D TROTTER & WEBBY, 1994

(Pl. 1, Figs. 16–17)

1994 *Belodina* sp. D TROTTER & WEBBY, Pl. 2, Figs. 16, 17, 22, 23.

Material: One compressiform element from each of locs 6, 8, 13.

Remarks: This species is distinguished by its markedly extended anterior margin, and resembles *Belodina* sp. D of TROTTER & WEBBY (1994).

Belodina? sp. E

(Pl. 1, Fig. 18)

Material: One compressiform element from Locality 8, Roscommon.

Remarks: The single element is poorly preserved, but is distinct in having an anterior margin similar to *B.* sp. D herein but with strongly prolate denticles.

Genus: *Drepanoistodus* LINDSTRÖM, 1971

Type species: *Oistodus forceps* LINDSTRÖM, 1955.

Drepanoistodus suberectus (BRANSON & MEHL, 1933)

(Pl. 2, Figs. 1–3)

- 1933 *Oistodus suberectus* BRANSON & MEHL, Pl. 35, Figs. 22–27.
1978 *Drepanoistodus suberectus* (BRANSON & MEHL); PALMIERI, Pl. 5, Figs. 14–31.
1988 *Drepanoistodus suberectus* (BRANSON & MEHL); NOWLAN & MCCrackEN in NOWLAN et al., Pl. 16, Figs. 19–22 (cum syn).
1989 *Drepanoistodus suberectus* (BRANSON & MEHL); MCCrackEN & NOWLAN, Pl. 2, Fig. 9.
1990 *Drepanoistodus suberectus* (BRANSON & MEHL); BERGSTRÖM, Pl. 2, Figs. 4–6; Pl. 3, Fig. 15.
1990 *Drepanoistodus suberectus* (BRANSON & MEHL); UYENO, Pl. 1, Figs. 13, 16–18.
1994 *Drepanoistodus suberectus* (BRANSON & MEHL); TROTTER & WEBBY, Pl. 5, Figs. 27–31.
1995 *Drepanoistodus suberectus* (BRANSON & MEHL); ZHEN & WEBBY, p. 282, Pl. 3, Figs. 8–10.
1995 *Drepanoistodus* sp.; FOWLER & IWATA, Fig. 2, 13.

Material: Twelve specimens from Locality 8, Roscommon.

Remarks: Elements from this fauna are either drepanodiform or suberectiform. There are insufficient specimens for comparison with the transition between suberectiform and drepanodiform described by ZHEN & WEBBY (1995, p. 282).

Genus: *Panderodus* ETHINGTON, 1959

Type species: *Paltodus unicastatus* BRANSON & MEHL, 1933.

Panderodus gracilis (BRANSON & MEHL, 1933)

(Pl. 2, Figs. 4–11)

- 1933 *Paltodus gracilis* BRANSON & MEHL, p. 108, Pl. 8, Figs. 20, 21.
1988 *Panderodus gracilis* (BRANSON & MEHL), NOWLAN & MCCrackEN, in NOWLAN et al., p. 21, Pl. 7, Figs. 1–10, 12, 13, 19 (cum syn).
1990 *Panderodus gracilis* (BRANSON & MEHL), UYENO, p. 69–70, Pl. 1, Figs. 14, 19, 20.
1990 *Panderodus gracilis* (BRANSON & MEHL), SAVAGE, p. 827, Fig. 7–10.
1994 *Panderodus gracilis* (BRANSON & MEHL), TROTTER & WEBBY, p. 483, Pl. 5, Figs. 1–4, 9, 10, 12–15.
1995 *Panderodus gracilis* (BRANSON & MEHL), ZHEN & WEBBY, p. 283, Pl. 3, Figs. 18–22.
1998 *Panderodus gracilis* (BRANSON & MEHL); ZHEN et al., Pl. 3, Figs. 8–12.

Material: One hundred and ninety specimens, sixteen from Loc. 3; three from Loc. 6; one hundred and forty from Loc. 8; nine from Loc. 9; seven from Loc. 12; five from Loc. 13, six from Loc. 14, nine from Loc. 28 and three from Loc. 31.

Remarks: This species is the most common in the WAF fauna constituting 23 % of the total. Using the nomenclature proposed by ORCHARD (1980), elements herein include M, Sc-Sb and Sa types. There is some variation in position and number of lateral costae and in degree of curvature of the cusp, a feature reported of this species from Cliefden Caves (ZHEN & WEBBY, 1995).

Panderodus nodus ZHEN, WEBBY & BARNES, 1998

(Pl. 2, Figs. 12–14)

- 1978 *Panderodus* sp. aff. *P. panderi*, PALMIERI, p. 23, Pl. 1, Fig. 13, possibly also Fig. 14, only.

- 1985 *Panderodus recurvatus* (RHODES), AN et al., Pl. 2, Figs. 17, 21.
1994 *Panderodus* sp. D, TROTTER & WEBBY, p. 484, Pl. 5, Figs. 24–25, doubtfully 26.
1995 *Panderodus* sp., ZHEN & WEBBY, p. 284, Pl. 4, Figs. 1–2.
1998 *Panderodus nodus* ZHEN et al., Pl. 3, Figs. 13–20.

Material: Thirty specimens from Loc. 8, Roscommon and two from Loc. 12, Manilla View 3.

Remarks: The WAF elements of this species are all more or less identical to the “long based” elements described by ZHEN et al., though there is a variation in degree of recurvature. It is noteworthy that “short based” elements are not represented in the WAF fauna. This new species, documented from the Bowan Park Group (ZHEN et al., 1998), is also known from the Fork Lagoon Beds in Queensland (PALMIERI, 1978), the Cliefden Caves Limestone (ZHEN & WEBBY, 1995) and the Malongulli Formation (TROTTER & WEBBY, 1994), and occurs in material recently recovered from the Uralba Beds in New England (FUREY-GREIG, in prep).

Panderodus panderi (STAUFFER, 1940)

(Pl. 2, Figs. 15–17)

- 1940 *Paltodus panderi* STAUFFER, Pl. 60, Figs. 8–9.
1988 *Panderodus?* *panderi* (STAUFFER); NOWLAN & MCCrackEN in NOWLAN et al., Pl. 7, Figs. 20, 23–25 (cum syn).
1990 *Panderodus panderi* (STAUFFER); UYENO, Pl. 1, Figs. 23, 29.
1995 *Panderodus panderi* (STAUFFER); ZHEN & WEBBY, p. 283–284, Pl. 3, Figs. 23–25.

Material: Twenty-four specimens from Locality 8 (Roscommon) and one from Loc. 13 (Manilla View 4).

Remarks: Two types of element are recognised, one is identical to similiform elements illustrated from the Canadian Eastern Arctic Archipelago (UYENO, 1990, Pl. 1, Figs. 23, 29), the other elements resemble “long based” elements from Cliefden Caves (ZHEN & WEBBY, 1995, Pl. 3, Figs. 23, 24). A noteworthy feature of the latter elements is a fine lateral carina preserved on some specimens.

Panderodus sp. 1

(Pl. 2, Fig. 18)

Material: Nine specimens from Locality 8, Roscommon.

Remarks: Falciform specimens closely resemble those illustrated from Cliefden Caves (ZHEN & WEBBY, 1995) and those assigned, as M elements, to *P. gracilis* from the Malongulli Formation (TROTTER & WEBBY, 1994, Pl. 5, Figs. 13, 14), having keel-like anterior and posterior margins, a feature not characteristic of *P. gracilis*.

Panderodus sp. 2

(Pl. 2, Fig. 20)

Material: Nine specimens from Loc. 8, Roscommon.

Remarks: The elements are short cones that have almost cylindrical cross-sections and a lateral furrow that extends to the apex.

Genus: *Periodon* HADDING, 1913

Type species: *Periodon aculeatus* HADDING, 1913.

Periodon grandis (ETHINGTON, 1959)

(Pl. 2, Figs. 21, 22; Pl. 3, Figs. 1, 2)

- 1959 *Loxognathus grandis* ETHINGTON, p. 281, Pl. 40, Fig. 6.
1981 *Periodon grandis* (ETHINGTON); LINDSTRÖM, in ZIEGLER, p. 243–244, Pl. 1, Figs. 13–18.
1989 *Periodon grandis* (ETHINGTON); MCCrackEN & NOWLAN, p. 1889, Pl. 3, Figs. 7–9.

- 1990 *Periodon grandis* (ETHINGTON); BERGSTRÖM, p. 11, Pl. 3, Fig. 7.
 1994 *Periodon grandis* (ETHINGTON); TROTTER & WEBBY, p. 484, Pl. 4, Figs. 13–14, 27, 28.
 1995 *Periodon grandis* (ETHINGTON); FOWLER & IWATA, Fig. 2, 1, 4–5 (only).
 1995 *Periodon grandis* (ETHINGTON); ZHEN & WEBBY, p. 284, Pl. 4, Figs. 3, 4.
 1998 *Periodon grandis* (ETHINGTON); ZHEN et al., Pl. 4, Figs. 19–21.

Material: Ten ramiform and 6 oistodontiform elements from Locality 8, Roscommon.

Remarks: Oistodontiform elements in this collection are similar to material reported from the Craighead Limestone (BERGSTRÖM, 1990), the Malongulli Formation (TROTTER & WEBBY, 1994) and the Cliefden Caves Limestone (ZHEN & WEBBY, 1995). Ramiform elements are of two types, some (Pl. 3, Fig. 1) resemble that illustrated from the Cliefden Caves Limestone (ZHEN & WEBBY, 1995, pl. 4, Fig. 4), others are identical to “prioniodini-form” elements from the Malongulli Formation (TROTTER & WEBBY, 1994, Pl. 4, Fig. 14) and from Akpatok Island (MCCRACKEN & NOWLAN, 1989, Pl. 3, Fig. 9). This species has been also reported from the Triangle Group of central NSW (FOWLER & IWATA, 1995).

Genus: *Phragmodus* BRANSON & MEHL, 1933

Type species: *Phragmodus primus* BRANSON & MEHL, 1933.

Phragmodus undatus BRANSON & MEHL, 1933

(Pl. 3, Figs. 3–9)

- 1933 *Phragmodus undatus* BRANSON & MEHL, p. 115–116, Pl. 8, Figs. 22–26.
 1966 *Phragmodus undatus* BRANSON & MEHL, PHILIP, p. 112, Fig. 1, 5, 6, 7, 9.
 1988 *Phragmodus undatus* BRANSON & MEHL, NOWLAN & MCCRACKEN in NOWLAN et al., p. 26–27, Pl. 10, Figs. 1–3, 6–7 (cum syn).
 1989 *Phragmodus undatus* BRANSON & MEHL, MCCRACKEN & NOWLAN, Pl. 3, Figs. 10–12.
 1994 *Phragmodus undatus* BRANSON & MEHL, TROTTER & WEBBY, p. 485, Pl. 6, Figs. 2–11.
 1998 *Phragmodus undatus* BRANSON & MEHL; ZHEN et al., Pl. 5, Figs. 1–4.

Material: Thirty-one specimens from Loc. 8, Roscommon.

Remarks: The material, dominantly Sc and M elements, is representative of the species, closely resembling material reported from the Trelawney Beds (PHILIP, 1966), the Fork Lagoons Beds (PALMIERI, 1978) and the Malongulli Formation (TROTTER & WEBBY, 1994). Elements typically lack preservation of the more delicate parts, a characteristic of the type material (BRANSON & MEHL, 1933). The specimens illustrated from the Trelawney and Fork Lagoons beds appear to be better preserved, possibly recording evidence for lower energy environments.

Genus: *Pseudobelodina* SWEET, 1979

Type species: *Belodina kirki* STONE & FURNISH, 1959.

Pseudobelodina dispansa (GLENISTER, 1957)

(Pl. 3, Figs. 10–13)

- 1957 *Belodus dispansus* GLENISTER, p. 729–730, Pl. 88, Figs. 14, 15.
 1978 *Belodina* sp. D, PALMIERI, Pl. 3, Figs. 18, 19, 22.
 1988 *Pseudobelodina? dispansa* (GLENISTER), NOWLAN & MCCRACKEN, in NOWLAN et al., p. 30–31, Pl. 12, Figs. 8–26.
 1989 *Pseudobelodina? dispansa* (GLENISTER), MCCRACKEN & NOWLAN, Pl. 4, Figs. 2–4.

- 1990 *Pseudobelodina dispansa* (GLENISTER), UYENO, Pl. 1, Figs. 24, 30.

- 1994 *Pseudobelodina dispansa* (GLENISTER), TROTTER & WEBBY, Pl. 1, Figs. 21–29, 31–33 (only).

Material: Nineteen specimens from Loc. 8, 2 from Loc. 12.

Remarks: These elements display a similar variation in degree of recurvature and cusp length as the material from the Malongulli Formation (TROTTER & WEBBY, 1994).

Genus: *Pseudooneotodus* DRYGANT, 1974

Type species: *Oneotodus? beckmanni* BISCHOFF & SANNE-MANN, 1958.

Pseudooneotodus beckmanni (BISCHOFF & SANNE-MANN, 1958)

(Pl. 3, Fig. 14)

- 1958 *Oneotodus? beckmanni* BISCHOFF & SANNE-MANN, p. 98, Pl. 15, Figs. 22–25.
 1981 *Pseudooneotodus beckmanni* (BISCHOFF & SANNE-MANN); NOWLAN & BARNES, p. 23, Pl. 2, Figs. 20, 21 (cum syn.).
 1990 *Pseudooneotodus beckmanni* (BISCHOFF & SANNE-MANN); UYENO, p. 99–100, Pl. 1, Figs. 36, 37.
 1994 *Pseudooneotodus beckmanni* (BISCHOFF & SANNE-MANN); TROTTER & WEBBY, p. 486, Pl. 4, Figs. 29, 30.
 1995 *Pseudooneotodus beckmanni* (BISCHOFF & SANNE-MANN); FUREY-GREIG, Pl. 1, Fig. 3.

Material: One specimen from locality 9 (Ningadoo South).

Remarks: The specimen from locality 9, though being less strongly recurved, closely resembles the material from the Vaureal Formation (NOWLAN & BARNES, 1981) and from the Malongulli Formation (TROTTER & WEBBY, 1994).

Pseudooneotodus mitratus (MOSKALENKO, 1973)

(Pl. 3, Fig. 15)

- 1973 *Ambalodus mitratus mitratus* MOSKALENKO, p. 86, Pl. 17, Figs. 9–11.
 1988 *Pseudooneotodus mitratus* (MOSKALENKO, 1973); NOWLAN & MCCRACKEN, p. 34, Pl. 16, Figs. 2–6.
 1990 *Pseudooneotodus mitratus* (MOSKALENKO, 1973); POHLER & ORCHARD, Pl. 6, Fig. 12.
 1994 *Pseudooneotodus mitratus* (MOSKALENKO, 1973); TROTTER & WEBBY, Pl. 4, Figs. 21, 22.
 1995 *Pseudooneotodus mitratus* (MOSKALENKO, 1973); ZHEN & WEBBY, Pl. 4, Figs. 16, 17.
 1998 *Pseudooneotodus mitratus* (MOSKALENKO, 1973); ZHEN et al., Pl. 5, Figs. 14, 15.

Material: Three elements from Loc. 8, one from Loc. 9.

Remarks: Previously illustrated specimens of this species display a degree of morphological elasticity, most notable in the forms illustrated by TROTTER & WEBBY (1994). Specimens herein most closely resemble those from the Cliefden Caves Limestone illustrated by ZHEN & WEBBY (1995) in having no nodes and appear to be more symmetrical than the type specimens from Siberia (MOSKALENKO, 1973).

Genus: *Strachanognathus* RHODES, 1955

Type species: *Strachanognathus parvus* RHODES, 1955.

Strachanognathus parvus RHODES, 1955

(Pl. 3, Fig. 16)

- 1955 *Strachanognathus parvus* RHODES, p. 132, Pl. 7, Fig. 16, Pl. 8, Figs. 1–4.
 1978 *Strachanognathus parvus* RHODES; LÖFGREN, p. 112, Pl. 1, Fig. 29.

- 1978 *Strachanognathus parvus* RHODES; PALMIERI, p. 27, Pl. 6, Figs. 27, 28.
 1980 *Strachanognathus parvus* RHODES; ORCHARD, Pl. 14, Figs. 34, 35.
 1990 *Strachanognathus parvus* RHODES; BERGSTRÖM, Pl. 1, Fig. 10.
 1994 *Strachanognathus parvus* RHODES; TROTTER & WEBBY, Pl. 4, Figs. 24–25.

Material: One damaged specimen from Locality 8 (Roscommon).

Remarks: This damaged element preserves sufficient detail for assignment to *S. parvus* but has some noteworthy features. The lateral costa on the cusp appears slightly twisted, a faint costa flanks the anterior denticle and small protrusion on the anterior face of the cusp (broken) may record development of a small extra denticle. SERPAGLI (1967, p. 99–100) regarded such variations as pathological.

Genus: *Taoqupognathus* AN, 1985

Type species: *Taoqupognathus blandus* AN, 1985.

Taoqupognathus tumidus TROTTER & WEBBY, 1994

(Pl. 4, Figs. 1–9)

- 1978 *Drepanodus? altipes* (?) PALMIERI, Pl. 2, Figs. 24–25.
 1994 *Taoqupognathus tumidus* TROTTER & WEBBY, p. 487–486, Pl. 7, Figs. 10–24.
 1998 *Taoqupognathus tumidus* TROTTER & WEBBY; ZHEN et al., Pl. 7, Figs. 1–9.

Material: Thirty-eight elements from Locality 8, two from Loc. 12 and two from Loc. 14.

Remarks: *T. tumidus* was first described from the Malongulli Formation by TROTTER & WEBBY (1994), being the youngest species in a lineage of three (*T. philipi* SAVAGE – *T. blandus* AN – *T. tumidus* TROTTER & WEBBY) known from the Late Ordovician of New South Wales (ZHEN & WEBBY, 1995). Elements in this collection essentially match the apparatus described by TROTTER & WEBBY, but display the following noteworthy features. Using the nomenclature of ZHEN & WEBBY (1995), the *T. tumidus* apparatus from New England includes an Sc6 element (Pl. 4, Fig. 3) with a bilobed posterior process (3 specimens) previously recognised only in *T. blandus*, and recently also reported from *T. tumidus* of the Bowan Park Group faunas (ZHEN et al., 1998). A few P elements (Pl. 4, Fig. 7) from Locality 8 display a prominent knob like protrusion of the anterior margin at the point of its posteriorly directed inflexion, while others from the same clasts (Pl. 4, Fig. 5) are identical to the type material. Two tall, relatively straight elements with a moderately bulged posterior margin (Pl. 4, Fig. 2) appear unique to the present collection and may represent a rare and unrecognised Sc element as they show similarity to the Sc5 element of *T. philipi*; one other tall element (Pl. 4, Fig. 9) with a strongly posteriorly directed cusp appears to be an aberrant form, possibly a P element. The Wisemans Arm collection is dominated by Sb elements (33%). *T. tumidus* has also been reported from the Sofala Volcanics of the Capertee Rise in the Lachlan Fold Belt (PERCIVAL, 1997).

Genus: *Yaoxianognathus* AN, 1985

Type species: *Yaoxianognathus yaoxianensis* AN, 1985.

Yaoxianognathus ani ZHEN, WEBBY & BARNES

(Plate 4, Figs. 10–14)

- 1994 *Ozarkodina sesquipetalis* NOWLAN & BARNES; TROTTER & WEBBY, p. 483, Pl. 6, Figs. 13–25.
 1998 *Yaoxianognathus ani* ZHEN, WEBBY & BARNES, Pl. 8, Figs. 1–16.

Material: Fourteen elements from locality 8 (Roscommon).

Remarks: The small collection of elements assigned to *Y. ani* include most elements documented by ZHEN et al. (1998), and closely resemble elements assigned to *O. sesquipetalis* by TROTTER & WEBBY (1994). This species occurs in the *T. tumidus* fauna of the Bowan Park Group, (ZHEN et al., 1998).

Yaoxianognathus? tunguskaensis

(MOSKALENKO, 1973)

(Pl. 4, Figs. 15–17)

- 1973 *Phragmodus? tunguskaensis* MOSKALENKO, 1973, p. 74, Pl. XII, Figs. 1–3.
 1983 *Phragmodus? (Spinodus?) tunguskaensis* MOSKALENKO; MOSKALENKO, Fig. 4.Z.
 1985 *Oulodus? tunguskaensis* (MOSKALENKO); AN, p. 108, Pl. 2, Fig. 9.
 1988 *Spinodus? sp. A* NOWLAN & MCCrackEN, p. 38, Pl. 17, Figs. 17–22.
 1990 *Phragmodus? tunguskaensis* MOSKALENKO; SAVAGE, Fig. 10.7, 10.8.
 1994 *Yaoxianognathus? tunguskaensis* (MOSKALENKO); TROTTER & WEBBY, Pl. 4, Figs. 18–20.
 1995 *Yaoxianognathus? tunguskaensis* (MOSKALENKO, 1973); ZHEN & WEBBY, p. 289–290, Pl. 6, Fig. 22.
 1998 *Yaoxianognathus? tunguskaensis* (MOSKALENKO); ZHEN et al., Pl. 6, Fig. 14, Pl. 7, Fig. 24, Pl. 8, Figs. 17–22.

Material: Twenty-one elements from Locality 8, Roscommon, one from Loc. 28 (doubtfully).

Remarks: The small number of elements in this collection are all more or less broken, but they display sufficient features for assignment to this little understood species. Eight specimens herein are identical to those from the Malongulli Formation (TROTTER & WEBBY, 1994), lacking small denticles between cusp and adjacent denticles and displaying a reduced anterior denticle. A single element (Pl. 4, Fig. 16) is similar to that illustrated from Cliefden Caves (ZHEN & WEBBY, 1995, Pl. 6, Fig. 22).

One broken element from Locality 8 is very similar to the Sd elements illustrated by ZHEN et al. (1998, Pl. 6, Fig. 14; Pl. 8, Fig. 18) in their recent multi element revision of this species.

Gen. et sp. indet. A

(Pl. 4, Fig. 18)

Material: One partial specimen from Locality 9, Ningadood South.

Remarks: This, the only platform conodont in the present fauna, is unfortunately broken. The preserved part is inadequate for assignment, but the arrangement of the short blade and nodes/ridges on the platform is similar to Pa elements of the Silurian genus *Aspidognathus*.

Gen. et sp. indet. B

(Pl. 4, Fig. 19)

Material: Three specimens from Locality 8.

Remarks: The three elements from Roscommon cannot be accommodated in any known Late Ordovician genus. They feature three bar-like processes, two of which

(posterior and anterior?) are relatively long with a reduced lateral? process. The long processes support four or five stout denticles with an elliptical cross-section. The short process is adenticulate. One element has slightly curved bars, two have straight ones.

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Plate 1

Manuskript bei der Schriftleitung eingelangt am 2. Oktober 1998

- Figs. 1–3: ***Ansella* sp.**
 Fig. 1: Asymmetrical element, AMF 103997, Loc. 8.
 Fig. 2: Adenticulate element, AMF 103998, Loc. 8.
 Fig. 3: Adenticulate element, AMF 103999, Loc. 8.
- Figs. 4–9: ***Belodina confluens* SWEET, 1979.**
 Fig. 4: Compressiform element, AMF 104000, Loc. 8.
 Fig. 5: Grandiform element, AMF 104001, Loc. 8.
 Fig. 6: Eobelodiniiform element, AMF 104002, Loc. 8.
 Fig. 7: Grandiform element, AMF 104003, Loc. 8.
 Fig. 8: Eobelodiniiform element, AMF 104004, Loc. 8.
 Fig. 9: ?juvenile grandiform element, AMF 104005, Loc. 8.
- Figs. 10–11: ***Belodina ? confluens.***
 Fig. 10: Grandiform element, AMF 104006, Loc. 6.
 Fig. 11: Grandiform element, AMF 104007, Loc. 13.
- Fig. 12: ***Belodina* sp. A.**
 Compressiform element, AMF104008, Loc. 31.
- Fig. 13: ***Belodina* sp. B.**
 Grandiform element, AMF 104009, Loc. 6.
- Figs. 14–15: ***Belodina* sp. C.**
 Fig. 14: Compressiform element, AMF 104010, Loc. 6.
 Fig. 15: Eobelodiniiform element, AMF 104011, Loc. 6.
- Figs. 16–17: ***Belodina* sp. D.**
 Fig. 16: Compressiform element, AMF 104012, Loc. 13.
 Fig. 17: Compressiform element, AMF 104013, Loc. 8.
- Fig. 18: ***Belodina* sp. E.**
 Compressiform element, AMF 104014, Loc. 8.

All scale bars = 100 microns.

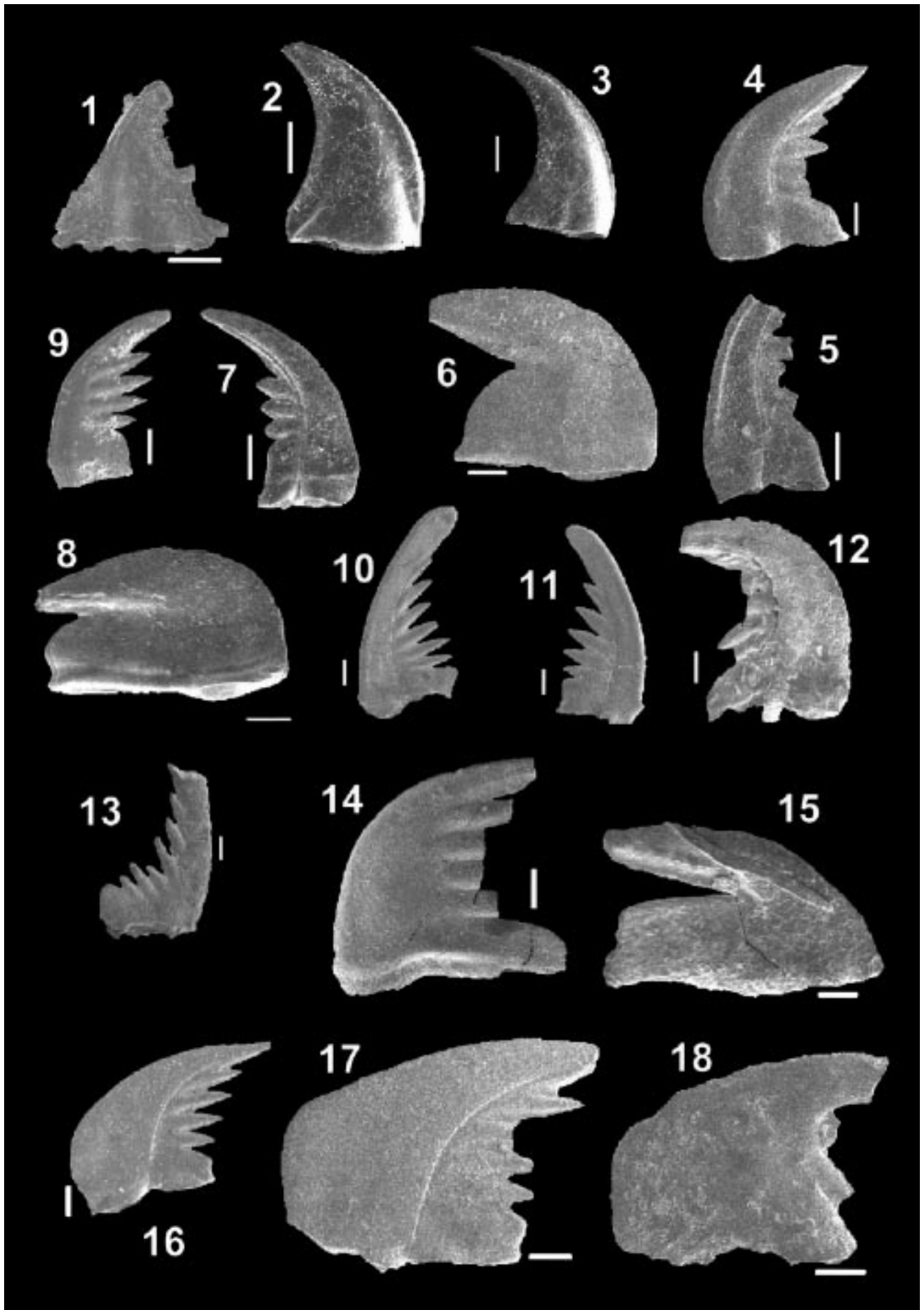


Plate 2

- Figs. 1–3: *Drepanoistodus suberectus* (BRANSON & MEHL, 1933).
Figs. 1,2: Drepanodiform element, AMF 104015, AMF 104016, Loc. 8.
Fig. 3: Suberectiform element, AMF 104017, Loc. 8.
- Figs. 4–11: *Panderodus gracilis* (BRANSON & MEHL, 1933).
Figs. 4,10: Sc-b elements, AMF 104018, AMF 104019, Locs. 8 & 3.
Fig. 5: Sb element, AMF 104020, Loc. 6.
Figs. 6,7,9,11: ?Sa elements: AMF 104021, AMF 104022, AMF 104023, AMF 104024, Loc. 8.
Fig. 8: M element, AMF 104025, Loc. 8.
- Figs. 12–14: *Panderodus nodus* ZHEN, WEBBY & BARNES, 1998.
Figs. 12–14: Long based elements, AMF 104026, AMF 104027, AMF 104028, Loc. 8.
Figs. 15–17: *Panderodus panderi* (STAUFFER, 1940).
Figs. 15,16: Long-based elements, AMF 104029, AMF 104030, Loc. 8.
Fig. 17: Similiform element, AMF 104031, Loc. 8.
- Fig. 18: *Panderodus* sp. 1.
M element, AMF 104032, Loc. 8.
- Fig. 19: *Panderodus* sp. 2
?tortiform element, AMF 104033, Loc. 8.
- Figs. 20,21: *Periodon grandis* (ETHINGTON, 1959).
Fig. 20: Oistodontiform element, AMF 104 034, Loc. 8.
Fig. 21: Ramiform element, AMF 104035, Loc. 8.

All scale bars = 100 microns.

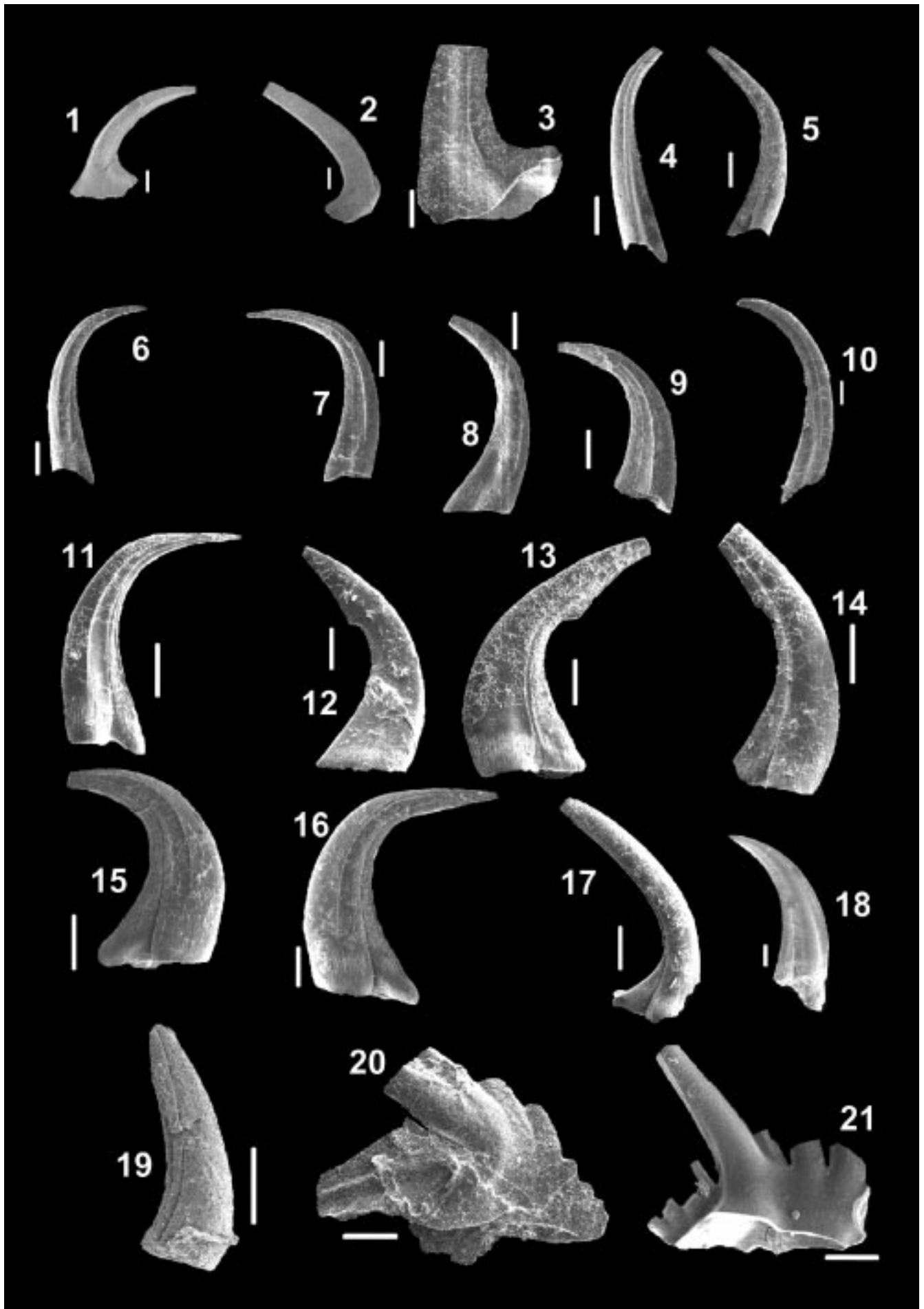


Plate 3

- Figs. 1–2: *Periodon grandis* (ETHINGTON, 1959).
Fig. 1: Ramiform element, AMF 104036, Loc. 8.
Fig. 2: ?element, AMF 104037, Loc. 8.
- Figs. 3–9: *Phragmodus undatus* (BRANSON & MEHL, 1933).
Fig. 3: Sb element, AMF 104038, Loc. 8.
Figs. 4–6: Sc elements, AMF 104039, AMF 104040, Loc. 8.
Fig. 7: ?juvenile element, AMF 104041, Loc. 8.
Figs. 8,9: M elements, AMF 104042, AMF 104043, Loc. 8.
- Figs. 10–13: *Pseudobelodina dispansa* (GLENISTER, 1957).
?grandiform elements, AMF 104044, AMF 104045, AMF 104046, 10, 12 Loc. 8, 11, Loc. 12.
- Fig. 13: ?grandiform element, AMF 104047, Loc. 8.
- Fig. 14: *Pseudooneotodus beckmanni* (BISCHOFF & SANNEMANN, 1958).
AMF 104048, Loc. 9.
- Fig. 15: *Pseudooneotodus mitratus* (MOSKALENKO, 1973).
AMF 104049, Loc. 8.
- Fig. 16: *Strachanognathus parvus* RHODES, 1955.
AMF 104050, Loc. 8.

All scale bars = 100 microns.
Note: 4,5 same scale.

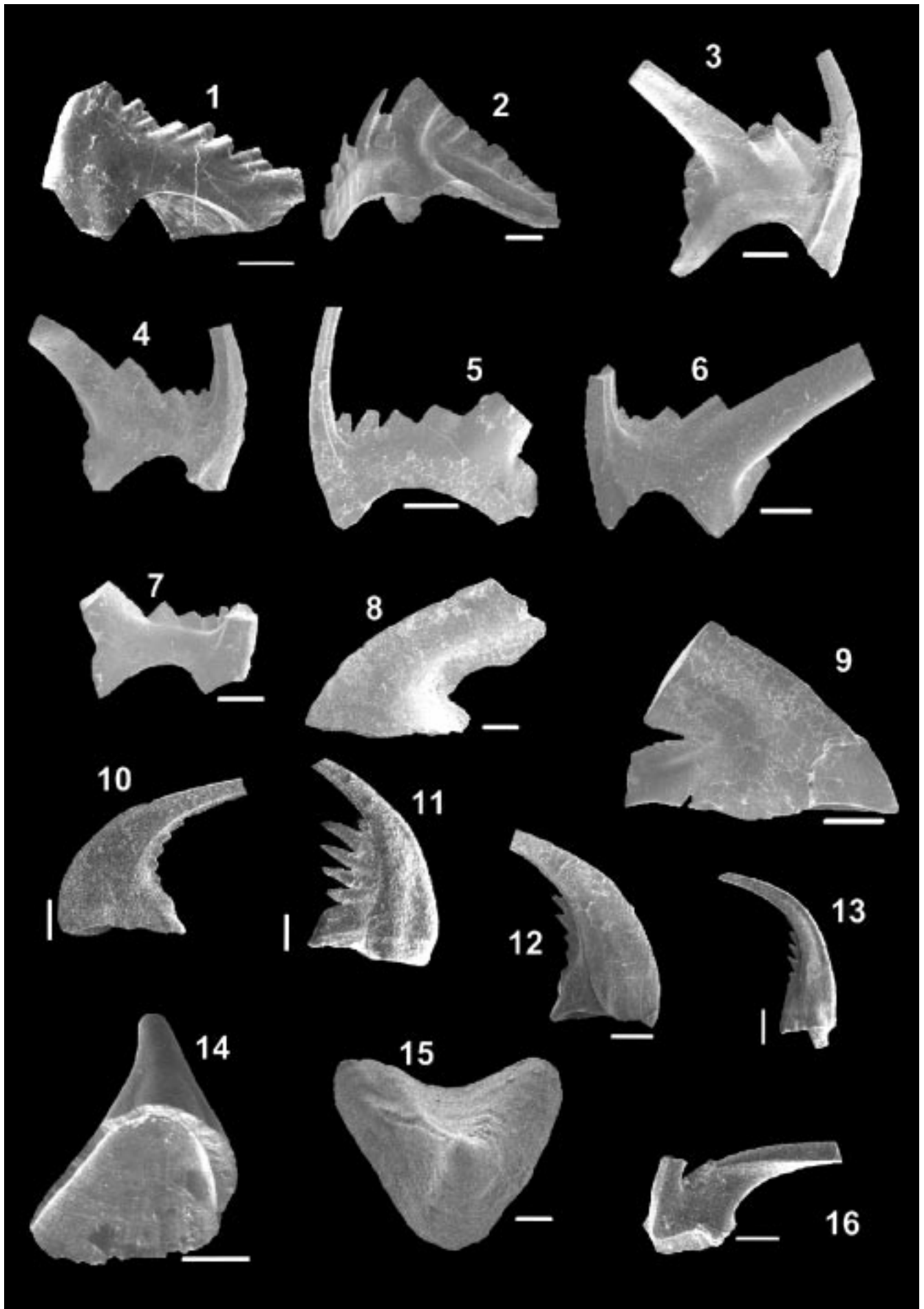


Plate 4

- Figs. 1–9: ***Taoqupognathus tumidus* TROTTER & WEBBY, 1994.**
Fig. 1: Sc3 element, AMF 104052, Loc. 12.
Fig. 2: Sc2 element, AMF 104053, Loc. 8.
Fig. 3: Sc6 element, AMF 104054, Loc. 8.
Fig. 4: M element, AMF 104055, Loc. 8.
Fig. 5: P element, AMF 104056, Loc. 8.
Fig. 6: ?Sc5 element, AMF 104057, Loc. 8.
Fig. 7: P element, AMF 104058, Loc. 8.
Fig. 8: Sb element, AMF 104059, Loc. 8.
Fig. 9: Tall element, AMF 104060, Loc. 8.
- Figs. 10–14: ***Yaoxianognathus ani* ZHEN, WEBBY & BARNES, 1998.**
Figs. 10,11: Pa elements, AMF 104061, AMF 104062, Loc. 8.
Fig. 12: Sd element, AMF 104063, Loc. 8.
Fig. 13: ?Sc element, AMF 104064, Loc. 8.
Fig. 14: Pb element, AMF 104065, Loc. 8.
- Figs. 15–17: ***Yaoxianognathus tunguskaensis* (MOSKALENKO, 1973).**
Figs. 15,16: Sc elements, AMF 104066, AMF 104067, Loc. 8.
Fig. 17: ?Sd element, AMF 104068, Loc. 8.
- Fig. 18: **Gen. et sp. indet. A.**
AMF 104069, Loc. 9.
- Fig. 19: **Gen. et sp. indet. B.**
AMF 104070, Loc. 8.

All scale bars = 100 microns.

