

Two pteridophyll remains, both about 3 cm long, are exposed at the surface. One specimen represents the terminal part of a last order pinna with a rachis 0.8 mm wide. The pinna consists of 5 nearly opposite pairs of pinnules which pass into a terminal part with indistinct lobes. The venation pattern is poorly preserved. The leaf is best determined as cf. *Callipteridium* pteridium, a genus of Permo-Carboniferous pteridosperms but could also represent a fragment of *Alethopteris*. Independent of any generic determination, the maximum stratigraphic age is Carboniferous; *Callipteridium* pteridium is typical for Stephanian with rare occurrences in the Lower Permian, *Alethopteris* occur in the Upper Carboniferous and Lower Permian. The stratigraphic age is therefore in the same range as the intrusion ages determined for the Zentralgneis in this area.

## **THE TWO PHASE REGION BETWEEN ORTHO- AND CLINOZOISITE**

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Analyses of natural coexisting clino- and ortho zoisite can give some constraints about the position of the two phase region in dependence of P and T. We analysed epidote minerals from rocks with different P-T-history (Dorfer Tal and Frosnitzal, Tauern: 20 kbar/600 °C with retrograde equilibration; Weissenstein, Münchberger Gneismasse: < 10 kbar/620 °C with retrograde equilibration; Schwingen, Münchberger Gneismasse: < 4 kbar, 400 °C). For the high P-T-rocks we could confirm a transition loop from - 15 to - 30 mole %  $Al_2Fe$  consistent with data by ACKERMAN & RAASE (1973), but significantly smaller than that given by ENAMI & BANNO (1980) and the experimental data by PRUNIER & HEWITT (1985). Retrograde reequilibration at low P and T yields epidote compositions inconsistent with the transition loop. Also, minerals from low P-T-rocks span the whole range between 5 and 30 mole %  $Al_2Fe$  of the proposed transition loop. We conclude that in addition to temperature, pressure plays a very important role; variations of mineral assemblages can be explained in a hypothetical P-T-x diagram.

Epidote minerals from the above mentioned localities show a complex growth zoning pattern which is only slightly modified by diffusion. This indicates that they may be potentially useful as petrogenetic indicators, but analyses have to be carried out with a back-scattered-electron image system in order to obtain interpretable results.

- ACKERMAN, D. & RAASE, P. (1973): Coexisting Zoisite and Clinozoisite in Biotite Schist from the Hohe Tauern, Austria. *Contr.Min.Petr.* 42, 333-341.
- ENAMI, M. & BANNO, S. (1980): Zoisite-clinozoisite relations in low- to medium-grade high-pressure metamorphic rocks and their implications. *Min.Mag.* 43, 1005-1013.
- PRUNIER, A.R. & HEWITT, D.A. (1985): Experimental observations on coexisting zoisite-clinozoisite. *Amer.Mineral.* 70, 375-378.