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1. Introduction

The differentiation of various facies zones developed during Triassic times ends in the Jurassic. During the Early and Middle Jurassic basin sediments are dominating. There is just some evidence indicating the existence of carbonate platforms (e.g. Lienz Dolomites, BLAU 1989). In contrast, during the Malmian great carbonate platforms with large shallow water sedimentation areas existed (Autochthonous Late Jurassic, Northern Calcareous Alps, Waschberg-Zone, Prättigau).

For a comprehensive overview only the most important and well known areas will be discussed. Some minor evidences for shallow water sedimentation areas which are derived from the study of pebbles are just mentioned below. The source of some Upper Jurassic pebbles of several flysch sediments (FLÜGEL & FAUPL 1987) or of sediments in the Helvelic Zone (FAUPL & SCHNABEL 1987) is not well known. Some large olistoliths in the "Buňlmergelserle" yield even a typical Late Jurassic algal assemblage (SCHWINGENSCHILÖGEL 1981). They could be derived from a now eroded, during Jurassic time uplifted block, surrounded by deeper marine sediments (FLÜGEL & FAUPL 1987).

Some comments on the Jurassic paleogeography may help for better understanding of the today's situation (fig. 1) being a result of Alpine orogeny. Before Alpine orogeny, during Cretaceous and Paleogene, the southernmost belt was

the Eastern Alpine Unit. This is the root of the Northern Calcareous Alps. To the north the Central – Tethyan Region with the Penninic ocean occurred. Part of this sedimentary realm is found e.g. in the Prättigau. Further to the north, the sedimentary unit of the Helvelic Zone was located. In the northernmost belt, fringing even southern parts of the Bohemian Massif, during the Late Jurassic a connection between Bavaria to Czechia existed. These sediments were even in the 19th century recognized as a non-Alpine element and were called "outer alpine", or "germanotype facies". Fossils like *Diceras* sp. or algal – sponge bioherms as they are typical for this most northern sedimentary area, are not known from the Alpine realm. During Cretaceous and Tertiary times these carbonate platforms were covered by several kilometers of molasse-type sediments. As a result of Alpine orogeny some parts of these (now covered) carbonate platforms were pressed up to the surface through a thick pile of molasse sediments, forming the "Klippen" of the Waschbergzone. These rootless "Klippen" which are now exposed at the eastern border of the Bohemian Massif (Ernstbrunn, Staatz, Falkenstein, Pavlovske vrchy,...) are the only remains of the "germanotype facies" at the surface.

The entire Eastern Alpine Unit was overthrust over the Penninic sediments during Alpine orogeny, therefore sediments which were primarily deposited south of the Penninic ocean are now in the north of it. Due to these movements, the Penninic Sulzfluh Limestone of the Prättigau shows strong diagenetic alterations, whereas the Plassen Limestone (which was overthrust over the former) is better preserved.





The Helvetic Zone, which may contain some Late Jurassic limestone pebbles and even olistoliths in the marly sediments, was overthrust in the same way by the Northern Calcareous Alps. Only a small line of Helvetic sediments (together with the Flysch zone) is visible between the Molasse zone and the Northern Calcareous Alps. This tectonic situation should be kept in mind with regard to paleobiogeographic considerations.

2. Lias

Although BÖHM (1992: 95) states that there are no evidences of algae during the Liassic in the Northern Calcareous Alps, he found some deeper marine stromatolites (BÖHM 1992; BÖHM & BRACHERT 1993 (=below the photic zone). These stromatolites could be caused by some blue - green algae, like *Frutexites* sp. (for discussion see BÖHM & BRACHERT 1993).

From the Early Liassic of the Lienz Dolomites (= western part of the Gaistal Alps), the Lavantler Breccia yields some components showing shallow-water influence. BLAU & SCHMIDT (1988) have recorded oncoids (1cm) as shallow-water indicators from this formation, without giving any detailed description, as clear prove of shallow-water. In addition, BLAU (1989) described "stromatolitic" crusts covering foraminifera. The dasyclad *Palaeodasycladus mediterraneus* (PIA) should be kept in mind as a significant fossil for Liassic sediments in the Southern Alps.

3. Dogger

In general for the Dogger a similar situation as for the Lias must be assumed. Basin sediments were dominating and there is no evidence for shallow - water

carbonate platforms. BÖHM & BRACHERT (1993) reported some deeper marine stromatolites from the "Klaus-kalk" in a comparable bathymetric position as in the Liassic described above. In conclusion BÖHM & BRACHERT (1993) found that *Frutexites* sp. is a common component of deep-water stromatolites, acting mainly as a dweller instead of a constructor of these stromatolites.

4. Malm

The Late Jurassic is an important geological period for calcareous algae. This is due to the Callovian transgression that overslepted the Middle Jurassic basins. As a result, great areas with shallow-water sedimentation with intensive carbonate sedimentation arose.

4.1. Autochthonous Late Jurassic

All the information on the autochthonous Late Jurassic is based on the intensive drilling activities of the past decades.

For the south-western part of the Bohemian massif NACHTMANN & WAGNER (1987) reported the development of an extensive carbonate platform during the Late Jurassic (fig. 2). In the western part of the basin some Oxfordian and Kimmeridgian algal-sponge banks up to 200 m thickness are capped by coral reefs. In the eastern part they interfinger with oolithic grainstones, sponge biostromes, coral reefs and sometimes with lagoonal, semi-restricted banks. These banks are characterised by "birdseyes" and "black pebbles", indicating a very shallow regime. During the Purbeckian (= Early Berriasian) the whole carbonate platform was covered by thin-bedded fine crystalline dolomites, cherty limestones,

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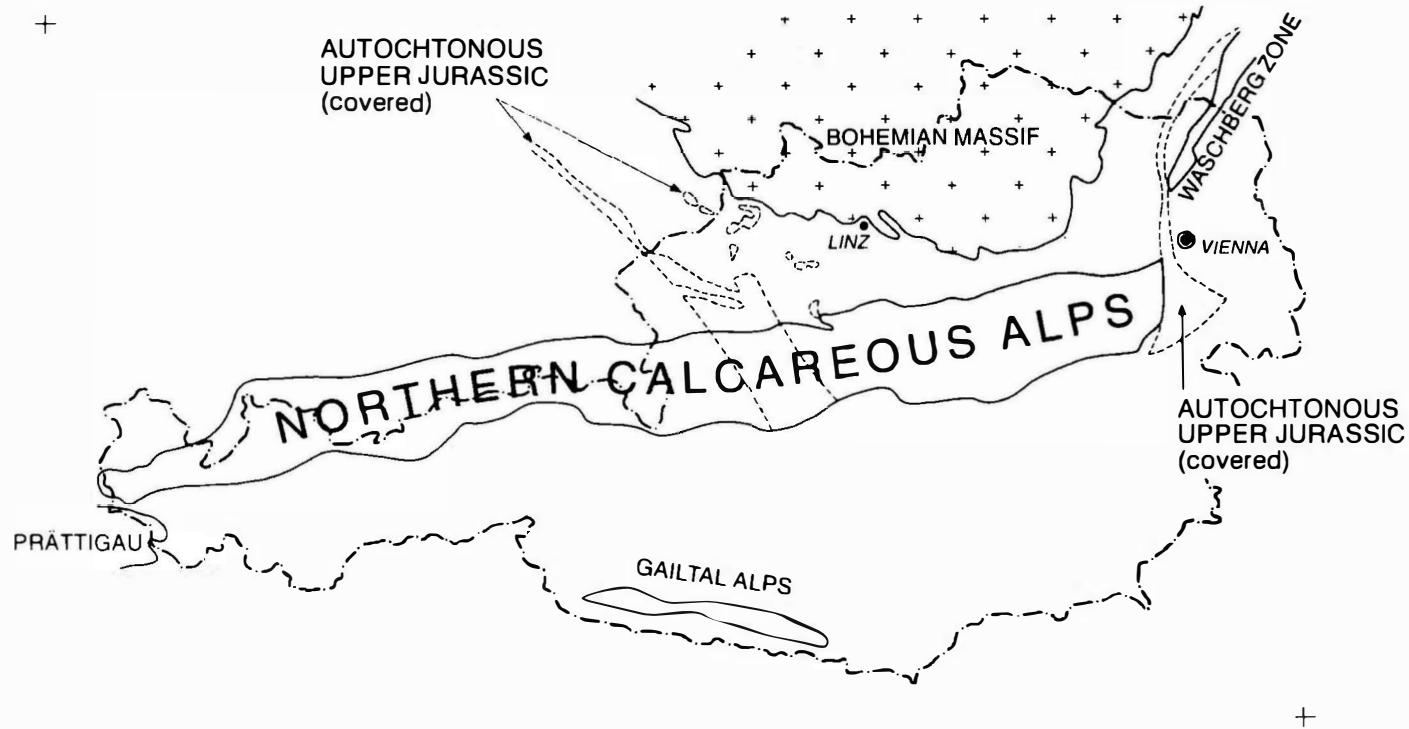
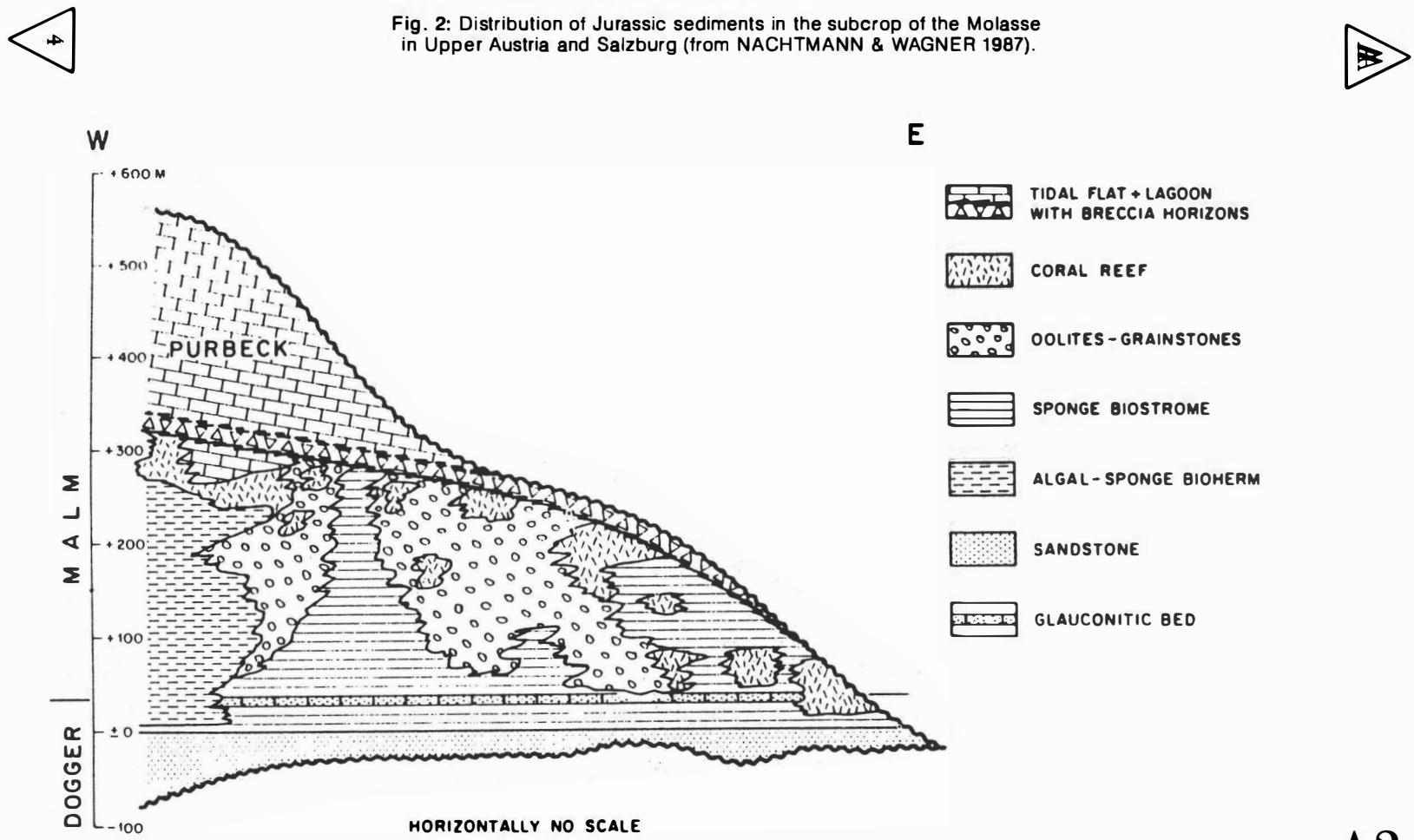


Fig. 1: Distribution of Jurassic sediments in Austria (after SAUER et al. 1992).



Fig. 2: Distribution of Jurassic sediments in the subcrop of the Molasse
in Upper Austria and Salzburg (from NACHTMANN & WAGNER 1987).



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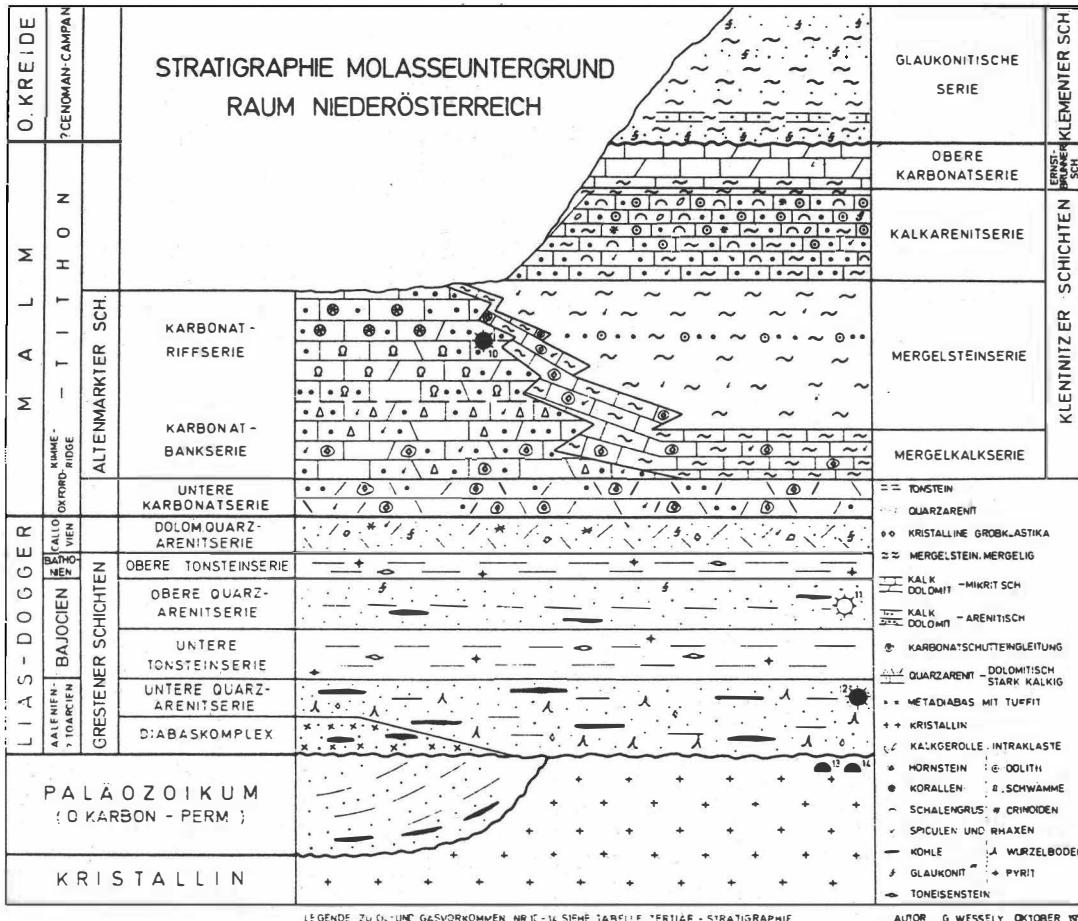


Fig. 3: Distribution of sediments in the subcrop of the Molasse in Lower Austria (from BRIX et al. 1977).



breccias and stromatolites. Characeae indicate some freshwater influence (NACHTMANN & WAGNER 1987). For the eastern part, fringing the Bohemian Massif, some detailed information was worked out by LADWEIN (1976). Additionally, a lot of information is available from Czechian geologists. A comprehensive overview is given by ELIAS & WESSELY (1990). During the Oxfordian and Kimmeridgian up to 750 m of platform carbonates (Altenmarkt Formation) developed (fig. 3). In the upper part some algal – sponge carbonates are found, the top of this formation is built by coral reefs (= "Karbonatriffserie"). Laterally the reef complexes are replaced sometimes by oolithic – bioclastic deposits (SAUER et al. 1992). During the Tithonian this sequence was covered by the terrigenous, bioclastic Kurdejov Formation (= "Kalkarenitserie"), overlain by the Ernstbrunn Limestone (= "Obere Karbonatserie").

**Flora of the Autochthonous Jurassic
from Upper Austria
(STEIGER, pers. comm.)**

Bacinella irregularis RADOICIC,
Lithocodium aggregatum ELLIOTT,
L. morikawai ENDO,
Tubiphytes sp.,
Globochaeta alpina LOMBARD,
Campbelliella striata (CAROZZI),
Clypeina jurassica FAVRE & RICHARD,
Macroporella praturloni DRAGASTAN,
Nipponophycus sp.,
Petrascula bursiformis ETALLON,
M. cf. embergeri BOUROULLEC &
DELOFFRE,
Salpingoporella annulata CAROZZI,
S. pygmaea (GÜMBEL),
Solenopora jurassica PFENDER.

4.2. Waschberg zone

This tectonic unit separates the Molasse zone in the west from the Vienna Basin in the east. The hilly character of the landscape is due to the Upper Jurassic thrust sheets of the Ernstbrunn Limestone and the Klentnice Formation, which are tectonically uplifted remains of the Autochthonous Jurassic. The latter is the equivalent of the Kurdejov Formation (= "Kalkarenitserie") and the Altenmarkt Formation in the underground. In fact, only a few details are known about the Klentnice Formation, because of lacking outcrops at the surface.

Fossils of the Tithonian Ernstbrunn Limestone (ZEISS & BACHMAYER, 1989; REHANEK 1987) are matter of scientific interest since more than 200 years (HAIDINGER 1775). More than 50 years ago dasyclad algae were described for the first time (BACHMAYER 1941). KAMPTNER (1951) described *Cayeuxia dörflesiana*, recently HOFMANN (1992) applied simple statistical methods on *Salpingoporella pygmaea* (GÜMBEL). Some detailed information is achieved by several works of ELIAS and ELIASOVA (e.g. ELIAS 1981, ELIAS 1992, ELIAS & ELIASOVA 1986).

**Algal flora of the Ernstbrunn Lime-
stone**
(ELIAS & ELIASOVA 1986, HOFMANN
1990, BACHMAYER 1941, KAMPTNER
1951, REHANEK 1987)

Bacinella irregularis RADOICIC,
Lithocodium sp.,
Thaumatoporella parvovesiculifera
RAINERI,
Koskinobullina socialis CHERCHI &
SCHROEDER,
Tubiphytes sp.,

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Marinella lugeoni PFENDER,
Acicularia sp.,
Girvanella sp.,
Solenopora sp.,
Arabicodium sp.,
Cayeuxia doerflesiana KAMPTNER,
C. piae FROLLO,
C. moldavica FROLLO,
Nipponophycus cf. *ramosus* YABE &
TOYAMA.
Griphoporella sp.,
G. ehrenbergi BACHMAYER,
Petrascula piae BACHMAYER,
P. bursiformis (BERNIER),
Macroporella praturloni DRAGASTAN,
Heteroporella sp.,
H. lemmensis BERNIER,
H. aff. lusitanica RAMALAHO,
H. morillonensis BERNIER,
Pseudocymopolia sp.,
Campbelliella striata (CAROZZI),
Neoteutloporella socialis
(PRATURLON),
Neoteutloporella sp.,
Teutloporella obsoleta CAROZZI,
Salpingoporella annulata CAROZZI,
S. katzeri CONRAD & RADOICIC,
S. pygmaea (GÜMBEL),
S. steinhauseri (CONRAD, PRATURLON
& RADOICIC),
Linoporella capriotica OPPENHEIM,
L. svilajensis SOKAC & VELIC,
L. aff. ellioti PRATURLON,
Triploporella sp.,
Clypeina solkani CONRAD & RADOI-
CIC,
C. jurassica FAVRE & RICHARD,
Actinoporella podolica ALTH,
Sarfatiella dubari CONRAD & PEY-
BERNES.

4.3. Prättigau

At the border between Austria and Switzerland, in the geographical area of the Rätikon at the top of the mountains

there are to be found some remains of a grey limestone called Sulzfluh Limestone. The most intensive study on algae of this Late Kimmeridgian to Middle Tithonian limestone which was originally deposited in an intrabasinal platform of the Penninic realm (pers. comm. OBERHAUSER) is the thesis of OTT (1969). Three different facies zones can be distinguished:
– micrite facies
– intrasparite facies (*Clypeina jurassica* FAVRE & RICHARD and *Campbelliella striata* (CAROZZI))
– oolith facies.

Algal flora from the Sulzfluh Limestone (OTT 1969)

Bacinella irregularis RADOICIC,
Cayeuxia austriaca FENNINGER,
C. doerflesiana KAMPTNER,
C. moldavica FROLLO,
C. americana JOHNSON,
C. kurdistanensis ELLIOTT,
C. piae FROLLO,
C. mediterranea HERAK,
Marinella lugeoni PFENDER,
Acicularia sp.,
Clypeina jurassica FAVRE & RICHARD,
Teutloporella aff. *obsoleta* CAROZZI,
Salpingoporella sp.,
S. mühlbergii (LORENZ),
Pseudoepimastopora jurassica
ENODO,
Campbelliella striata (CAROZZI).

4.4. Northern Calcareous Alps

The study of the Plassen Limestone, the Tressenstein Limestone and the Oberalm Beds started in the sixties by members of the University of Graz. The relation of these different formations within the carbonate platform is shown in fig. 4.





FENNINGER (1966) was the first to distinguish a micritic from a sparitic dominated Plassen Limestone facies. Some detailed microfacial analysis was made later by STEIGER & WURM (1980) who found a highly differentiated facies pattern (fig. 5).

Algal flora from the Plassen Limestone, Barmstein Limestone and Tressenstein Limestone

(FENNINGER 1966, FENNINGER & HOLZER 1972; , FENNINGER & HÖTZL 1967, FLÜGEL 1964; FLÜGEL & FENNINGER 1966, HÖTZL 1966; STEIGER & WURM 1980; MOSSBAUER 1989; MANDL & HOFMANN 1993)

Bacinella irregularis RADOICIC,
Lithocodium morikawai ENDO,
L. aggregatum ELLIOTT,
Pseudolithocodium sp.,
Tubiphytes sp.,
Thaumatoporella parvovesiculifera RAINERI,
Koskinobullina socialis CHERCHI & SCHRÖDER,
Muniera baconica (DEECKE),
Cayeuxia sp.,
C. austriaca FENNINGER,
C. anae DRAGASTAN,
C. moldavica FROLLO,
C. piae FROLLO,
Pycnoporidium lobatum YABE & TOYAMA,
Nipponophycus ramosus YABE & TOYAMA,
Arabicodium sp.,
Solenopora sp.,
S. cf. sudakensis MASLOV,
S. jurassica BROWN,
Actinoporella podoloica ALTH,
Macroporella sp.,
Pseudoclypeina sp.,
Salpingoporella johnsoni (DRAGASTAN).

ygmaea (GÜMBEL),
S. annulata CAROZZI,
S. grudii CONRAD & RADOICIC,
S. sellii (CRESCENTI),
Clypeina jurassica FAVRE & RICHARD,
C. solkani CONRAD & RADOICIC,
C. parvula CAROZZI,
Petrascula bursiformis ETALLON,
Nipponophycus sp.,
Pseudoepimastopora sp.,
Pseudoepimastopora jurassica ENDO,
Griphoporella sp.,
Macroporella praturloni DRAGASTAN,
Campbelliella striata (CAROZZI),
Heteroporella lemmensis BERNIER,
Teutloporella obsoleta CAROZZI,
Neoteutloporella socialis
(PRATURLON),
Cymopolia sp.

5. Paleobiogeographical conclusions

Starting in the last century when geologists recognized the different faunistic character of the "Alpine" Plassen Limestone and the "outer - alpine" Ernstbrunn Limestone, this challenge is now up to the algae. Comparing the algae listed above, it is evident, that most of them are common in the Alpine, the Penninic and in the Outer Alpine realm. Many of them are even cosmopolitan like *Salpingoporella annulata* CAROZZI. As an exception, up till now only *Triploporella* sp. and *Linoporella* sp. are known only from the Ernstbrunn Limestone, so they can be designed as "outer - alpine" genera.



6. Acknowledgements

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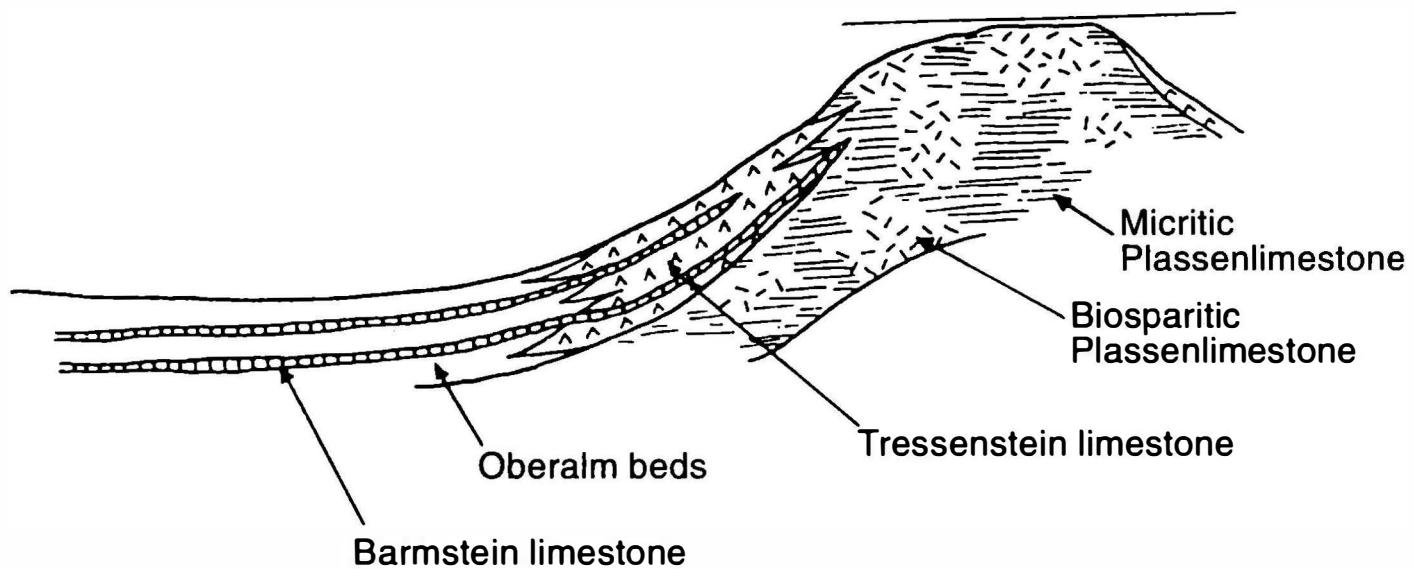
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H. Reuss has to be thanked for his permission to carry out works in the animal-park of Ernstbrunn.

Fig. 4: Geological situation of the Malmian carbonates in the Northern Calcareous Alps (from FLÜGEL & FENNINGER 1966).

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Facies model of the Upper Jurassic Plassen limestones

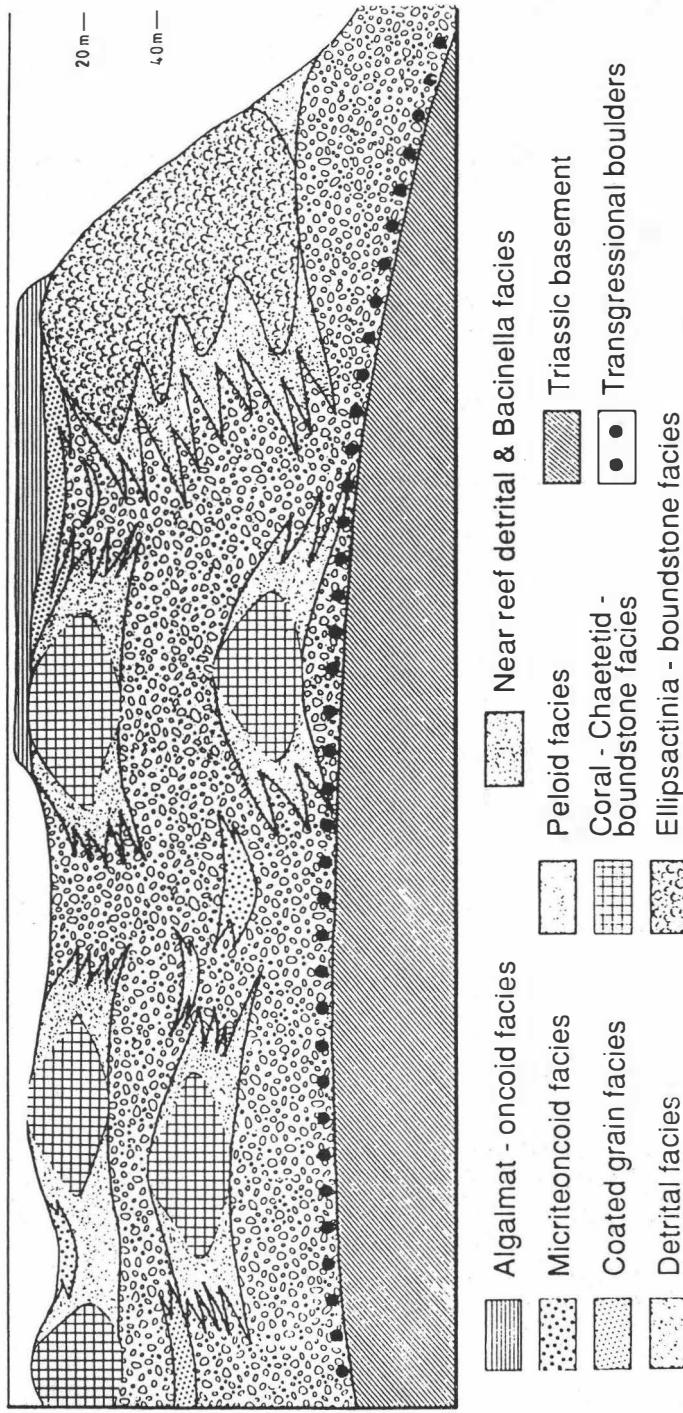


Fig. 5: Facies model of the Late Jurassic Limestones (from STEIGER & WURM 1980).



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